

**HEAT VARIATIONS CAUSED BY GROUNDWATER FLOW
IN GROWTH FAULTS OF THE SOUTH TEXAS
GULF COAST BASIN**

APPROVED:

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IN GROWTH FAULTS OF THE SOUTH TEXAS
GULF COAST BASIN**

by

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THESIS

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Abstract

Zones of above average subsurface temperatures have been noted in certain areas of the Gulf Coast basin. Their cause has often been credited to geopressure, which presumably traps heat because of higher porosity and consequently lower thermal conductivity. I determined the temperature distribution in a portion of South Texas by collecting and analyzing over 1600 bottom-hole temperature measurements. The analysis included correcting the temperatures with the Kehle correction scheme; constructing isothermal surfaces by both interpolating and extrapolating the data, and Kriging the result. Temperature profiles were plotted for twelve subregions of the study area. The greatest temperature anomalies are associated with the Tertiary Wilcox growth fault zone, and the simple presence of geopressure is insufficient to account for the temperature anomaly.

Numerical modeling indicates that growth faults act as zones for concentrated vertical flow. The upwelling of deep basinal fluids advects heat and causes the high temperatures observed in the growth fault zone. The model, by Smith (1983), is two-dimensional, finite element, steady state, and couples heat and fluid transport. It indicates that the source of these fluids is deep, perhaps over 20,000 feet (6096 m) below the surface. The modeling results also suggest that an unidentified region of high thermal and hydraulic conductivity could exist coastward of the Wilcox faults at a depth of about 15,000-20,000 feet (4572-6096m).

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INTRODUCTION

In the Gulf Coast Basin, as well as in most sedimentary basins, temperatures are unevenly distributed. The source of heat is the earth itself, which radiates heat from within in a relatively uniform fashion as a result of dispersed radioactive decay. There are many factors which cause uneven thermal patterns to emerge from this uniform heat source. These can be grouped into the following broad categories: heat transport by basinal fluids, heat transport through the solids, the presence of plutons or other heat sources or sinks, chemical reactions, and mass and energy exchanges across basinal boundaries. Of these, conduction and transport of heat via the movement of ground water, which saturates the sediments of the basin, are perhaps the most important processes in the Gulf Coast. Water has a high specific heat (ie., it can absorb heat without greatly raising its temperature). Thus flowing ground water has the potential to transport a significant amount of heat. Conversely, the heat content itself influences groundwater movement by altering the fluid's hydraulic properties. Thus one process influences and is in turn

influenced by the other. These are said to be coupled. This thesis explores the importance of coupled ground water and heat transport in controlling the deep basinal thermal regime of the Gulf Coast Basin.

Knowledge of groundwater movement in the deep basin has practical consequences in both the energy and mineral industries, as well as in academic inquiries into basinal diagenesis. It is, however, often difficult to observe groundwater movement, especially where rates are slow and/or in the remote deep-basin. It is less difficult to measure temperatures at such points. The coupled flow phenomena provides a method by which groundwater movement in the deep basin can be inferred from available temperature data. This is the approach of this research.

The research required two distinct stages. The initial step was data collection and analysis. These are presented in Part I. Part I includes discussion of the data, their collection and preparation, descriptions of the thermal patterns that emerged from analysis of these data, and implications of these patterns for groundwater flow. After the completion of the first stage, it was evident that too many factors were involved in the thermal regime of the basin to make any definitive conclusions possible from this largely empirical approach. However, the data were consistent and anomalous thermal trends were evident. Therefore, the second stage required computer modeling of the basin using the acquired thermal data. The details of the model, its results, and its implications are covered in Part II. Finally, the analysis allows us to make some definite conclusions about heat flux and groundwater flow in the deep portions of the Gulf of Mexico Basin and to offer

suggestions for further investigation.

PART I

1.1 BACKGROUND

1.1.1 THE STUDY AREA

The study was conducted on a regional scale, covering a large portion of South Texas (Fig 1.1). The total area is over 17,000 square miles (44,030 km²) and includes all or part of the counties of Webb, Zapata, Jim Hogg, Kleberg, Duval, Jim Wells, Nueces, Lasalle, McMullen, Live Oak, Refugio, San Patricio, Aransas, Bee, Atascosa, Karnes, Goliad, and Victoria. The area features low topographic relief, with greatest elevations of about 600 feet (183 m) to the west, dropping gradually to sea level at the coast. The maximum depth investigated is 20,000 feet (6096 m), which is wholly within Tertiary sediments.

1.1.2 REGIONAL GEOLOGY

The onshore lower Tertiary Gulf Coast section is composed of several terrigenous clastic wedges, which thicken down dip toward the Gulf of Mexico. The sediments prograded over and basinward of the rapidly subsiding Cretaceous Stuart City shelf margin, where

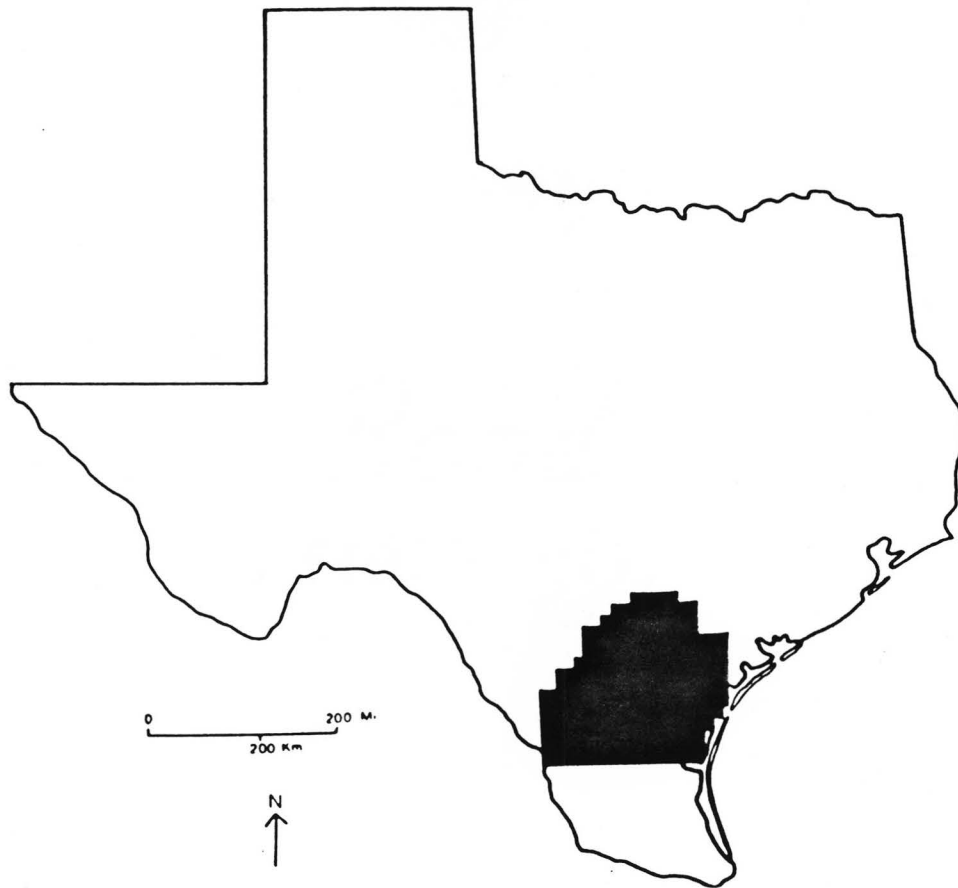


Figure 1.1: State of Texas with study area indicated in black.

contemporaneous growth faulting ensued in bands along the shelf edge. As sedimentation and subsidence continued, the shelf edge prograded seaward, establishing other growth fault zones (Fig 1.2). The major growth fault zones have been named after large sediment packages associated with them. The three major ones are the Wilcox, Vicksburg and Frio growth fault trends as shown on the cross section in Fig 1.2 which also identifies other major clastic groups.

1.1.3 THERMAL HISTORY

Potentially dominant heat sources are recent tectonic events, which can cause high basement heat fluxes, and volcanic activity, which causes localized heat-source variations.

The most recent tectonic event involving the Gulf Coast was the late Triassic/early Jurassic rifting episode. This event probably featured typical rift-level heat fluxes. Figure 1.3 (Royden and others, 1980) indicates a return to background heat-flux levels by Cretaceous time. Therefore, "normal" levels of heat flux have presumably prevailed since the Cretaceous. While local Cretaceous and Tertiary volcanic activity may have perturbed this pattern, their effects were minimal.

1.1.4 GENERALIZED FLOW SYSTEMS

Generalized groundwater flow through Gulf Coast sediments can be conceptualized as occurring in three distinct zones closely correlated with depth. Flow in each of these zones is controlled by different mechanisms. The three, from shallowest to deepest, are the

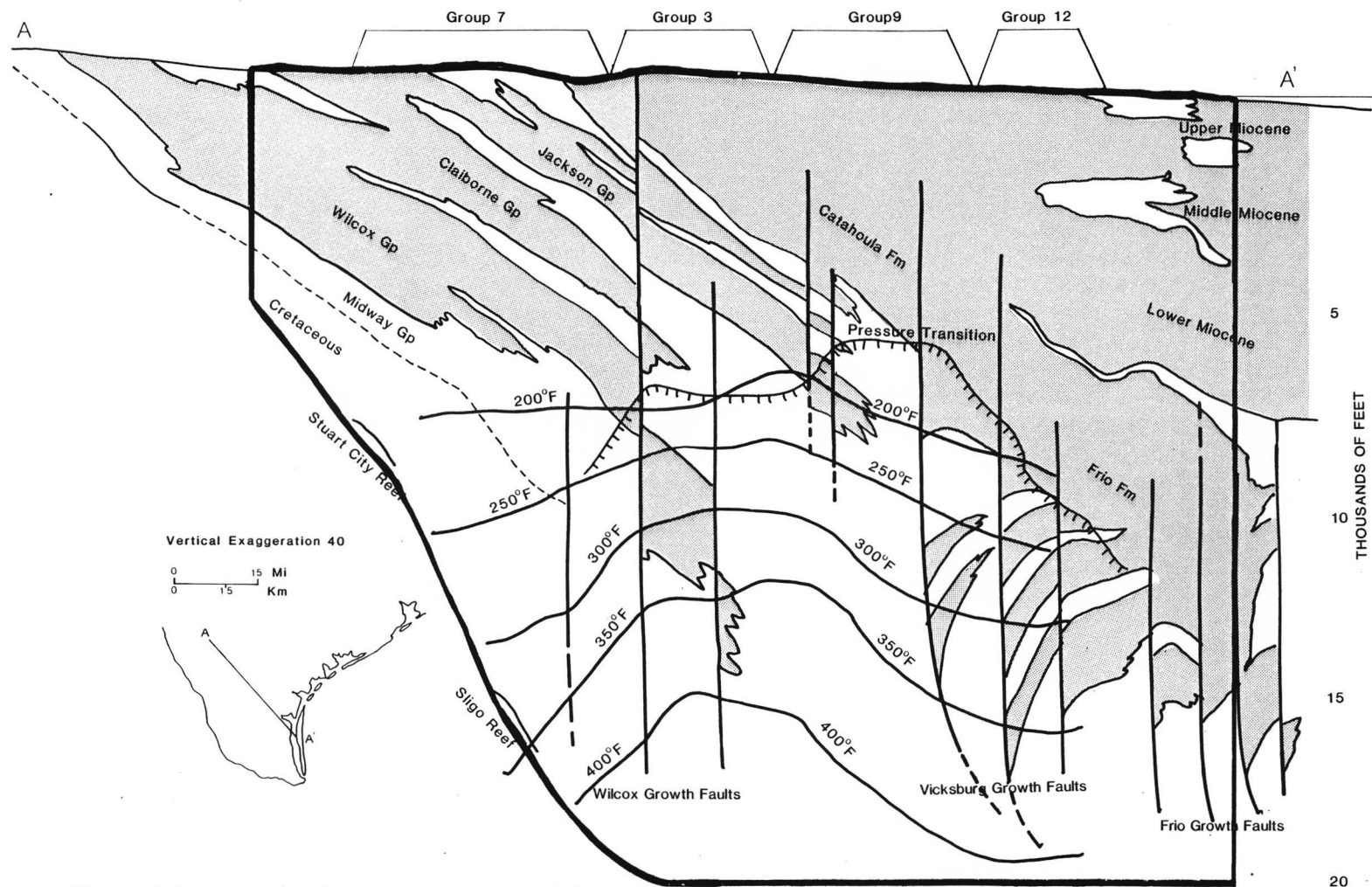


Figure 1.2: Generalized cross section of south Texas Gulf Coast. Grey pattern shows the sandier units, the rest of the section being predominantly argillaceous (modified from W.,E. Galloway, written communication). Also indicated are isotherms and the top of the geopressure Transition zone as determined by drilling mud weights. Heavy outline indicates position of finite element grid used in modeling (Part 2).

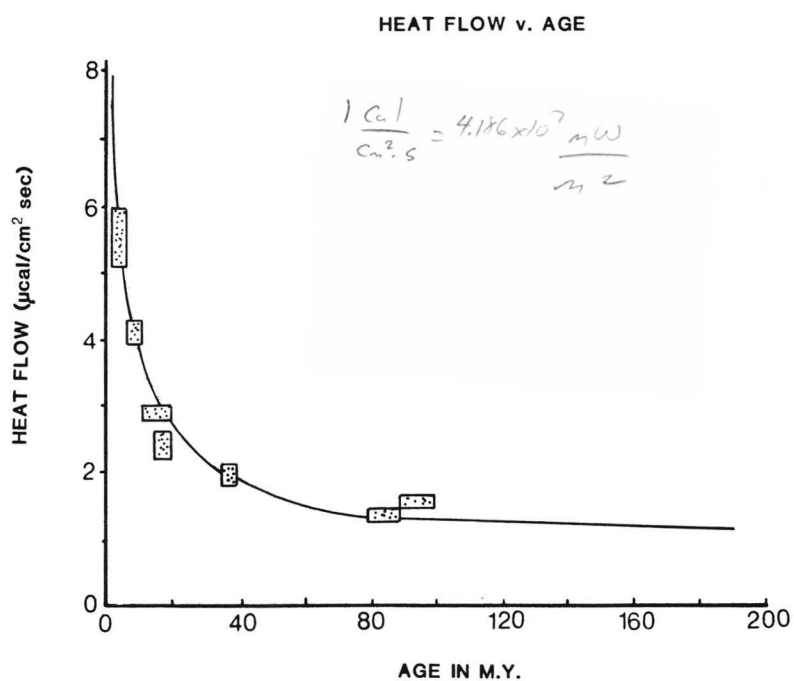


Figure 1.3: Generalized plot of mean heat flow versus time after a rifting event. Stippled boxes indicate various heat flow data (redrawn from Royden and others., 1980)

meteoric, compactional/overpressured and thermobaric zones. These are shown for a generalized sedimentary basin in Figure 1.4 from Galloway and Hobday (1983).

Meteoric Zone. Groundwater flow in the meteoric zone is controlled by the elevation of the water table (i.e., ground water moves from areas of high hydraulic potential, usually under topographic highs, to areas where the potential is low, usually at rivers and streams). It has been demonstrated by Back (1966) for the Atlantic Coastal Plain and also by Smith and others (1982) for the Oakville aquifer of Texas that most ground water discharges to rivers. This is contrary to the common misconception that meteoric water flows only down dip towards the coast, a condition that occurs primarily near the coast where there is little topographic relief and in deeper, more inactive portions of the meteoric zone. For these cases, discharge is by diffuse, upward, cross-formational flow in the vicinity of the shoreline and is not concentrated near rivers and streams. The term *meteoric* implies water of an atmospheric origin, but connate water may also enter the meteoric system.

Compactional/overpressured zone. The compactional/overpressured system lies below the meteoric one. Here groundwater movements are not significantly influenced by surficial conditions, but rather are controlled by excess pore-fluid pressure gradients. Excess pressures are partially caused by the weight of the accumulated overlying sediments which causes this zone to compress or compact. As the sediment grains come closer together, the intergranular

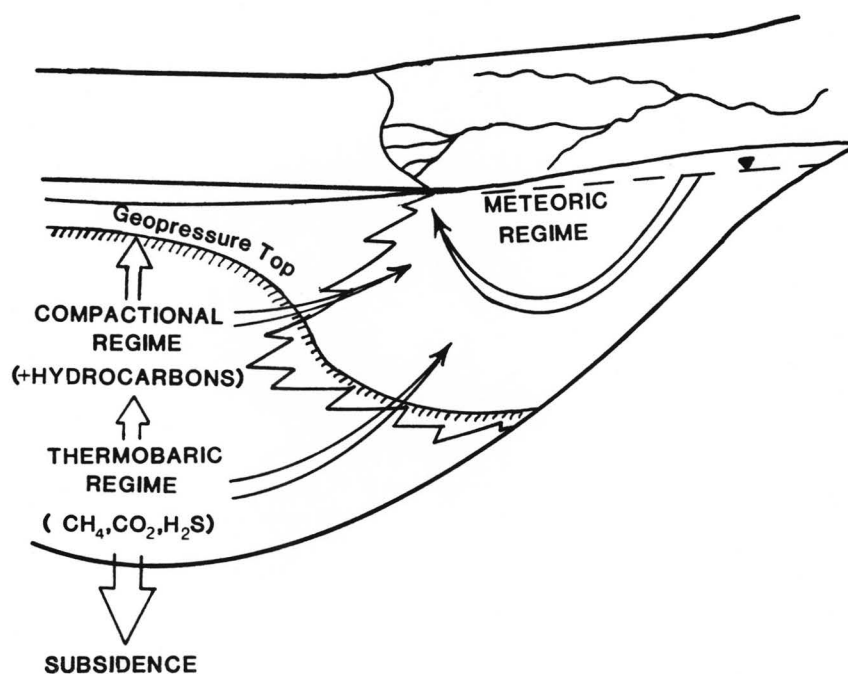


Figure 1.4: Generalized hydrogeologic model of a sedimentary basin (redrawn from Galloway and Hobday, 1983).

spaces become smaller and the pore fluid is expelled concomitantly. If permeability is not sufficiently high, as is the case in Gulf Coast clays and shales, the fluid becomes trapped for a time. Higher porosities may thus be maintained in the sediment, kept open by trapped pore fluids which are also subjected to greater than normal pressures. This circumstance is variously known as *overpressuring*, *excess pressuring*, or *geopressuring* and represents a compactional disequilibrium. Eventually, the trapped fluids will bleed off, the sediments will compact, and equilibrium will be reestablished, unless the pressuring processes are continued (Gibson, 1956; Magara, 1976; Sharp, 1976; Sharp and Domenico, 1976; and Keith and Rimstidt, 1985). Other mechanisms have been proposed as contributing to excess pressuring, including aquathermal pressuring (Barker, 1972), mineral phase transformations such as clay dehydration (Burst, 1969; Weaver & Beck, 1971; Bruce, 1984), and hydrocarbon maturation (Hedberg, 1980).

Much of the Gulf Coast Basin is overpressured. The hachured line in Figure 1.2 indicates the top of overpressure derived from mud weights. Figure 1.5 shows generalized plots of fluid pressure versus depth for the Gulf Coast, depicting inflection points where pressure deviates from the normal hydrostatic condition. The sediments beneath these points are considered to be overpressured, although discontinuous zones of lesser overpressures may exist above this depth.

Thermobaric Zone. The thermobaric regime is transitional with the compactional regime and encompasses the deepest parts of the basin. This regime is defined as the zone where the temperatures and pressures

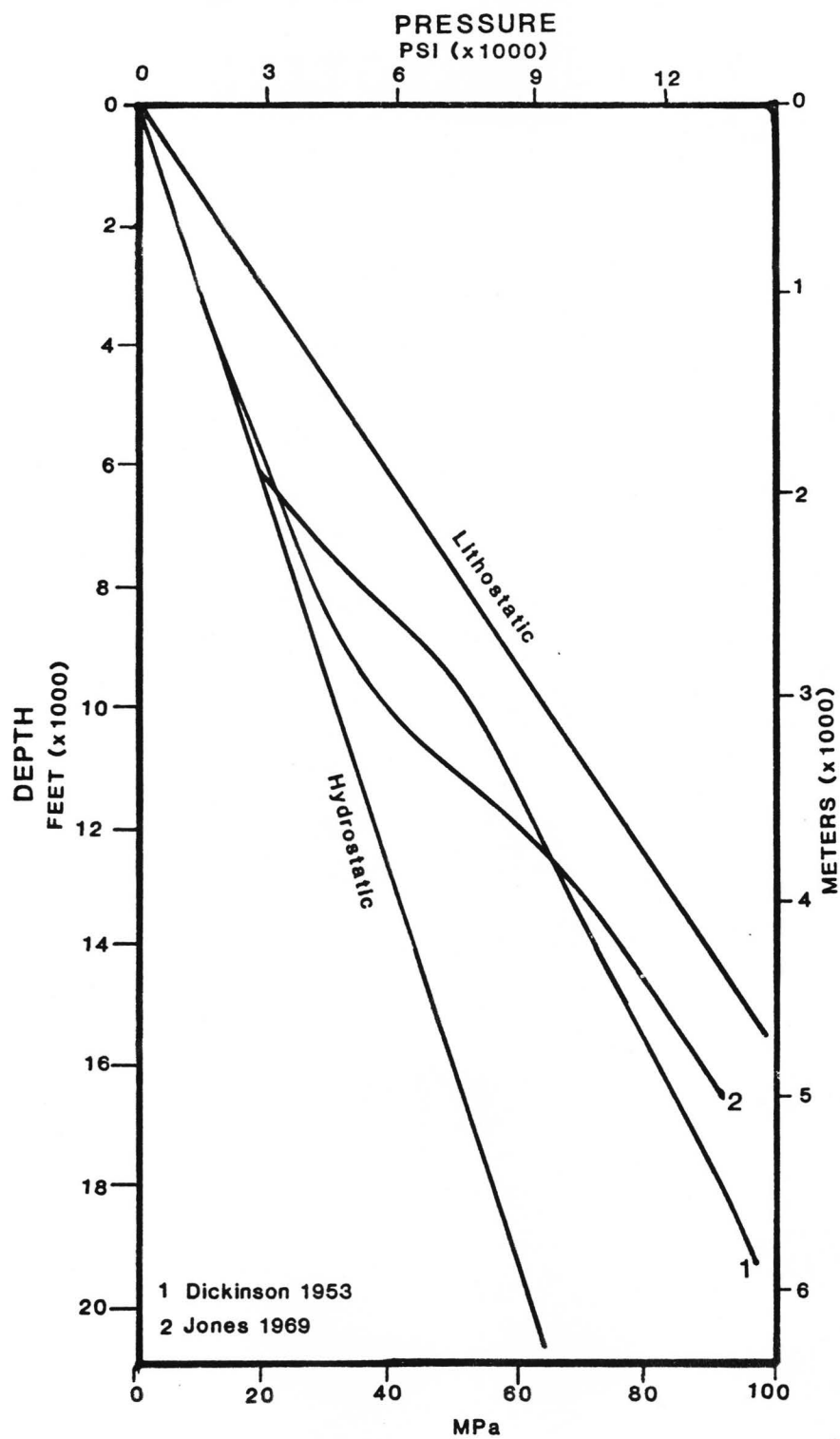


Figure 1.5: Generalized variation of fluid pressure versus depth in the Gulf Coast.

are great enough to cause low-grade metamorphism. The subsequent release of metamorphic fluids is presumed to be the major hydrologic process. Rocks above and/or below the Louann Salt may currently be producing water in addition to large quantities of carbon dioxide as the result of metamorphic reactions. The importance of these fluids on the hydrology of the rest of the basin is unknown because direct sampling of this portion of the basin has not yet been accomplished.

1.2 DATA COLLECTION AND PREPARATION

The data collected for this research consists of over 1600 bottom-hole temperatures (BHTs) from over 725 separate wells (see pocket map for locations). The BHT readings were taken off of well headers, i.e., the first page of an electric log. The headers also list other information about the well including the Company that drilled it, the depths of each run, the land owner, sometimes the elevation of the well head, and its location. The logs listed in Appendix A are all from the files of the Bureau of Economic Geology (BEG), but are in the public domain and may be purchased from the appropriate well libraries. To collect them, all logs held by BEG for wells within the study area were examined. Of those, all that were readable, that included the well head elevation, and that had BHTs above 170°F were recorded. It should be noted that data for shallow measurements below 170°F were not recorded, though such measurements did exist on many logs. It was later apparent that this information would have been quite useful, but time

constraints ruled out their acquisition. Any additional research should include adding shallow measurements to the data base.

1.2.1 TEMPERATURE CORRECTIONS

A bottom-hole temperature is a temperature reading taken at the bottom of a well. In the case of deep oil wells, such as those used in this study, a logging tool equipped with a thermometer is lowered periodically as the hole is deepened. Thus several runs, each with a separate BHT reading, are commonly available from a deep well.

When a well is logged, the bottom-hole temperature is just one of many different types of measurements taken. Many different measuring instruments are usually lowered into the hole strung together in a pipe-like bundle known as a *sonde*. When logging, drilling operations must be halted, and the sondes are lowered into the hole without removing the drilling mud. There are many sources of error that are introduced when taking the bottom-hole temperature, mostly linked to temperature equilibration problems. When drilling, the drilling mud which fills the hole and surrounds the drill stem is circulated to remove the drilling debris and to keep the drill bit cool. As a consequence, the rocks surrounding the hole are also cooled. When logging, mud circulation is halted, but the time elapsed between halted circulation and temperature measurement is usually not enough for the bottom-hole temperature to reequilibrate with the actual temperature of the rock at that depth. Of course, it would be useful to record this elapsed time so that some estimate of the degree of equilibration could be made, but

unfortunately this is seldom included on the well log. It is clear, however, that some sort of correction is needed for the BHT reading.

The correction scheme adopted for these data is that of Kehle (1971), an empirical correction determined by comparing equilibrium and well-log bottom-hole temperatures from 602 selected West Texas and Louisiana wells. The apparent temperature from well-log BHTs was subtracted from the temperature obtained using equilibrium BHTs and the difference was plotted as a function of depth. A polynomial was fitted to these data using a least square error criterion, resulting in the correction curve shown in Appendix B.

1.2.2 CREATING THE COMPUTER FILE

In order to use the data in computer applications, a computer data file was created. The file was first loaded with well identification numbers and corresponding location coordinates. This was done using a digitizing tablet and BEG software. The well maps were placed on the digitizing tablet, and the well positions were recorded by manually positioning a cursor over the selected wells. An arbitrary numbering system was used to keep track of each well, and a notebook was used to record all the corresponding log information. When all the wells were digitized, the elevation, temperature, and depth readings recorded in the notebook were free-formatted into the file from a terminal, one computer line per well. The completed file contained over 725 lines, each line containing one well's identification number, location coordinates, multiple BHT and depth readings, and its elevation.

1.2.3 MANIPULATING THE FILE

Once the raw data was entered into the file, a program was written to calculate the Kehle temperature corrections for each paired BHT and depth reading. The program created a new file which contained the corrected temperatures in a formatted form. Appendix B shows the correction scheme.

In order to produce isothermal surfaces, the depth data had to be converted to a depth relative to sea-level, and interpolated or extrapolated to arbitrarily chosen temperatures. I choose 50°F increments from 200°F to 400°F. A program was written to produce files for each isothermal surface from the main data file. Since the thermal trends which were being investigated were assumed to vary with depth, it was essential to limit the amount of interpolation or extrapolation to a fairly narrow temperature range. The program thus requires the user to specify a temperature range limit. If a well has no BHT values within the specified temperature range of that isotherm, then that well is disregarded. If there is a value within the range, then the depth to the isotherm is either interpolated between it and the next closest temperature/depth pair that the isotherm value straddles, or, if there are no other values, the data is extrapolated to the isotherm. The range was set at 30°F for the 200, 250, and 300°F isotherms, and was 50°F for the 350 and 400°F isotherms, for which data are more scarce. Appendix C shows the FORTRAN program used to create the isothermal data files.

Data scatter had to be eliminated to contour isotherms. This was accomplished by the geostatistical technique of kriging. The kriging process smooths the data by analyzing the variation in the data and weighting each data point accordingly. The technique also provides a determination of the accuracy of the contoured surface by supplying standard deviations at gridded intervals. Appendix D provides extensive information on the procedures for kriging.

Finally, to produce contour maps and three-dimensional drawings, Radian Corporation's CPS1 graphics subroutines were employed. Appendix E briefly covers the use of CPS1, which is sophisticated and cumbersome. Large amounts of computer and personal time were consumed before the figures presented in this chapter were successfully produced. The temperature profiles and linear regressions were executed in the statistics package SPSS.

1.3 OBSERVED THERMAL PATTERNS

The following series of twenty illustrations (Figs. 1.6-1.10, a through d) depicts the temperature patterns that were resolved from the data. These patterns are viewed from the perspective of five isotherms (200, 250, 300, 350, 400 °F). Four illustrations are shown for each isotherm: a, a three-dimensional representation of the surface; b, the corresponding contour map; c, the selected data; and d, a contour map of the standard deviations for the isothermal surface. With each succeeding temperature increment the depths increase, as one would expect. Also increasing with temperature is the prominence of a thermal

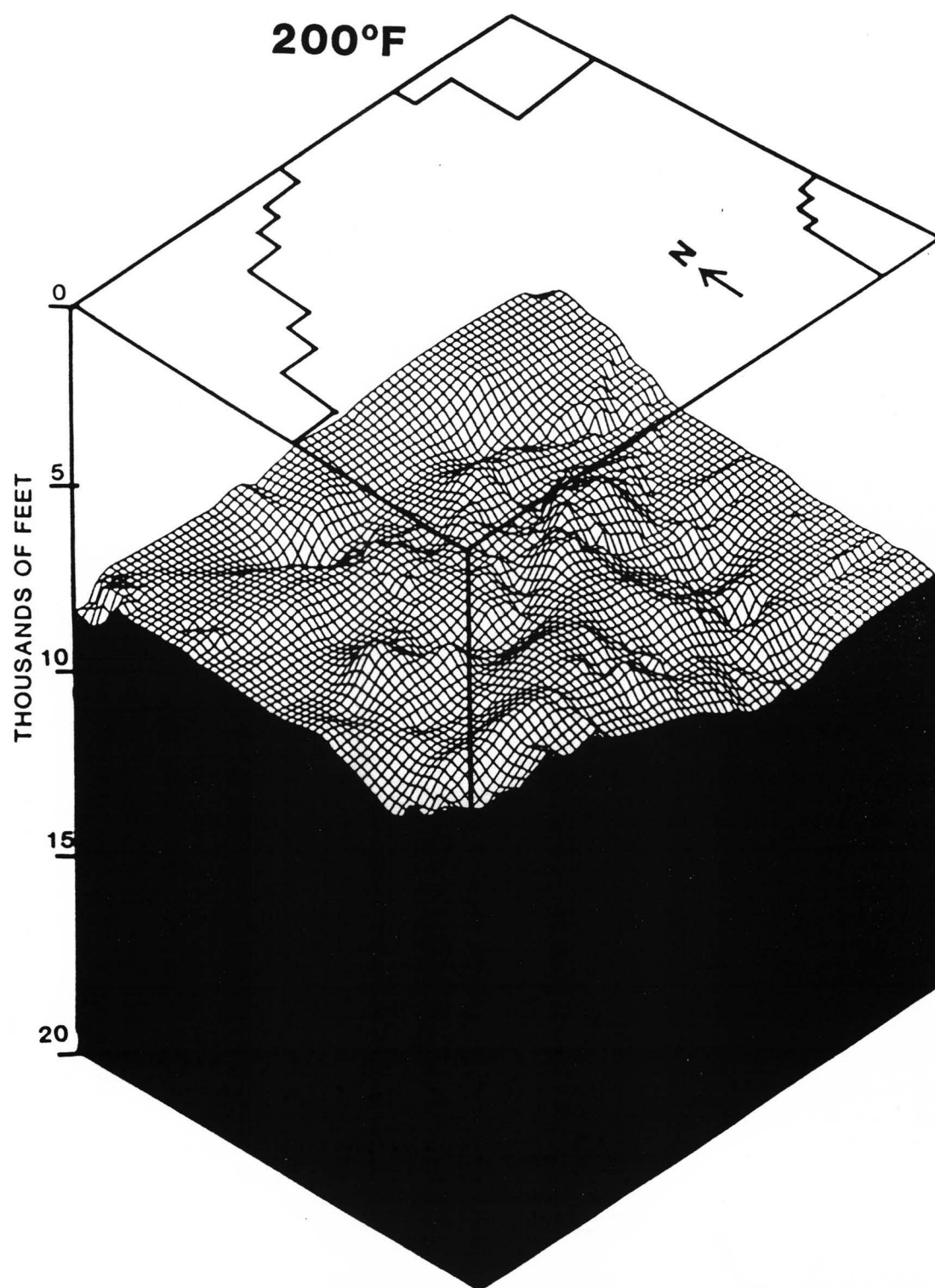


Figure 1.6a: Isotherm from corrected and Kriged BHT data with study area projected above.

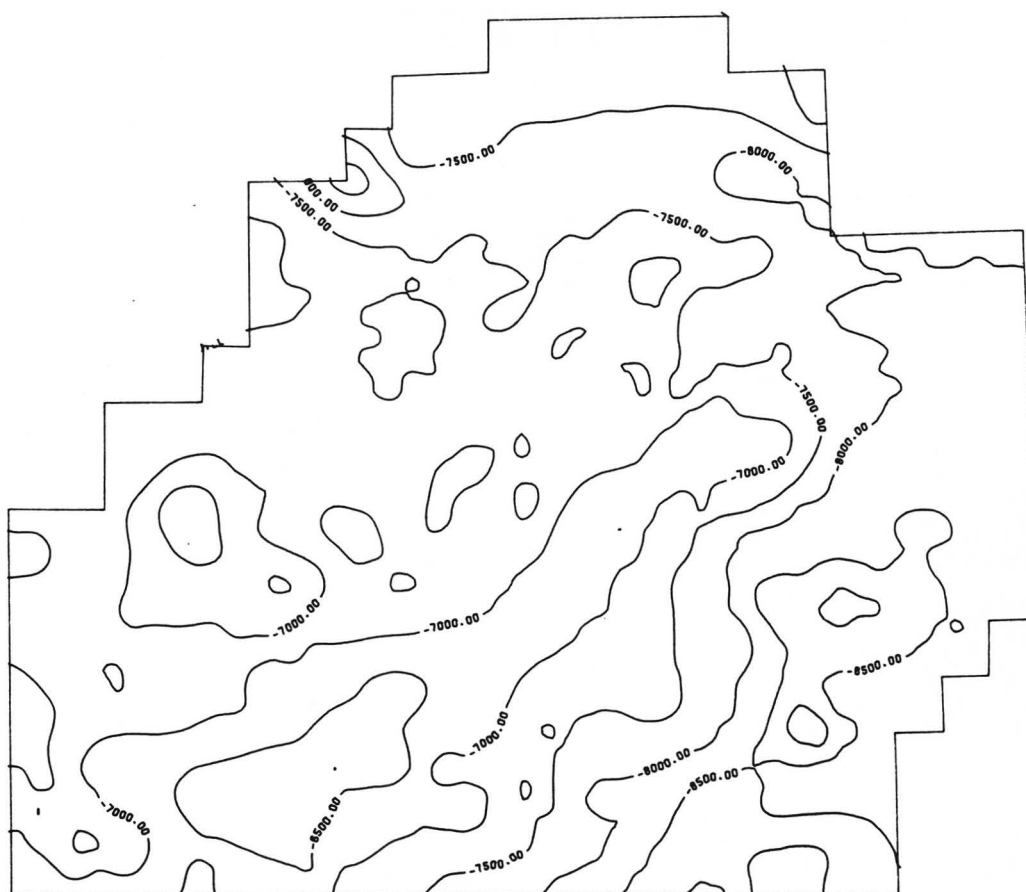


Figure 1.6b: Contoured 200°F isotherm within the study area.

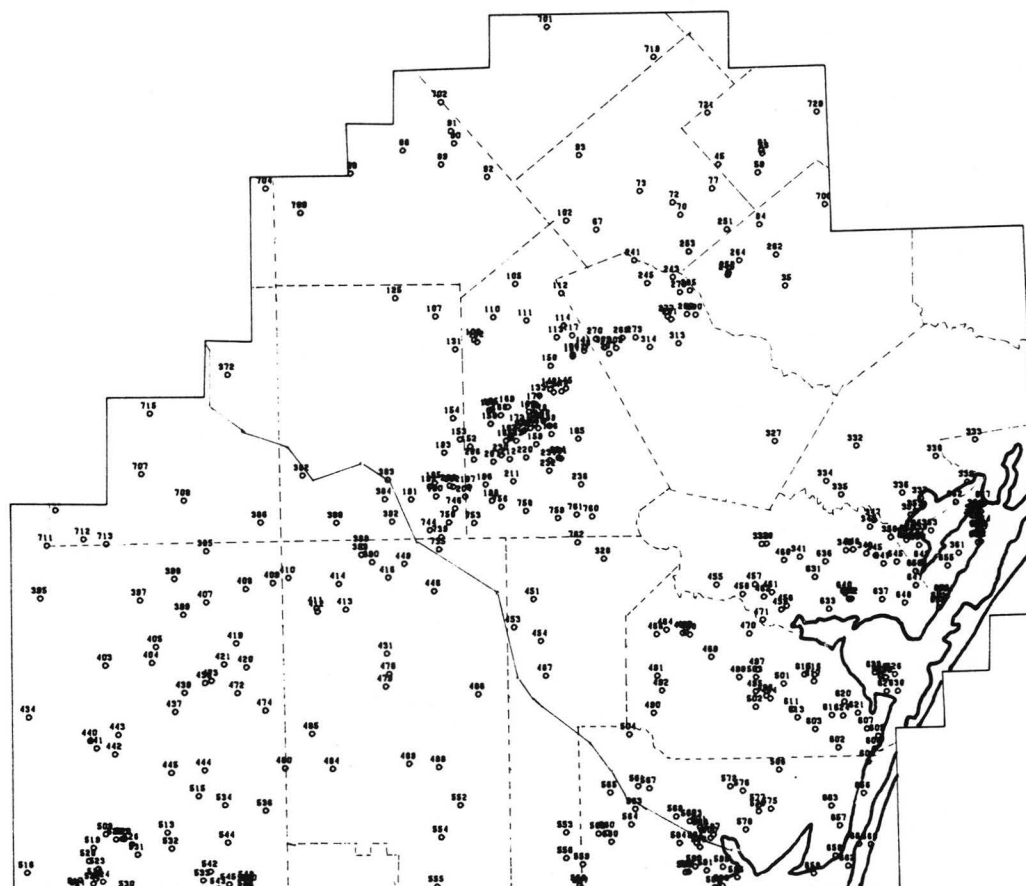


Figure 1.6c: Well locations for the 200°F isotherm.

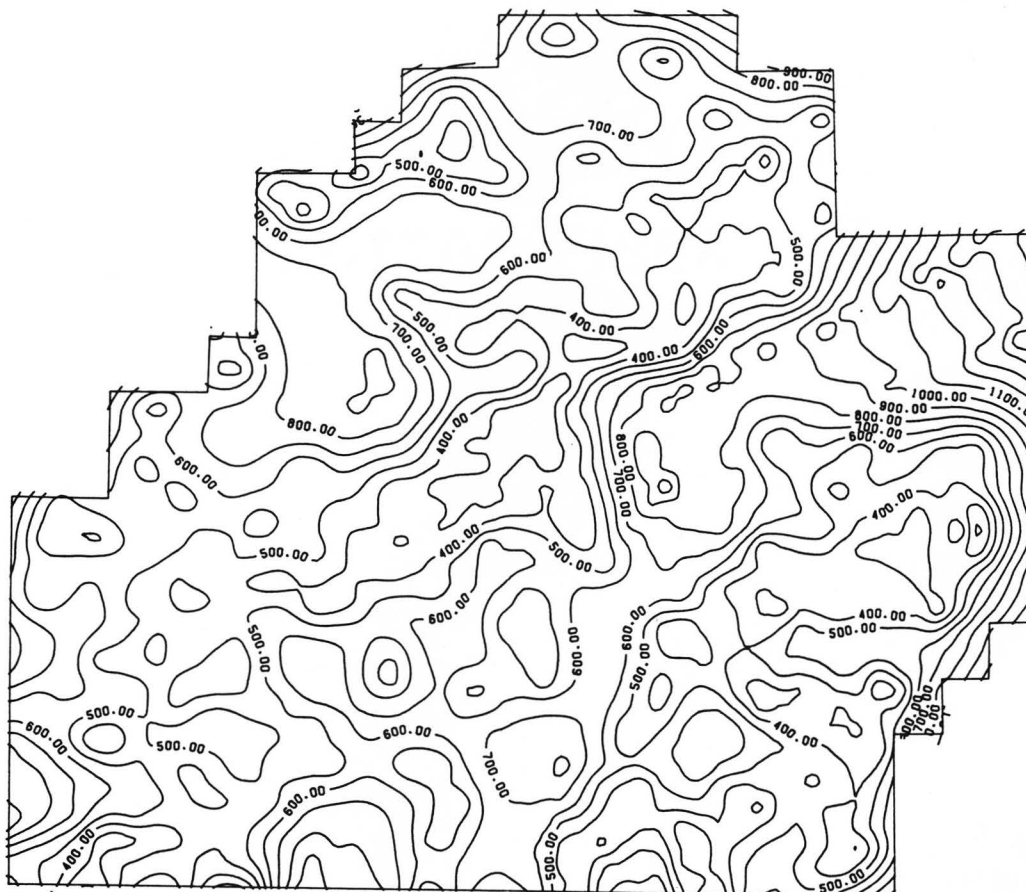


Figure 1.6d: Standard Deviation of contoured 200°F isotherm (computed by Kriging).

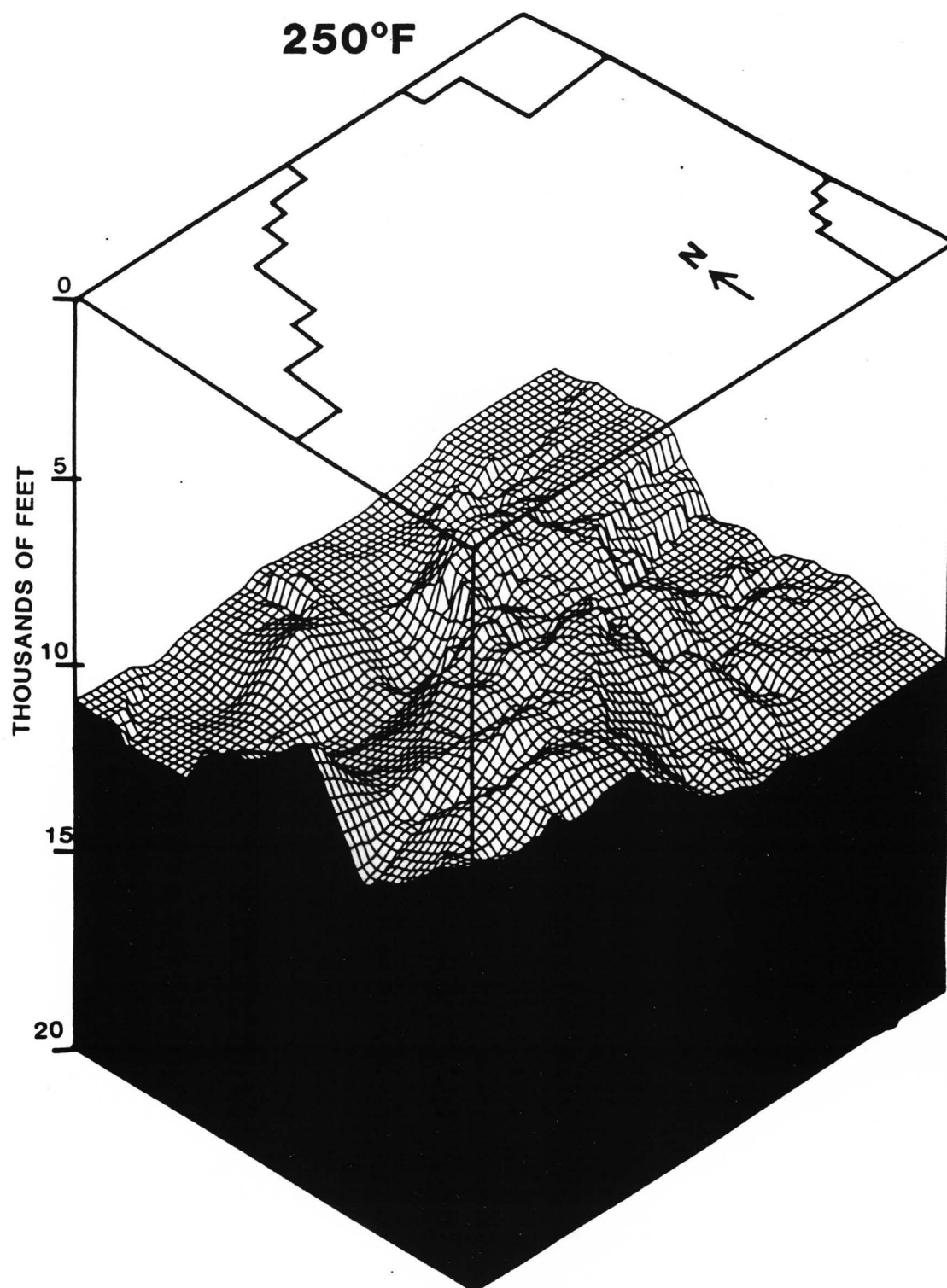


Figure 1.7a: Isotherm from corrected and Kriged BHT data with study area projected above.

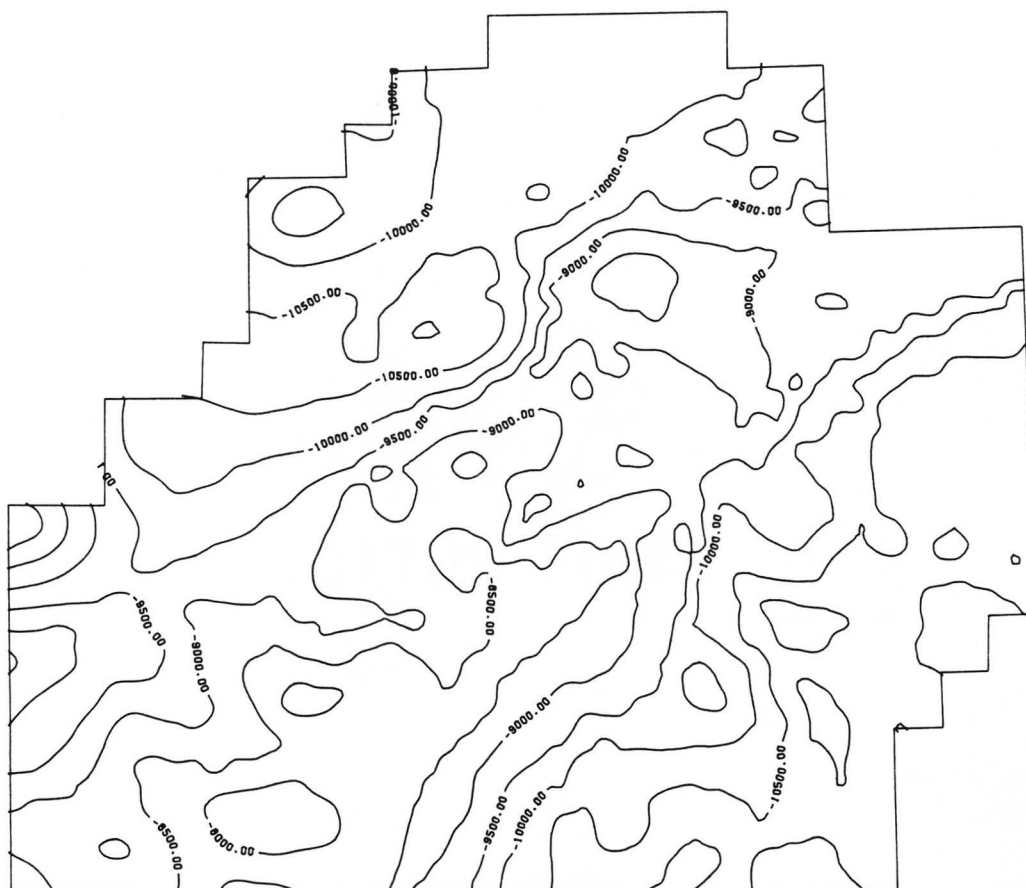


Figure 1.7b: Contoured 250°F isotherm within the study area.

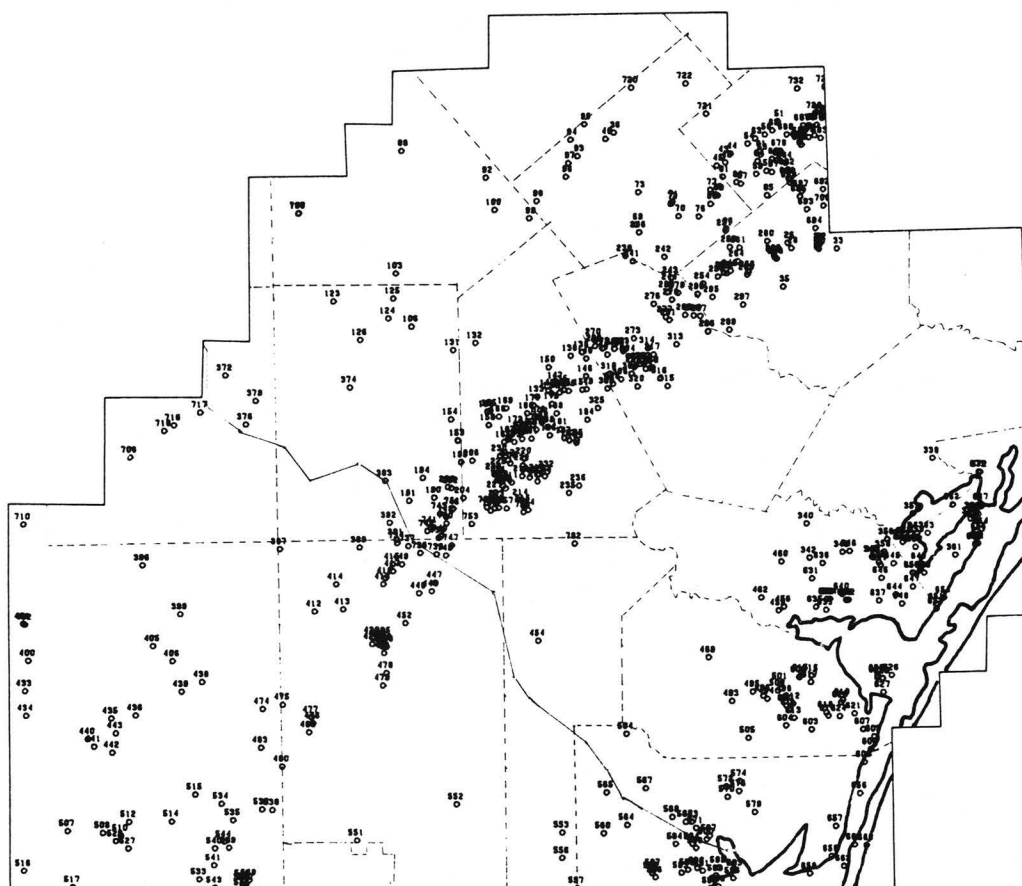


Figure 1.7c: Well locations for the 250°F isotherm.

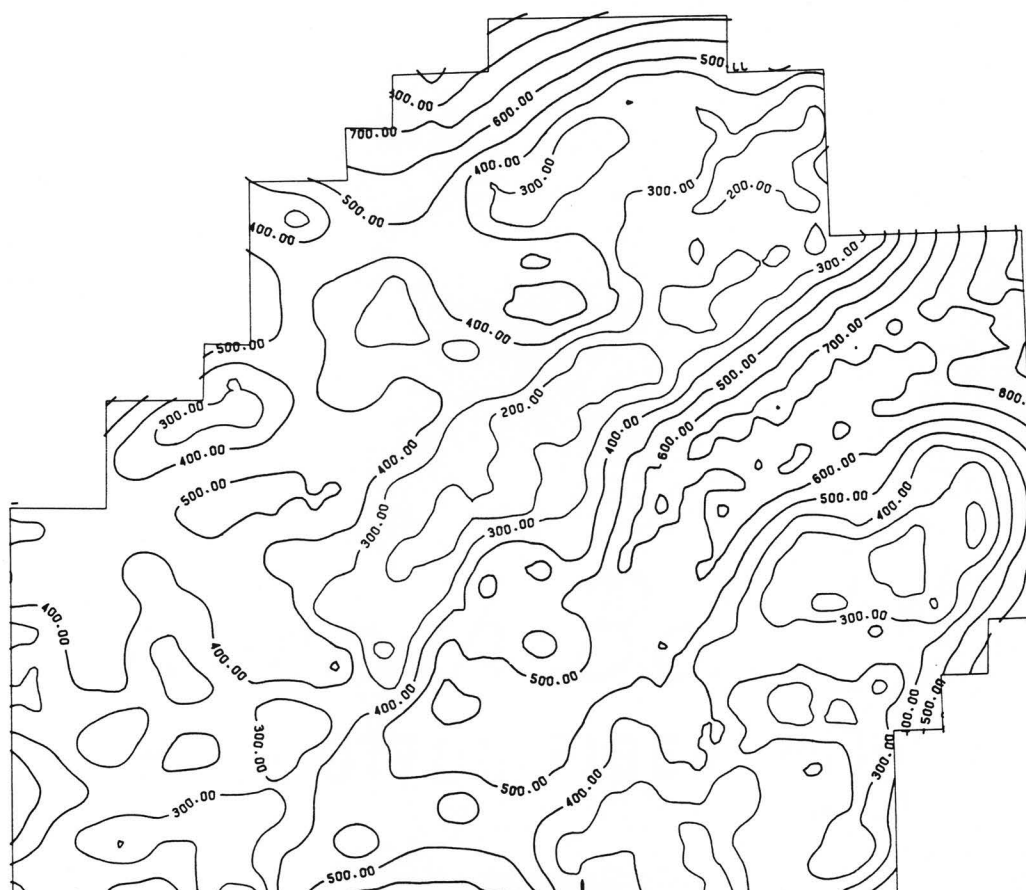


Figure 1.7d: Standard Deviation of contoured 250°F isotherm (computed by Kriging).

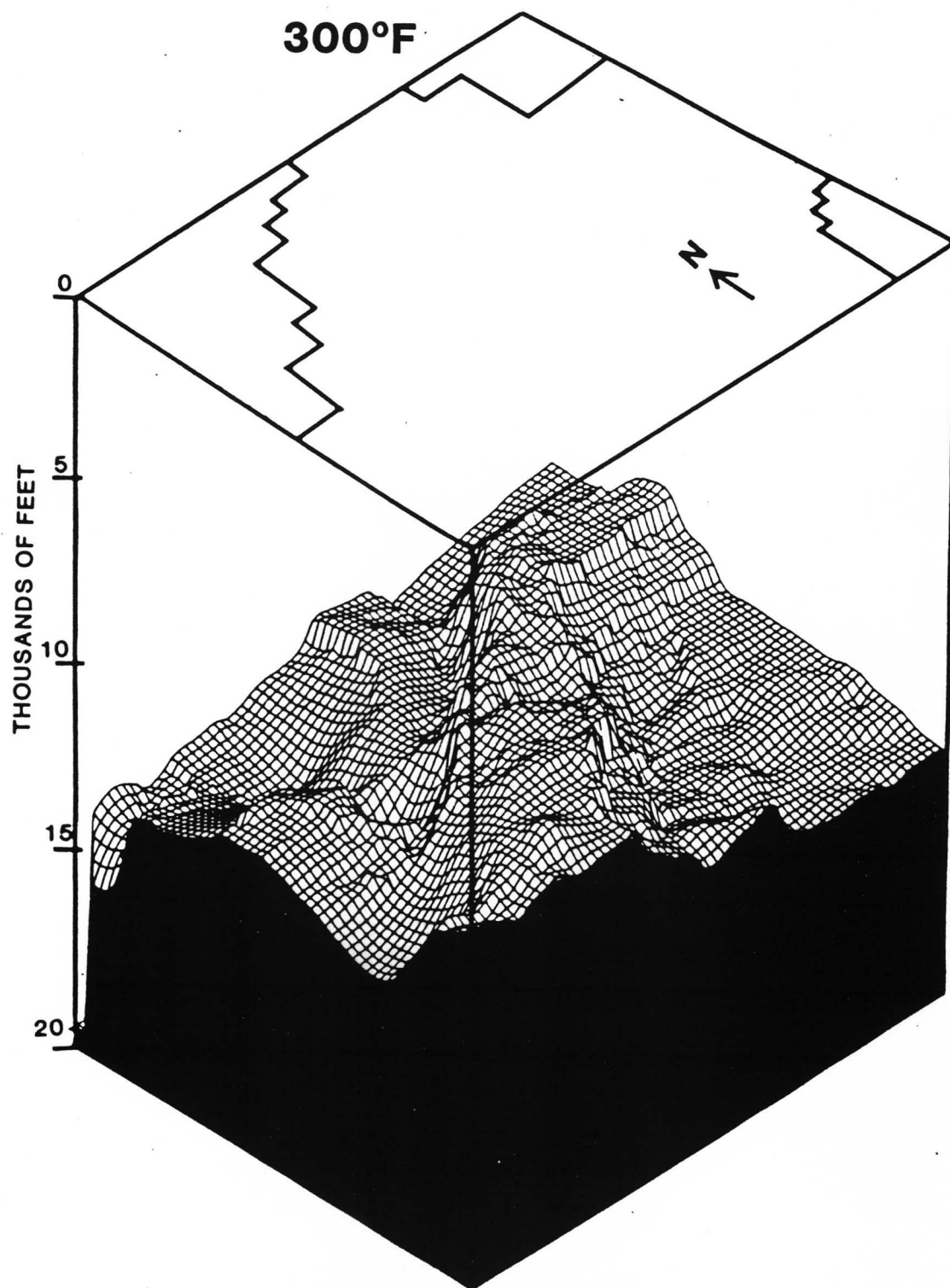


Figure 1.8a: Isotherm from corrected and Kriged BHT data with study area projected above.

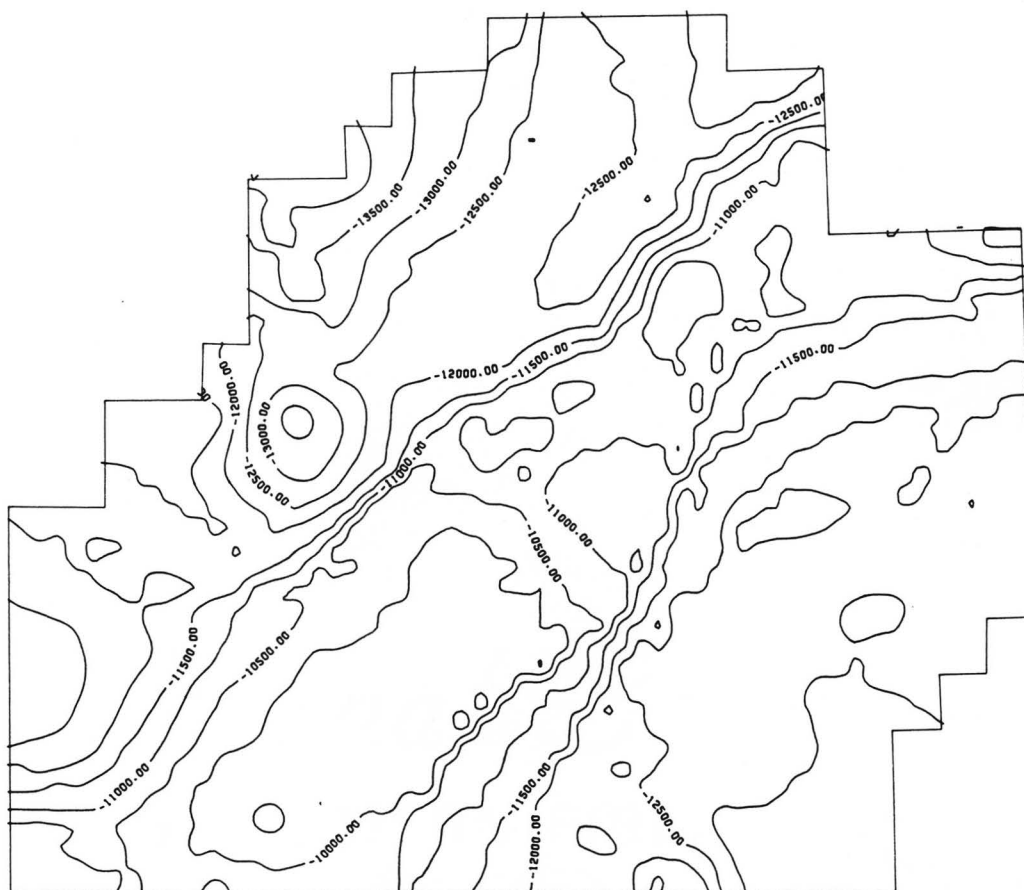


Figure 1.8b: Contoured 300°F isotherm within the study area.

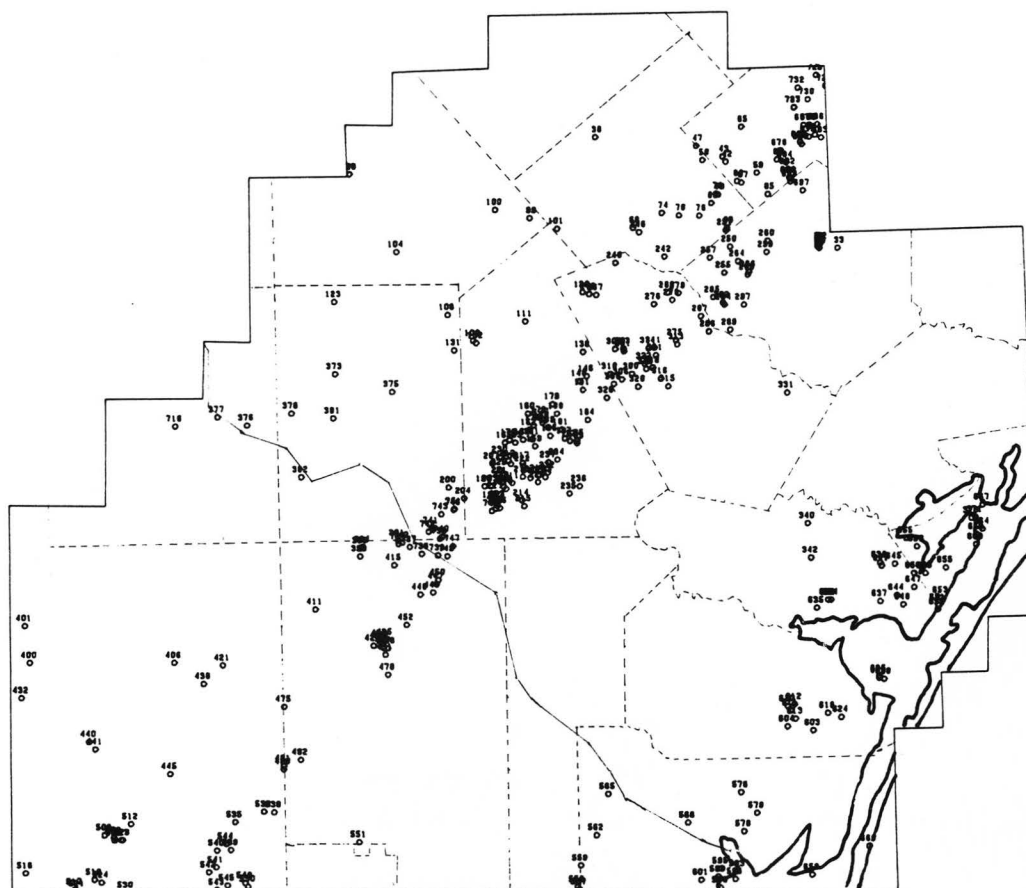


Figure 1.8c: Well locations for the 300°F isotherm.

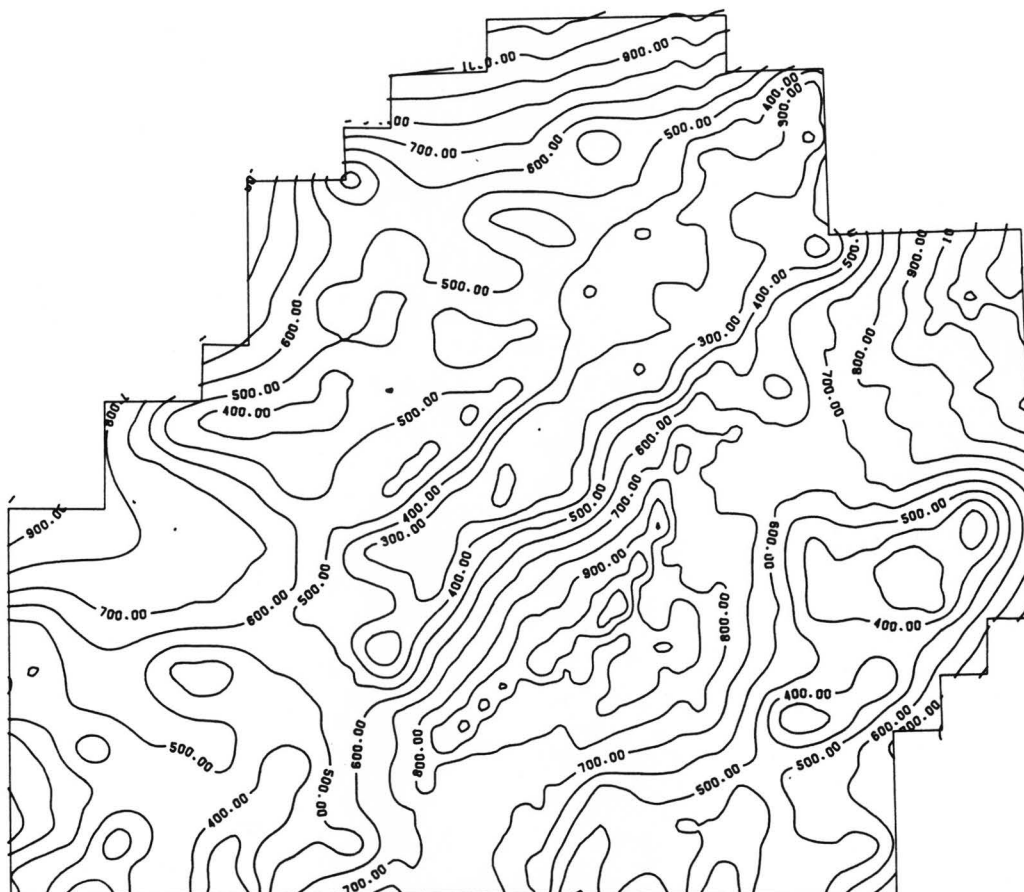


Figure 1.8d: Standard Deviation of contoured 300°F isotherm (computed by Kriging).

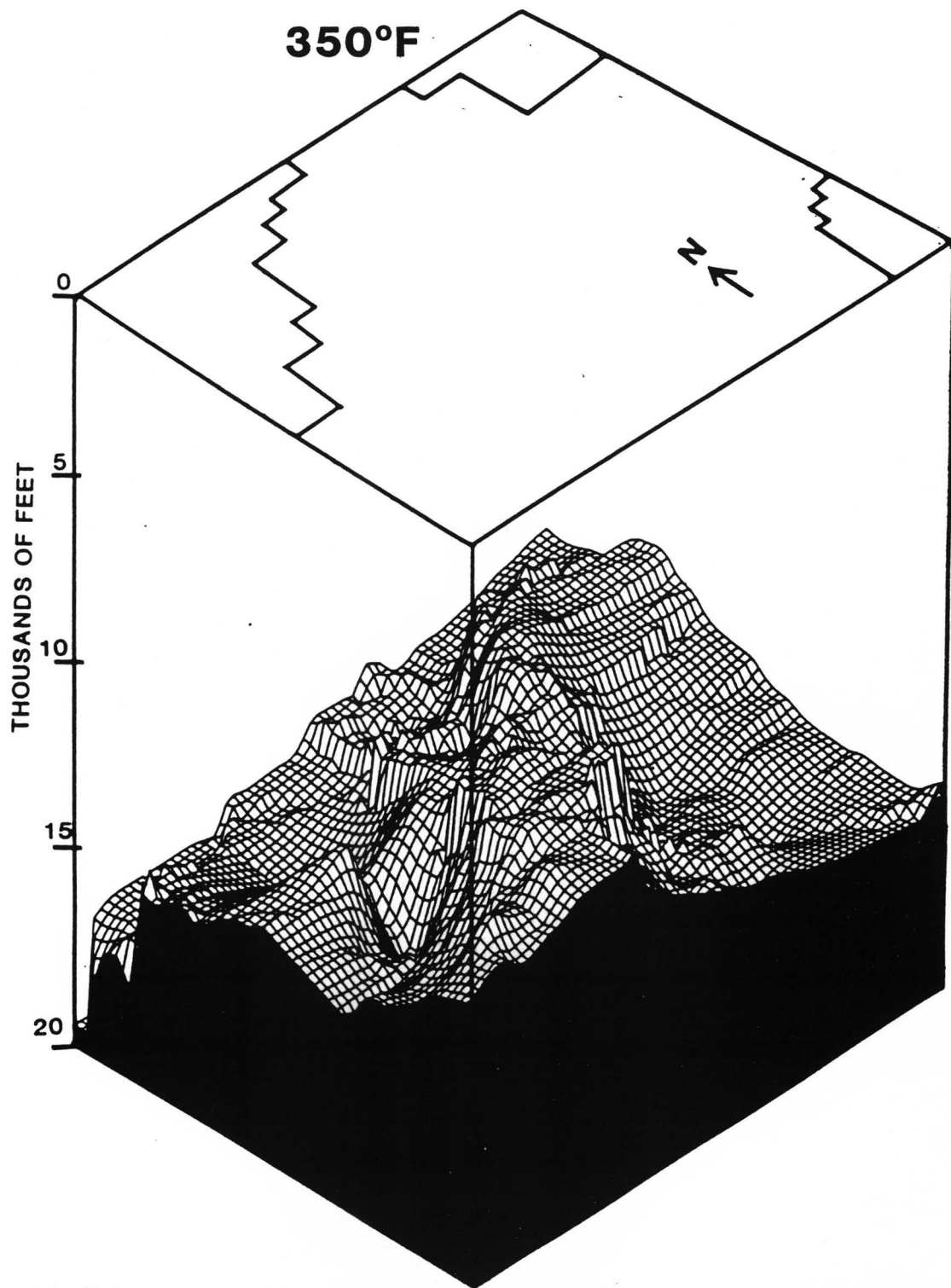


Figure 1.9a: Isotherm from corrected and Kriged BHT data with study area projected above.

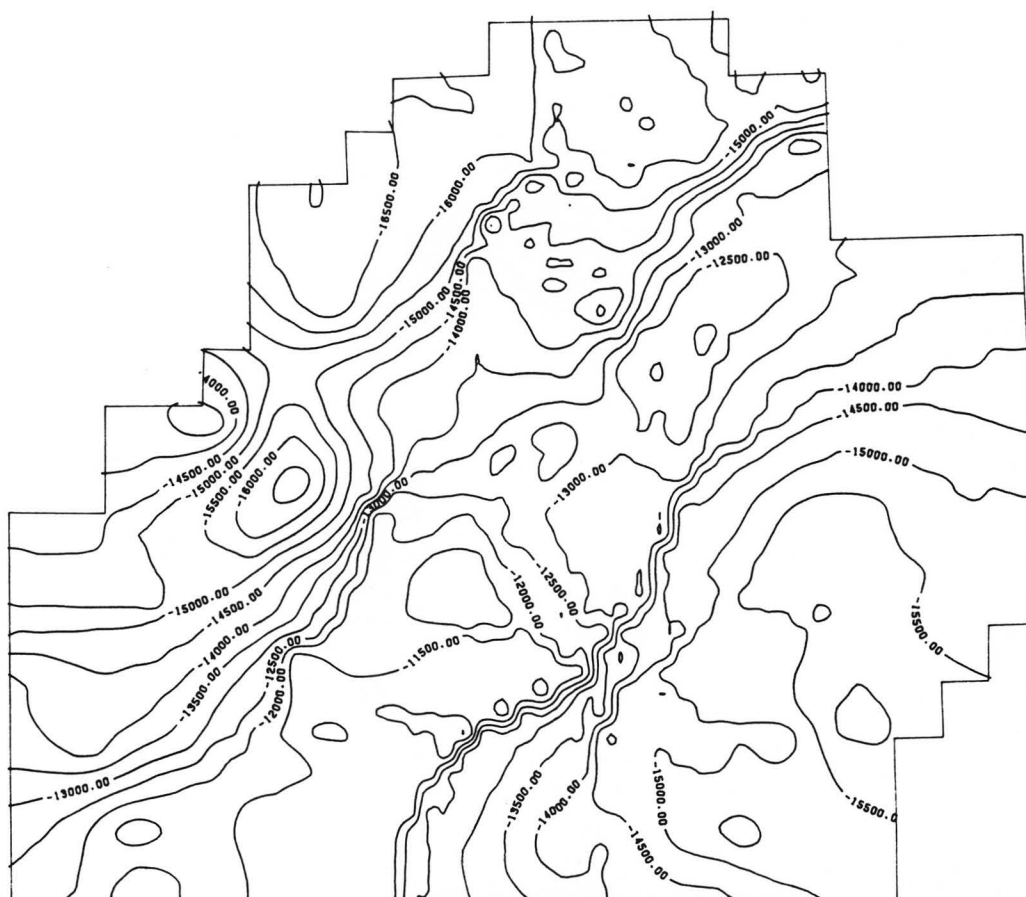


Figure 1.9b: Contoured 350°F isotherm within the study area.

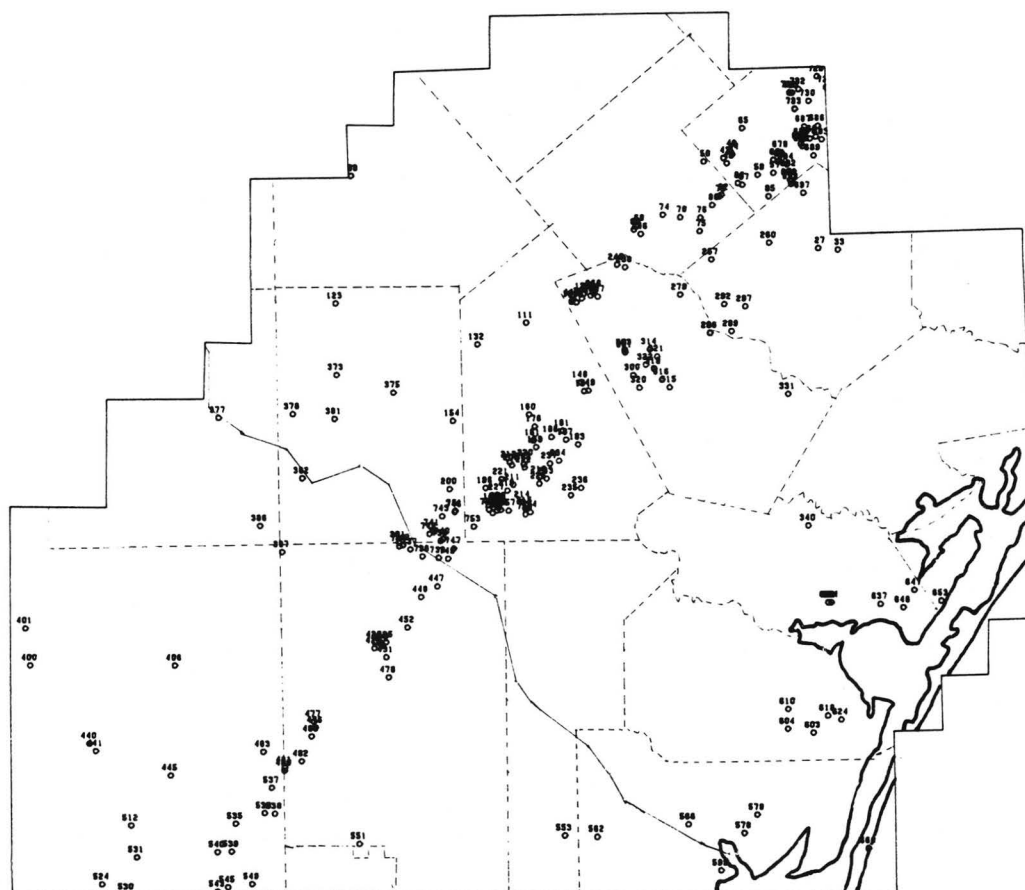


Figure 1.9c: Well locations for the 350°F isotherm.

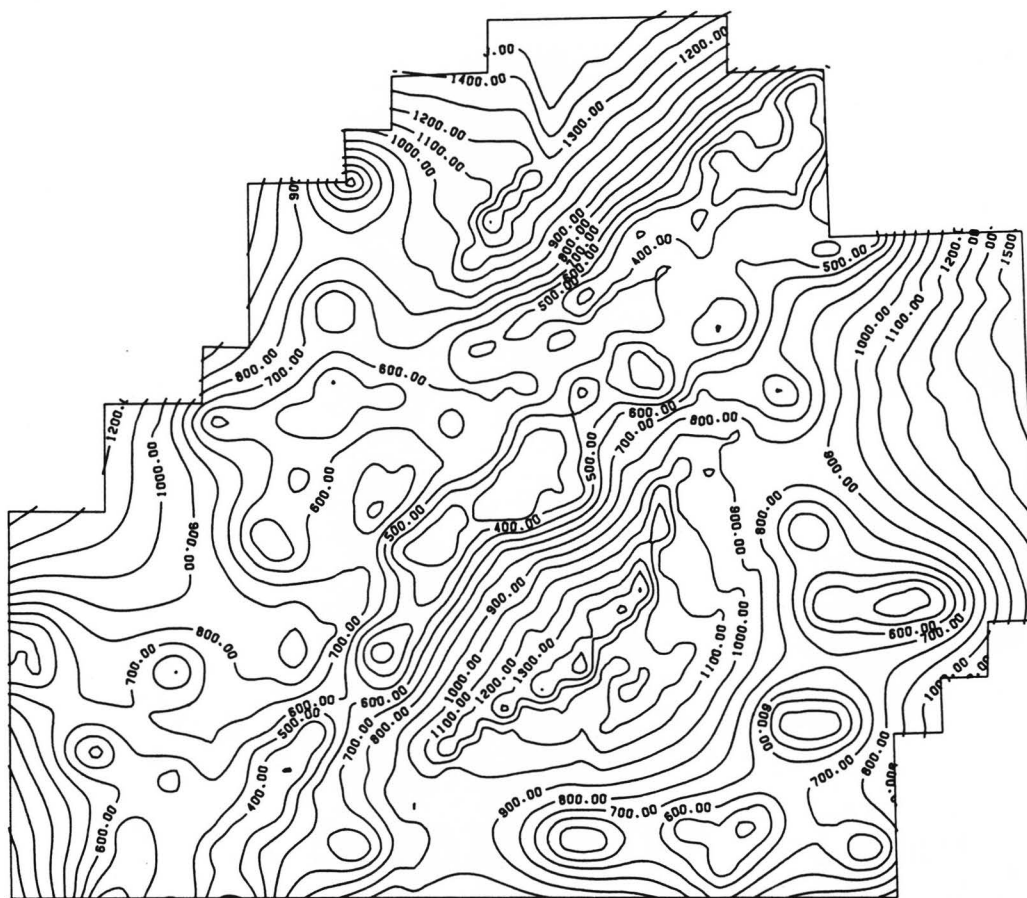


Figure 1.9d: Standard Deviation of contoured 350°F isotherm (computed by Kriging).

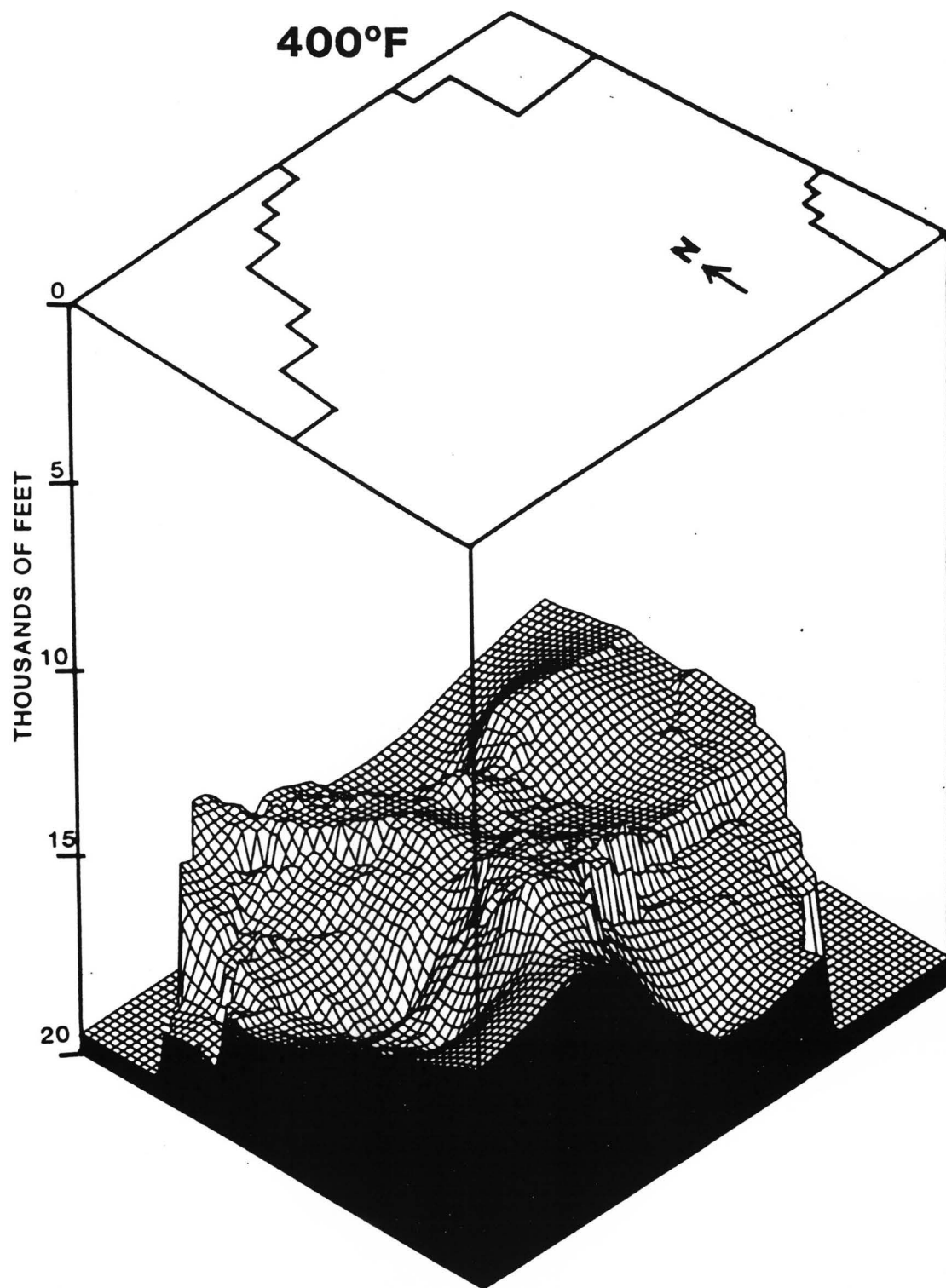


Figure 1.10a: Isotherm from corrected and Kriged BHT data with study area projected above.

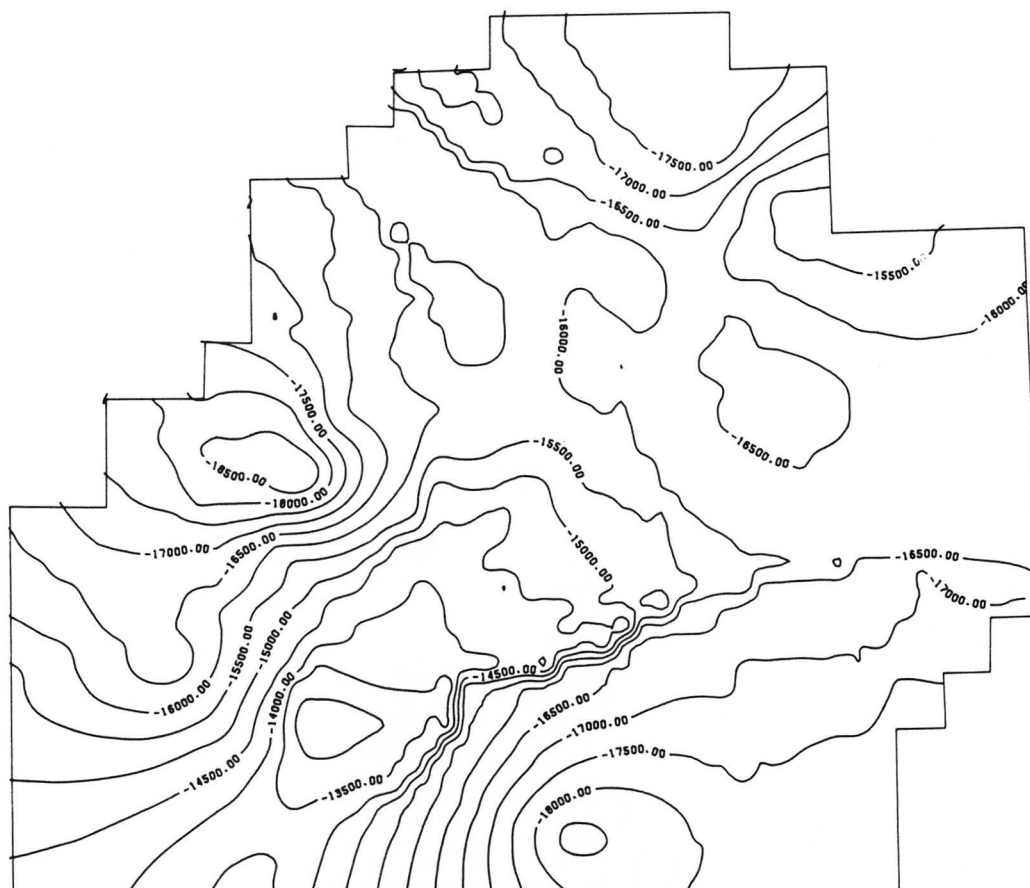


Figure 1.10b: Contoured 400°F isotherm within the study area.

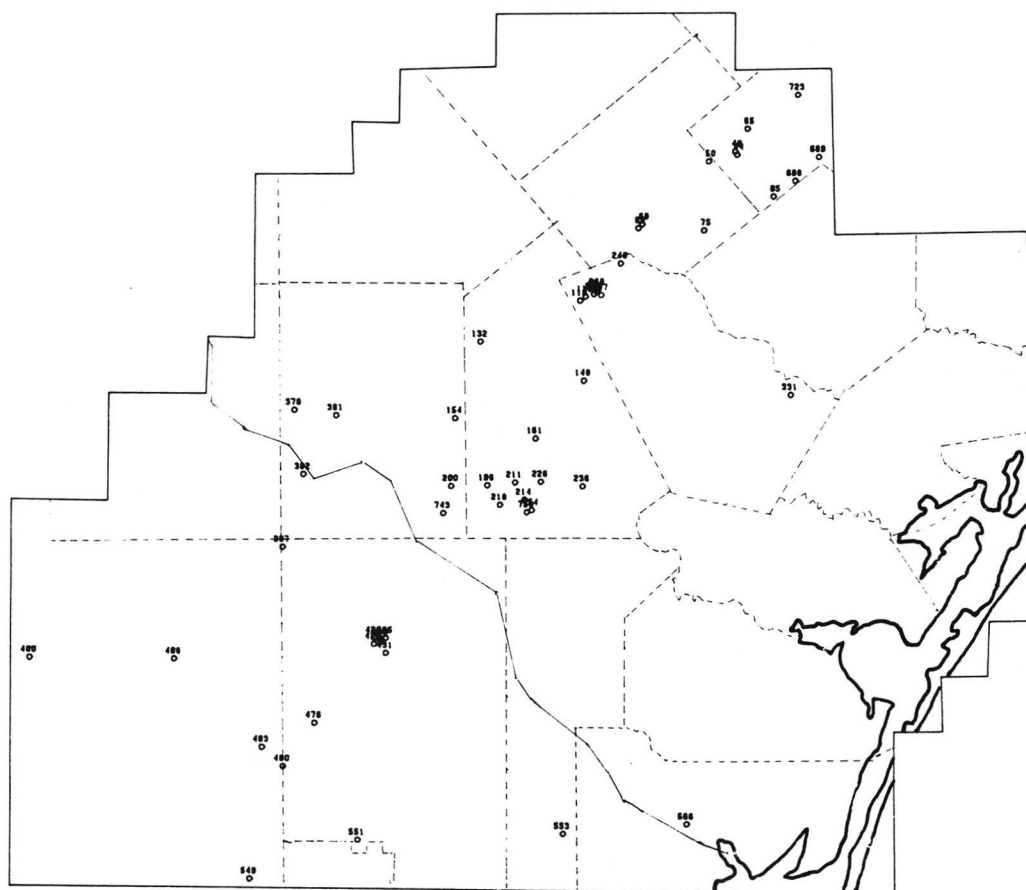


Figure 1.10c: Well locations for the 400°F isotherm.

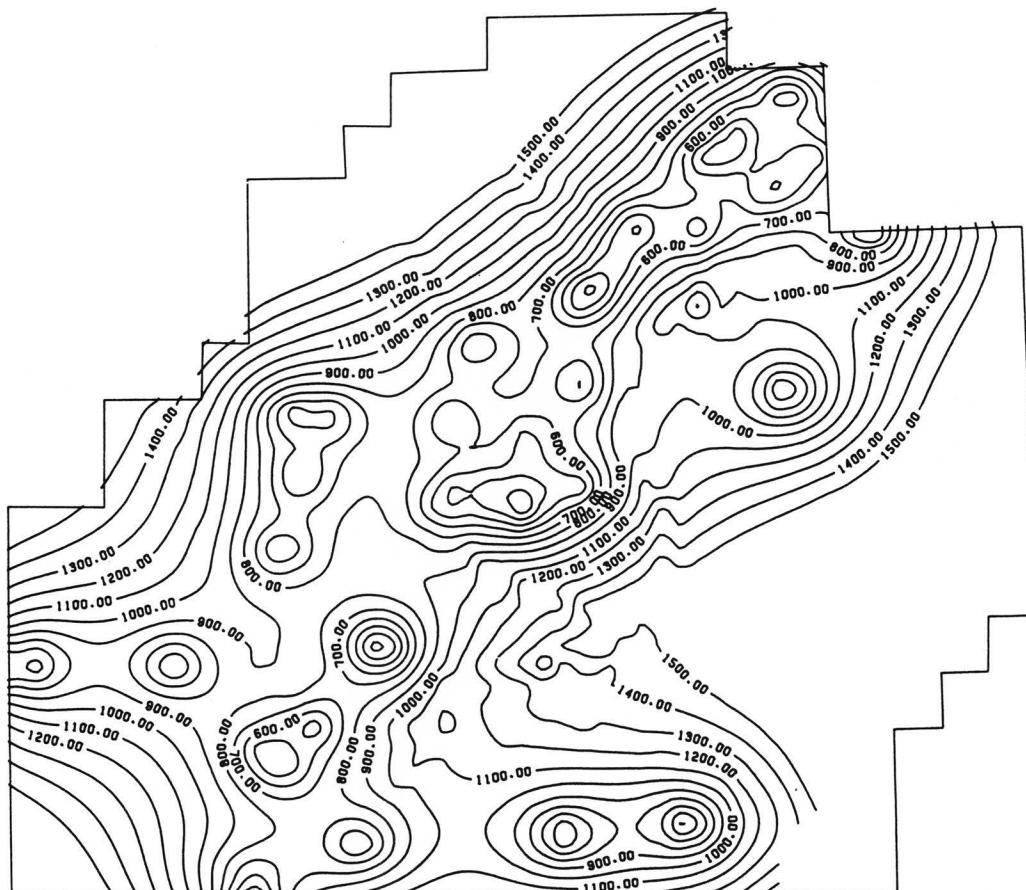


Figure 1.10d: Standard Deviation of contoured 400°F isotherm (computed by Kriging).

ridge trending SW subparallel to the coast. This ridge corresponds to the Wilcox growth fault zone shown in Fig 1.11. It indicates that high temperatures occur shallower near the fault zone than in the surrounding areas and that this effect begins at about 10,000 feet (3048m) and becomes more pronounced at increasing depths.

The isothermal surfaces are informative but they obscure the actual data points. Another way to observe the thermal trends is via a series of charts plotting depth versus temperature for the actual corrected BHT measurements. For this purpose, the study area was divided into 12 subareas, shown in figure 1.11. The subareas were drawn to correspond to the Wilcox growth fault trend (subareas 1 through 5), the Vicksburg/Frio growth fault trend (subareas 10, 11 and 12), and the remaining outlying areas. These temperature profiles appear in numeric order from figure 1.12 to 1.23. On this plot, the slope of a line corresponds to a thermal gradient. The steeper the slope, the lower the gradient. The dashed line on each plot indicates what is generally considered the normal earth gradient of $1.64^{\circ}\text{F}/100\text{ Ft}$ ($30^{\circ}\text{C}/\text{km}$). The + signs indicate individual corrected BHT measurements. The solid line is a linear regression for the data. The correlation coefficients among the regressions were all above .90. Fig. 1.24 summarizes all the regression lines and also tabulates the gradients represented by each regression line for each subarea. Not surprisingly, the Wilcox growth fault zone is characterized by the highest gradients, which far exceed normal gradients. Fairly high gradients are also found in the Vicksburg/Frio

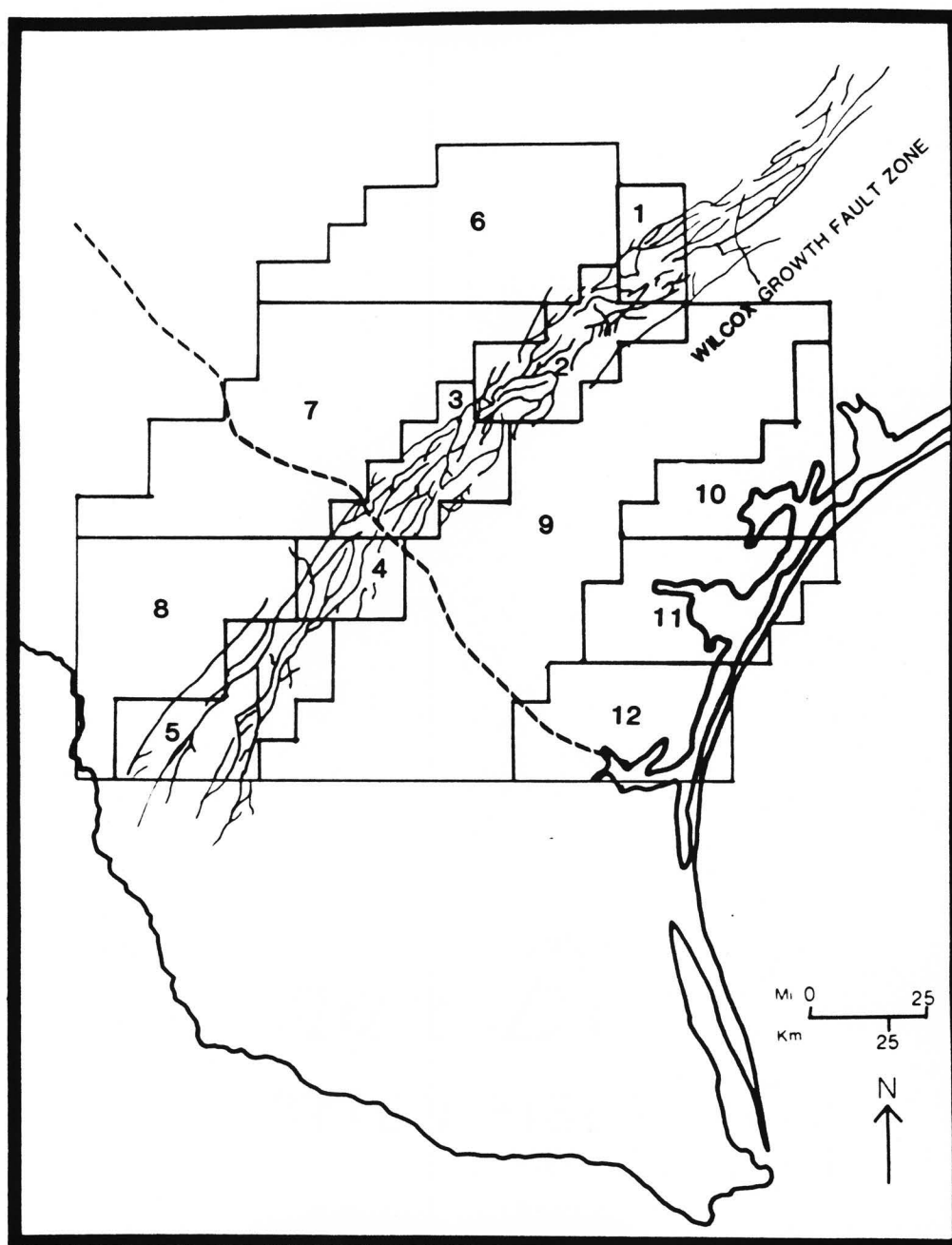


Figure 1.11: Study area showing the twelve subareas and the distribution of the Wilcox growth faults. Dashed line shows the location of the type cross section (Figure 1.2).

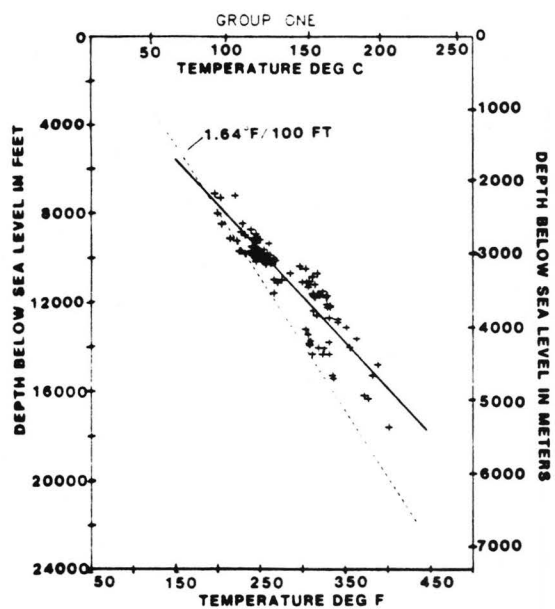


Figure 1.12

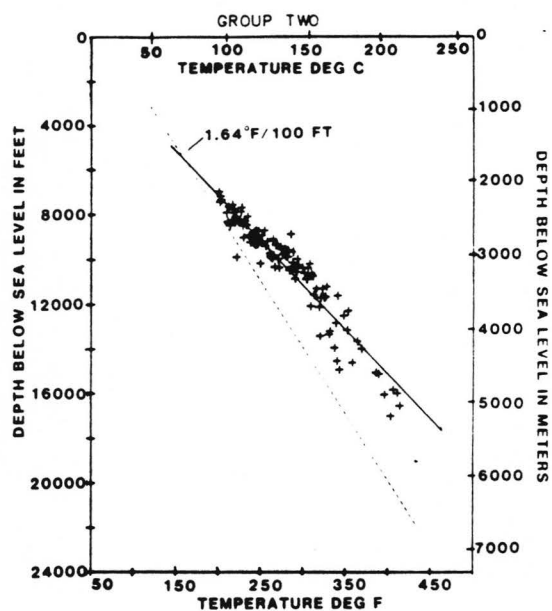


Figure 1.13

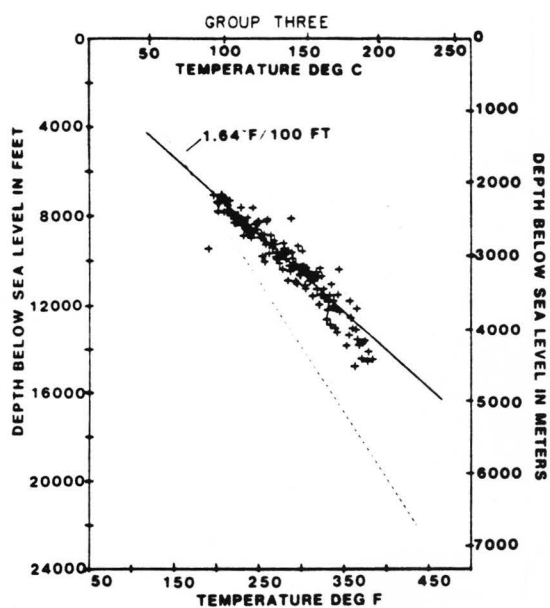


Figure 1.14

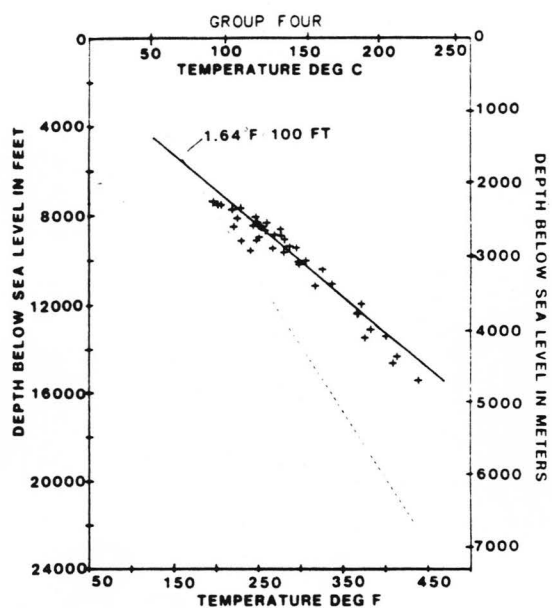


Figure 1.15

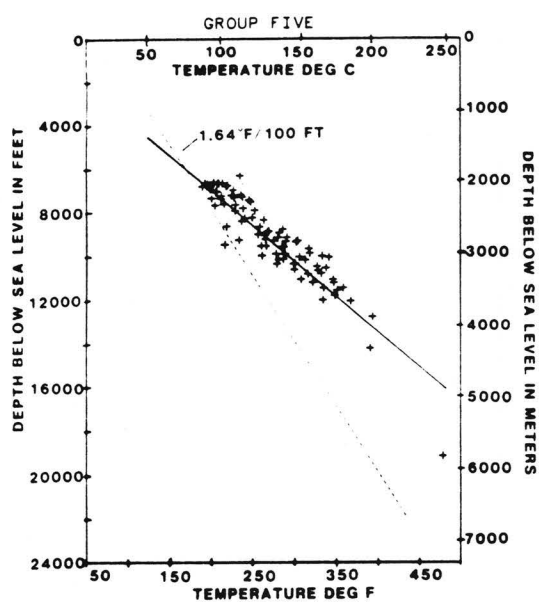


Figure 1.16

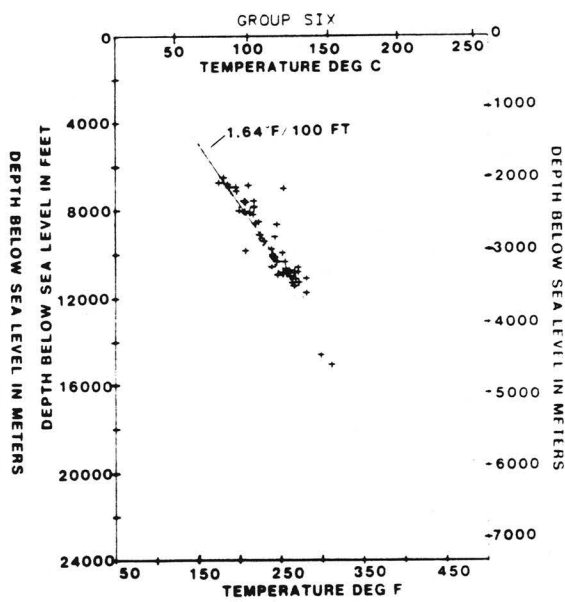


Figure 1.17

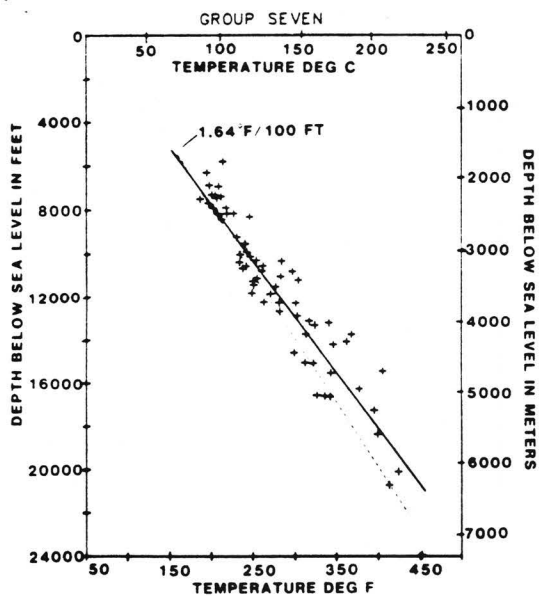


Figure 1.18

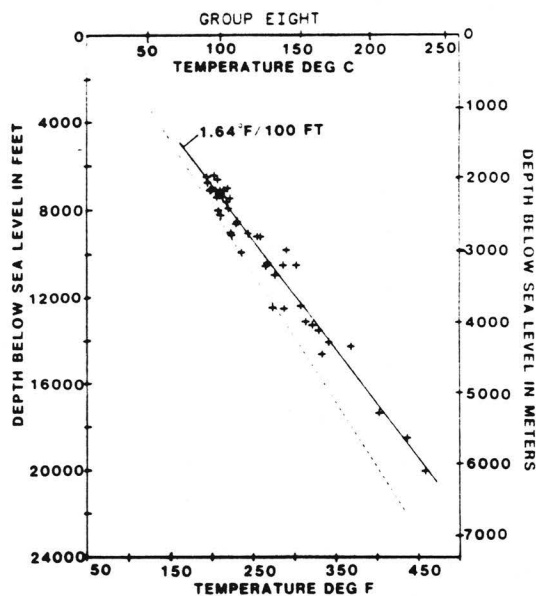


Figure 1.19

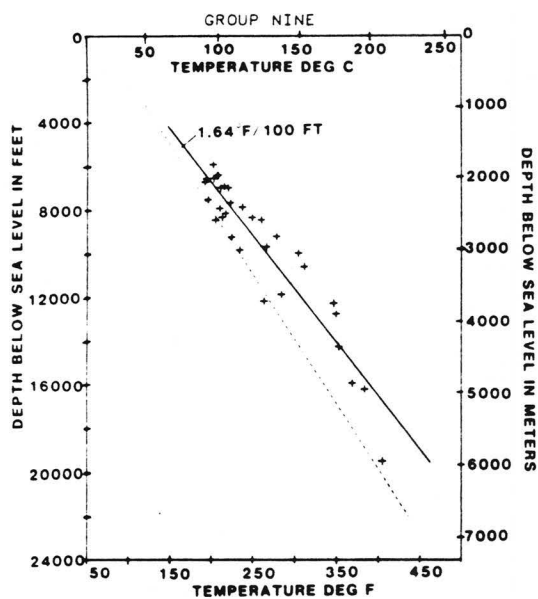


Figure 1.20

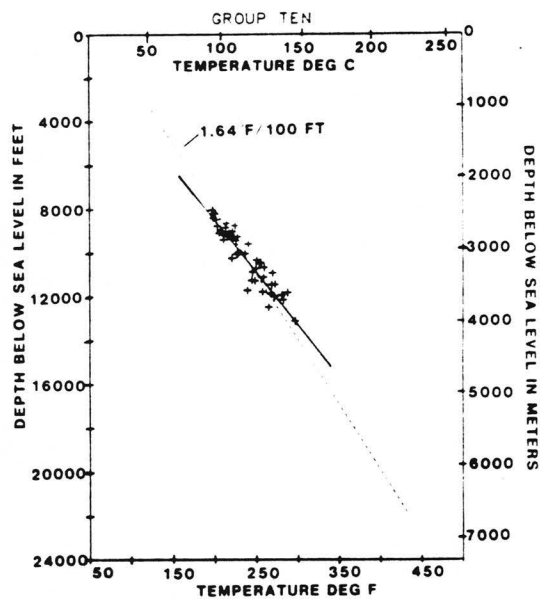


Figure 1.21

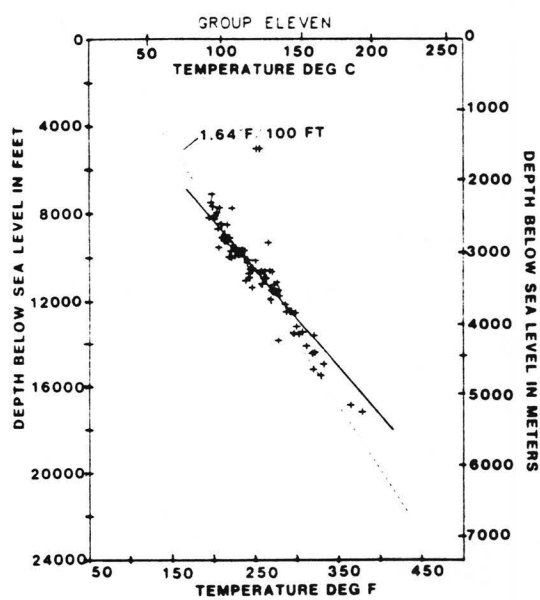


Figure 1.22

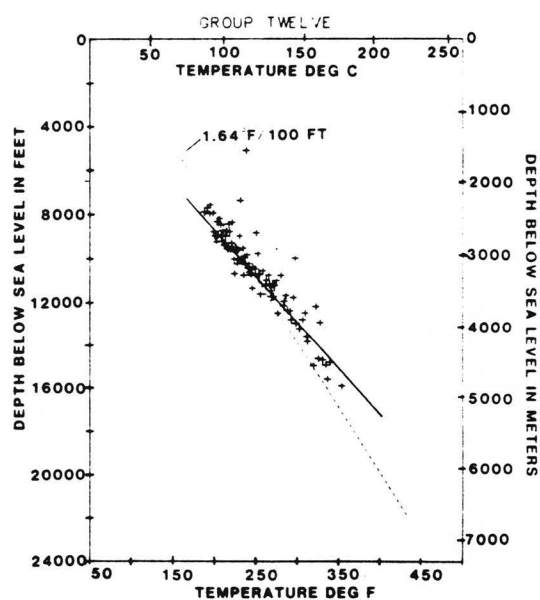


Figure 1.23

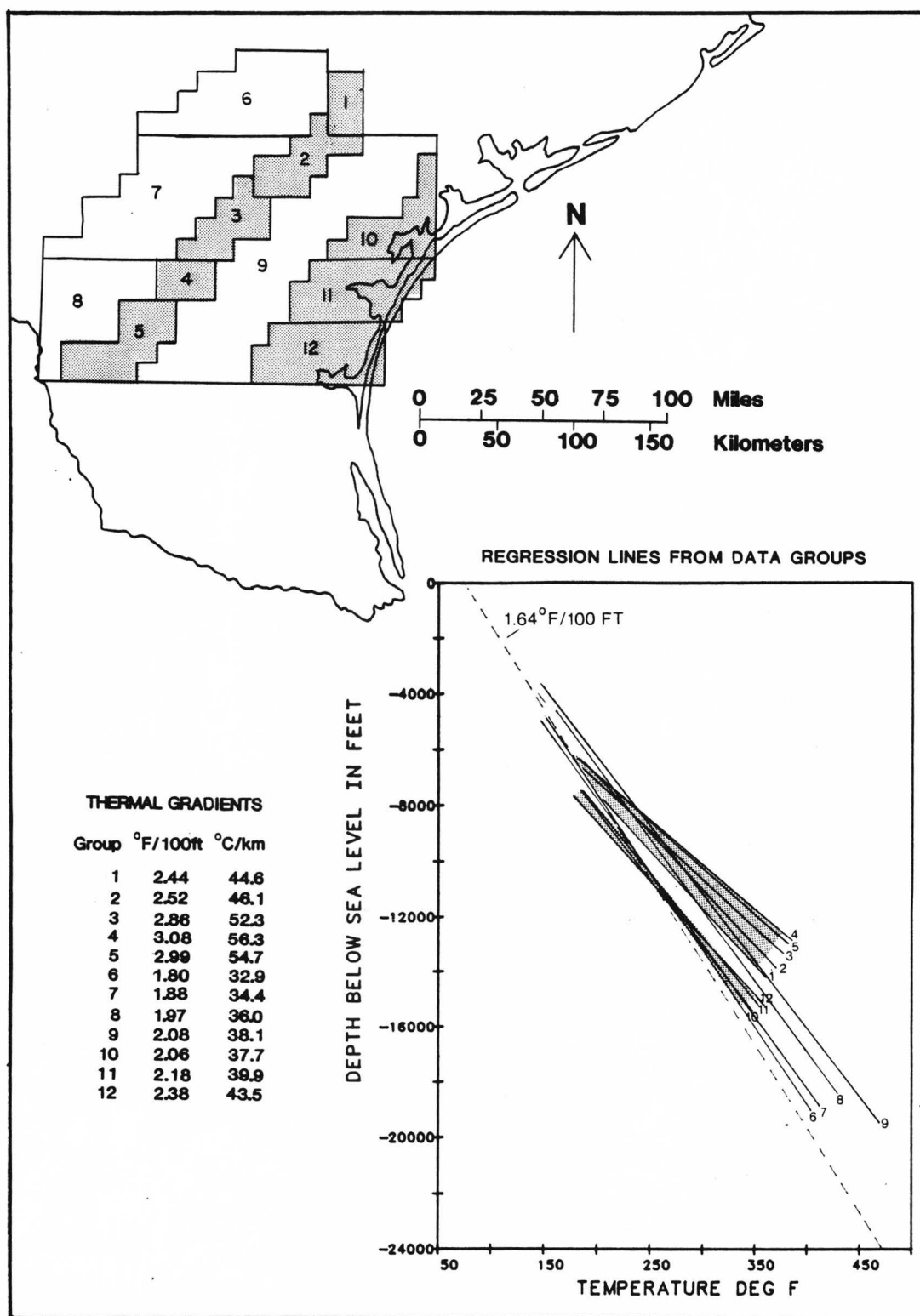


Figure 1.24: South Texas showing subareas where temperature data were collected and their corresponding geothermal gradients, generated by linear regression.

fault zone. The summary of regression lines also reveals a trend of higher gradients toward the southwest.

1.4 GROUNDWATER FLOW

One of the most frequently invoked scenarios to explain anomalously high temperatures in the Gulf Coast was first described by Lewis and Rose (1970) who called on overpressuring as a major contributing factor. The correlation between overpressuring and abrupt increases in thermal gradients is hard to deny. On the temperature profiles of figures 1.12 through 1.23 the departure of the data trend from the normal gradient (dashed line) corresponds closely with the top of geopressuring indicated by the hachured line in figure 1.2 (derived from mud weights). The Lewis and Rose model attributes this to decreased thermal conductivity of the overpressured zone (i.e., the ability of this zone to conduct heat is diminished), giving it insulating properties relative to the surrounding sediments. The decreased conductivity results from the higher porosity and consequent water content of these overpressured sediments. Water has a lower thermal conductivity than mineral grains.

This scenario is not completely consistent with the known configuration of the basin and the temperature data. Since overpressuring largely results from a compactional disequilibrium, its presence correlates to the presence of extensive marine shale bodies whose low permeability causes the disequilibrium. According to the scenario, these same areas should exhibit a temperature build up because

of their insulating characteristics, and the more extensive the shales are, the larger the thermal insulating effect should be. But, in fact, the marine shales are more extensive on either side of the Wilcox fault zone than in the zone itself, which features significant sand accumulation (see Fig. 1.2). Yet the thermal anomaly is greatest within the growth fault zone and declines sharply on either side. Increased thermal gradients are also noted in the other major growth-faulted regions, the Vicksburg and Frio zones, where marine shales are least prevalent. One must therefore conclude that simple thermal conductivity variation is not sufficient to explain the observed patterns. Furthermore, the presence of the growth faults is clearly associated with the development of high-temperature anomalies within geopressure, so that fluid movement along the faults is an attractive mechanism for heat transfer. For reasons stated in section 1.1.3, isolated regions of high basement heat flux were ruled out as a possible factor.

1.4.1 EVIDENCE FOR VERTICALLY MOVING FLUIDS

There is much circumstantial evidence that fluids move upward along growth faults. Galloway (1984) and Smith and others (1982) described evidence in Uranium deposits in the Oakville Sandstone (Miocene) for pulses of deep-basinal fluids from adjoining growth faults. Similar findings from temperature and salinity data were reported by Mumme and Ferrel (1979) for growth faults in Louisiana, and by Kingston (1985) for Oakville associated growth faults. Furthermore, the occurrences of oil often suggests involvement of faults in migration.

Philippi (1965) discusses chemical evidence from shallow oil fields indicating its similarity to deep-dispersed oil, again suggesting vertical migration of fluids.

Petrographers have been confounded by the abundance of cements in Gulf Coast sediments, which imply that large volumes of water have been flushed through the system. Oxygen isotopes suggest that the fluids responsible for this cementation were from high temperature regions, as is also suggested by the presence of sphalerite (K. L. Milliken, personal communication, 1985), and sphene and laumontite (Boles, 1980; Richman and others, 1980) in minor cements. Again, the evidence points toward fluids upwelling from great depths.

Therefore it is not unreasonable that growth faults could act as conduits for upwelling fluids. The result of fluid movement of this type would be transport of heat from the compactional or thermobaric zones toward the surface. This could cause a thermal ridge like the one associated with the Wilcox growth faults. If so, the temperatures observed in this zone suggest that flow is occurring at depths even greater than 20,000 feet (6096 m), steadily diminishing to 10,000 feet (3048 m). This is indicated by the initial prominence and steady dissipation of the thermal ridge through this depth range.

A similar, albeit more subdued pattern is evident in the temperature profiles from the Vicksburg/Frio fault trend. The above-average gradients occur at greater depths because the geopressure system is likewise deeper. The lithology in the Vicksburg/Frio trend is more complex than in the Wilcox trend, and its flow regime may

therefore involve more than the simple vertical flow proposed for the Wilcox growth faults. For example, free convection has been proposed for the extensive normally pressured sand packages in the Frio (Blanchard and Sharp, 1985; Bodner and others, 1985). Free convection would cause temperatures to exceed normal values at the top of the convection cell, while causing below normal temperatures at its base. Hence the low temperatures near the top of geopressuring which are indicated in the temperature profiles 1.21-1.23 may result from the lower portion of a convection cell. The mechanism of the moving boundary effect (Sharp and Domenico, 1976) could also be invoked. When sediment accumulations are sufficiently rapid, deposited material remains cool relative to its burial depth until it can thermally equilibrate. The low temperatures near the top of geopressuring could be explained in this way. Without the shallow temperature data, possible flow mechanisms in this region are conjectural.

It should be noted that free-convection may also be occurring on a regional scale in the basin. Petrologists and sedimentary geochemists have estimated that large volumes of water are needed to account for the volume of cements in the basin (Bjørlykke, 1983; Wood and Hewett, 1982; Land and Dutton, 1979; Sibley and Blatt, 1976). The flow volumes indicated are so large that the recirculation of basinal fluids is the only reasonable means by which they can be supplied. Convection cells might integrate growth faults in their circulation, which presumably would be

confined to the compactional system since fluids could not freely convect downward across the geopressure transition.

1.4.2 REGIONAL TRENDS

An interesting trend is observable in figure 1.24. The summary of regression lines from each subarea shows a strong relationship between geothermal gradient and proximity to the Mexican border. The further southwest the subarea is located, the higher its geothermal gradient tends to be. This is observed in regressions lines 1-5, 6-7, and 10-12. The explanation for this phenomena is unknown. A possible cause is increased heat flux across the basement towards the southwest. Why such a variation in heat flux should occur is unknown. Other possibilities include a systematic regional gradation of lithology. Such a gradation seems unlikely because the trend is perpendicular to the progradation of facies and is therefore correlated to the location of deposition, and not the type of deposition. If regional free convection is somehow integrated into the flow pattern of the major fault trends, one might speculate that the regional variations in the three-dimensional shape of the basin might affect large scale convection cells, which could in turn produce regional variations in geothermal gradients.

1.5 SUMMARY

Data collected for this research consists of bottom-hole temperatures from 750 wells. It is clear that a high-temperature ridge exists in the overpressured portion of the Wilcox growth fault zone. The

anomaly begins to emerge at 10,000 feet (3048 m) and becomes more pronounced at increasing depths. Although high porosities associated with geopressures have been linked to lower thermal conductivities and higher geothermal gradients, this mechanism alone does not explain why the temperatures and gradients are highest in the fault zones, and lower in adjacent areas with similar geopressures. Instead, the fault zones themselves seem to be important, serving as conduits for vertical flow. The Wilcox faults correspond to greater temperatures than do the Vicksburg/Frio faults, perhaps due to the greater geopressuring in the Wilcox faults or other differences in their configuration and location. In addition, thermal gradients increase to the southwest for unknown reasons. The increase may be linked to increasing basement heat flux, but is not understood. Free convection may also be integrated into the flow pattern of the faults on a regional or local scale. The gradients in the Frio/Vicksburg trend, for example, may in part be caused by a shallow, local convection cell. The regional variation of thermal gradients may likewise be tied to a changing configuration of regional convection cells.

The evidence presented in this section is empirical. Although it presents a clear indication that vertical fluid movement in growth fault zones might be occurring, such an approach falls short of quantifying the flow. Variables such as thermal conductivity, hydraulic permeability, porosity, and the interplay between the meteoric, geopressured, and thermobaric zones, must be considered, and can be treated better by a quantitative model.

PART 2-THE MODEL

The temperature data discussed in the previous section support the hypothesis that vertical fluid movement is concentrated along fault-zones in the Gulf Coast Basin. However, because of the many different influences on heat distribution within the basin, it is difficult to distill empirically from these thermal patterns a definitive model of the deep-basinal flow regime. In order to account for the many complicating factors in an integrated and realistic manner, a computer model is used to analyze the relative importance of certain factors vis-à-vis the others. This approach is termed a sensitivity analysis. In the case of observed thermal patterns, a model can, for example, help determine whether or not the concentrated vertical movement of fluids is an essential condition for producing the known temperature field.

I have adopted a computer model for this purpose. This chapter is devoted to the modeling phase of research, and encompasses all its aspects, from the design of the program itself, to its

implementation, and the subsequent conclusions. For clarity, the computer code which performs the modeling operations shall henceforth be referred to as the *program*, and the design of the mesh and its various parameters to which the program is applied shall be termed the *model*.

2.1 MODEL SPECIFICATIONS

The overall goal of modeling was to determine the role of the groundwater flow regime in determining the basin's thermal regime. Using temperature distribution as the primary data, a model was implemented to inversely solve for fluid flow. A two-dimensional program by Smith (1984) was obtained for the study. It features coupled heat/fluid flow at steady-state conditions and calculates densities and viscosities as temperature and pressure dependent variables.

Since the model was two-dimensional, only cross-sections could be modeled. For this purpose, a detailed cross-section compiled from well-logs by W. E. Galloway (written communication, 1985) was chosen as the type section. This particular section was one of the best available and also traversed the study area from which the thermal data was collected. Fig. 2.1 shows its location in the study area and Fig. 1.2 shows a modified version of the cross-section. A second cross-section by Wesselman (1983) which traverses a nearby area was also used as a cross reference. This cross-section also showed temperature data, which generally agreed with my data, but were

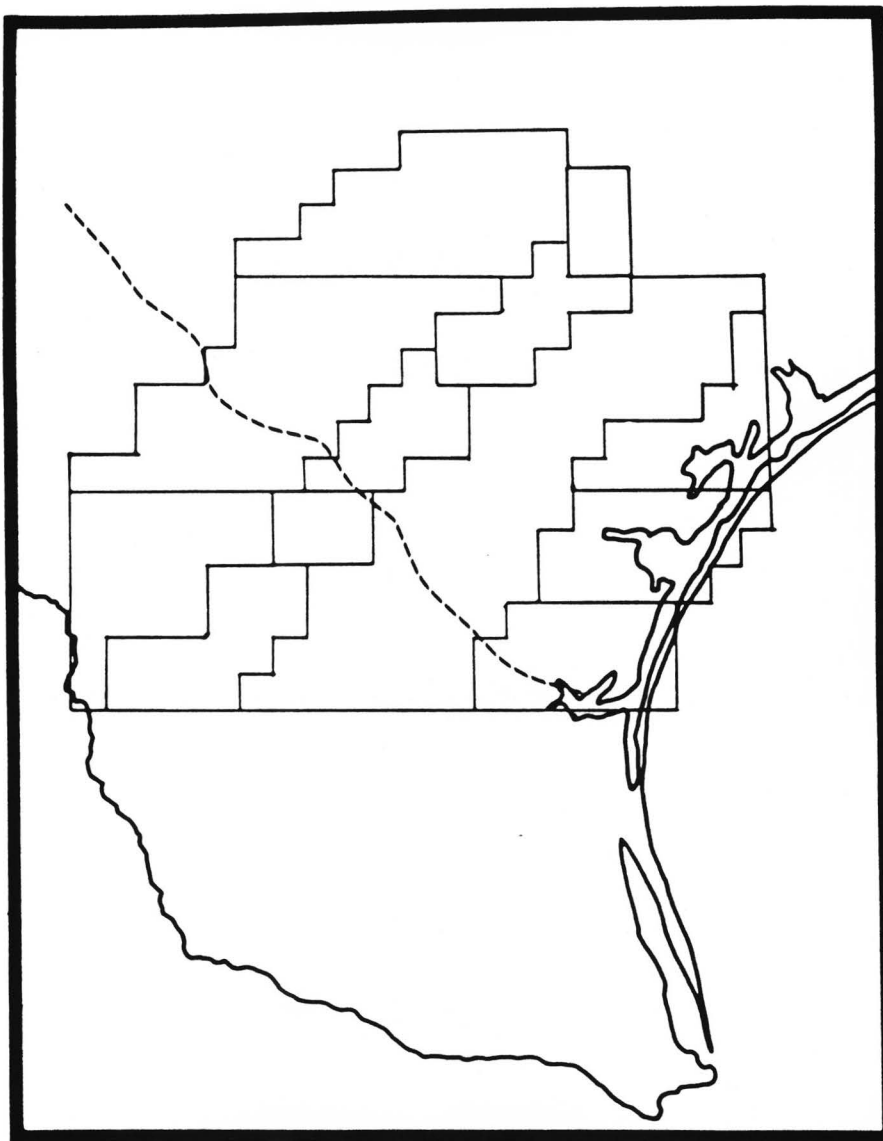


Figure 2.1: Dashed line indicates location of the type cross section (Figure 1.2) in the study area.

derived from fewer points and were not statistically treated. The cross-section also provided pressure data, which also correlated well with pressure data taken from mud-weights. These data consist of contours of the geopressure transition zone and were used to position the top of the geopressure zone on the cross-section.

2.2 FEATURES OF THE MODEL

The first part of the following discussion covers the mathematical basis of the modeling program including its underlying assumptions. The mathematics are covered in greater detail by Smith (1984). A second section deals with the boundary conditions and assumptions.

2.2.1 EQUATIONS OF HEAT AND FLUID FLOW

The building blocks of any computer model are the equations on which it operates. Smith's model uses two general equations, one describing fluid flow, and the other heat flow. The fluid flow equation can be further subdivided into two parts, one for describing the driving forces for flow, and the other to state the conservation of mass at steady-state conditions. The former is of the form:

$$q_i = \frac{-k_{ij}\rho_0 g}{\mu} \left[\frac{\partial h_j}{\partial x_j} + \rho_r \frac{\partial z}{\partial x_j} \right] \quad \text{where } \rho_r = \frac{\rho_f}{\rho_0} - 1 \quad (1)$$

where q_i is the specific discharge in the coordinate direction i , k_{ij} is the

intrinsic permeability tensor of the porous medium, μ is the dynamic viscosity of the fluid, g is the gravitational constant, h is the hydraulic head, ρ_0 is a reference density of water, and ρ_r is a relative density of water derived from ρ_0 and ρ_f , the nodal density of the fluid. Taken together, the permeability tensor, density, gravitational constant, and viscosity terms form the hydraulic conductivity K . The equation can be seen as a version of Darcy's law $q = K \nabla h$ where ∇h , the hydraulic gradient, is reduced to two components, the pressure head plus the elevation head. The elevation head is modified in this case by the relative density term which adjusts for buoyancy affects.

The flow equation must be combined with a continuity equation. The continuity equation is simply the mathematical statement that fluid mass is conserved, ie., in any given portion of the modeled area, the amount of fluid flowing in is equal to the amount flowing out. This equation is of the form:

$$-\frac{\partial}{\partial x} [p r q_x] + \frac{\partial}{\partial z} [p r q_z] = 0 \quad (2)$$

fluid continuity equation

The equation is set equal to zero to specify steady-state conditions. One equation is obtained from the two by substituting the expression for q from (1) into equation (2), this step being omitted here for clarity.

The heat transport equation is composed of three basic terms, set equal to zero (steady-state). The equation is:

$$\begin{aligned}
 & \frac{\partial}{\partial x} \left\{ (n\lambda_{xx}^f + (1-n)\lambda_{xx}^s) \frac{\partial T}{\partial x} + (n\lambda_{xz}^f + (1-n)\lambda_{xz}^s) \frac{\partial T}{\partial z} \right\} \\
 & \quad \text{(x direction heat conduction and dispersion in the solid-fluid composite)} \\
 & + \frac{\partial}{\partial z} \left\{ (n\lambda_{zx}^f + (1-n)\lambda_{zx}^s) \frac{\partial T}{\partial x} + (n\lambda_{zz}^f + (1-n)\lambda_{zz}^s) \frac{\partial T}{\partial z} \right\} \\
 & \quad \text{(y direction heat conduction and dispersion in the solid-fluid composite)} \\
 & - \rho f C_f \left[q_x \frac{\partial T}{\partial x} + q_z \frac{\partial T}{\partial z} \right] = 0 \quad (3) \\
 & \quad \text{(advection of heat with fluid)}
 \end{aligned}$$

where T is temperature, C_f is the specific heat of the fluid, n is the porosity, λ_f is the conduction-dispersion tensor for fluid, λ_s is the heat conduction tensor for solid. The assigned longitudinal and horizontal thermal dispersivities were 100 and 10 meters respectively. The first two terms of the equation account for thermal conduction and dispersion in the saturated medium. The third term accounts for the advection of heat with the fluid.

For the above series of equations, viscosity and density are both figured as functions of temperature and pressure. For density, an iterative procedure is used, based on equations given by Keenan and others, (1978). The basic equation which is solved for density ρ is:

$$p = pRT \left[1 + pQ + p^2 \left(\frac{\partial Q}{\partial p} \right)_{\tau} \right] \quad (4)$$

where p is pressure, R is the gas constant, T is temperature on the Kelvin scale, τ is $1000/T$ and Q is defined as:

$$Q = (\tau - \tau_c) \sum_{j=1}^7 (\tau - \tau_{aj})^{j-2} \left[\sum_{i=1}^8 A_{ij} (p - p_{aj})^{i-j} + e^{-E_p} \sum_{i=9}^{10} A_{ij} p^{i-9} \right] \quad (5)$$

where $\tau_c \equiv 1000/T_{\text{critical}}$ and $\tau_{aj} = \tau_c$ (if $j=1$) or $=2.5$ (if $j>1$) and $p_{aj} = 0.634$ (if $j=1$) or $=1.0$ (if $j>1$). A_{ij} are coefficients which are tabulated in the appendix of Keenan and others (1978).

Viscosity is determined from a set of equations presented by Watson and others (1980). The domain of validity for the equations is given by ($0 < \text{Temp (Deg C)} < 900$) for ($0 < \text{Pres (MPa)} < 300$). The viscosity equation is:

$$\mu = \mu_0(T) \exp \left[p^* \left\{ \sum \sum a_{ij} X^i Y^j \right\} \right] \quad (6)$$

where μ is viscosity, T is temperature on the Kelvin Scale; $p^* = p/p_r$ where p is the density, p_r is a reference density; $X = T^*-1$ where $T^* = T/T_r$, T_r being a reference temperature; $Y = p^* - 1$; a_{ij} are coefficients tabulated in Watson and others (1980); and μ_0 is given by the equation:

$$\frac{\mu_0(T)}{10^{-6}\text{Pa.s}} = \sqrt{T^*} \left[\sum_{k=0}^3 \frac{a_k}{T^*} \right]^{-1} \quad (7)$$

where a_k are coefficients which are also tabulated in the reference.

The two dimensional model features anisotropic permeability, and temperature/pressure dependent density and viscosity. It employs the Galerkin finite-element method using linear basis functions applied to triangular elements. The reader is referred to the the works of Wang and Anderson (1982) and Pinder and Gray (1977) for a complete development of this method. For more about the application of finite-elements in heat transport problems see Mercer and others (1975), Andrews and Anderson (1979) and Li (1980).

To use the model, a finite-element mesh is designed to conform to the area being studied. The mesh is a series of polygons, each being an element. At the corners of each polygon are nodes. Fig. 2.2 shows the mesh used for this study. The following procedure was used to couple the heat and fluid flow equations: A purely conductive thermal regime is initially assumed and the temperature field is calculated by the heat transport equation (3), setting the advective term to zero. Appropriate viscosities and densities at the initial temperatures and pressures are then calculated for each element. These are entered into the fluid-flow equation (1)+(2) to obtain a set of heads at the nodal points. The specific discharge q is then calculated for each node (1). These are plugged back into the advective term of the heat-flow equation (3), and the whole equation is then

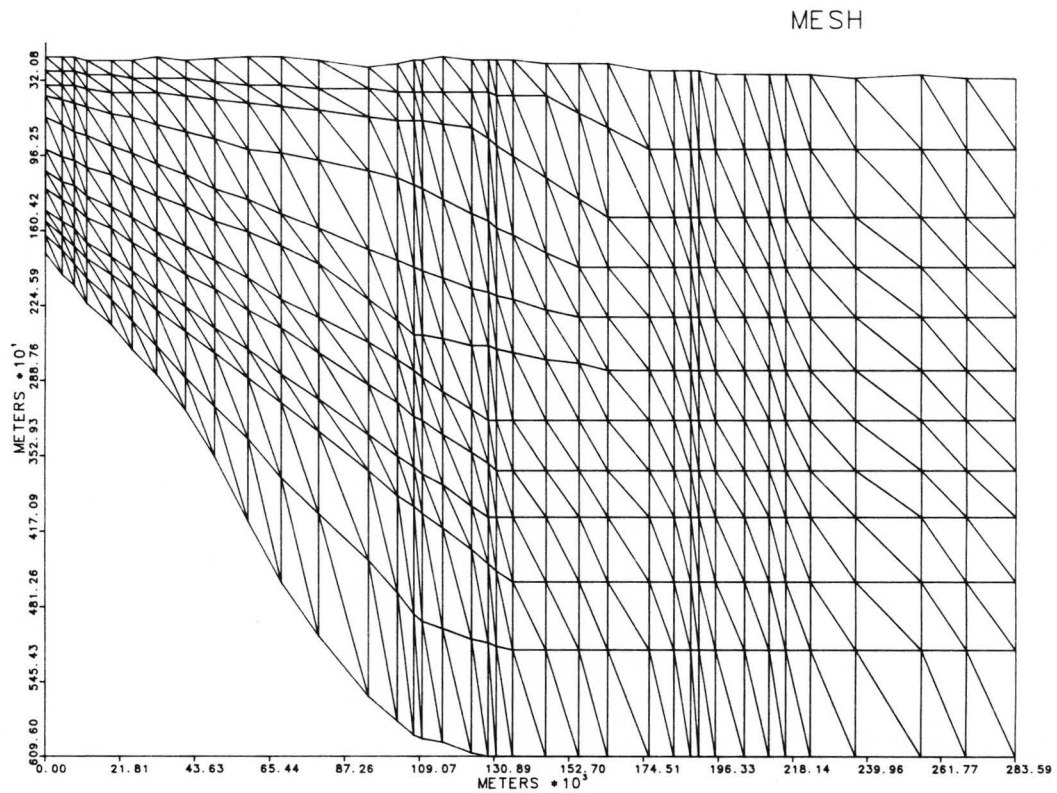


Figure 2.2: The finite element mesh used to model the type section (Figure 1.2), shown at a vertical exaggeration of 40.

recalculated, entering into an iterative sequence. The iterations are halted when the maximum change in nodal temperatures is less than a given tolerance. The tolerance was set at $.5^{\circ}\text{C}$.

2.2.2 BOUNDARY CONDITIONS, ASSUMPTIONS AND LIMITATIONS

The basic boundaries of the model are the four sides of the mesh, representing the bottom, surface, northwest and southeast edges of the cross section.

Basal Boundary. The basal boundary is treated by the program as impermeable with a fixed heat flux value across it. It was assumed that the temperature perturbations of the basin are not the result of any basal heat source or sink, such as a pluton or a tectonic feature. Thus the basal heat flux was treated as uniform across the width of the mesh. The value used was 60 mW/sq m , which is a typical value for stable continental interiors (Reiter 1982, Epp and others, 1970, Smith, written communication, 1985). The Cretaceous boundary was taken as the basement of the basin, and although it is not impermeable, its permeability is probably very low. Note that the northwestern portion of the mesh slopes abruptly downward from shallow depths. This is intended to correspond with the position of the ancestral cretaceous shelf edges of the Stuart City and Sligo reefs. Further southeast, the boundary of the mesh is drawn horizontal at a constant depth of about 19,000 feet (5790 m). The actual depth to the Cretaceous here is not accurately known, but is perhaps as deep as 40,000 feet (12200 m). However, it was impractical to include greater

depths in the model because of the lack of well control and temperature data.

Surface Boundary. The upper boundary of the mesh is a constant head boundary with a fixed temperature at 20°C. Because the temperature is fixed, the model can not account for any local heating at the water table due to upwelling fluids nor can it account for possible thermal springs. Since the meteoric system is not the prime focus of this investigation, these limitations are inconsequential.

Lateral (Vertical) Boundaries. The sides of the mesh are treated as impermeable (no flow) and are also thermally nonconductive. This is a somewhat unrealistic configuration. The mesh does not extend to the outcrop of many of the lower formations, nor does it go far offshore. Fluid and heat transfer could occur across both boundaries in nature. To resolve this problem, fluid flux terms were added to selected nodes on each of these boundaries.

Flux Nodes. Several flux nodes were assigned to the lateral boundaries. Two on the northwest boundary near the outcrop belt of the Wilcox formation were assigned a combined flux of -0.00014 Kg/s , which is about 1/2 ml per hour. The negative sign signifies injection of fluid. On the other side of the mesh were placed three constant flux nodes which withdrew fluid at the equal but opposite rate of $+0.00014 \text{ Kg/ms}$. This arrangement simulated the passage of fluid into the section from outcrop, and also from the section into the Gulf. These fluxes proved to be rather unimportant in their effect on the pressure and temperature fields, and consequently their assigned values could be

varied considerably without appreciable influence. Aside from the two negative flux nodes on the outcrop side of the mesh, other recharge of the meteoric zone was ignored.

Because the model is steady-state and noncompacting, flux nodes were needed to simulate compaction-driven fluid in the geopressured zone. The model was first designed without such nodes and the result was a complete absence of geopressuring. Several distributions of these flux nodes were tried, and it was determined that fluxes assigned just to the basal nodes worked best.

Mathematical Assumptions. As mentioned above, the program calculates appropriate densities and viscosities as functions of both temperature and pressure (equations 4 -7). The subroutines which handle these calculations are equipped to treat pure water in fluid, steam, or supercritical phases. However, multiphase flow is not allowed. In addition, the fluid is required to be in thermal equilibrium with the medium at all times. Thermal conductivity of the rock is assumed to be constant, although in reality it varies slightly as a function of temperature. Thermal conductivity of water is treated as a function of temperature. Porosity is also assumed to be nontransient.

Chemistry. Chemical processes are extremely difficult to couple with groundwater models because of the numerous calculations which even simple chemical models require. Including such calculations in an iterative sequence across an entire mesh would result in unmanageable computation times. As a result, both density and fluid flow calculations are somewhat handicapped. Fluid samples

collected from Tertiary formations in the Gulf (Fisher 1982, Land, personal communication, 1985) indicate that the salinity (ppm) of formation waters is significantly correlated to its in-situ density. The modeling program assumes only pure water. In an attempt to circumvent the problem, a reference density corresponding to a 50,000 ppm solution at standard state (1.053 gm/cm^3) was substituted for the reference density of pure water in one test run. This had no effect on the modeled temperatures, and only a subtle effect on the modeled pressures, mainly in the meteoric zone. In the modeling program, the meteoric zone can not be assigned a separate reference density from the geopressured section. Likewise, viscosity and density subroutines were written for pure water only. Hence, high ppm pore fluids could not be realistically considered in the study. Additionally, clay dehydration reactions may contribute a significant quantity of fluid at particular horizons. Some simulation runs of the model were conducted using assigned fluid flux nodes in areas of critical temperatures and pressures to simulate such fluid release, but this failed to produce correct temperatures or heads (simulating the process in this way did not take into account the heats of reaction).

Assessing Thermal Conductivity. The model requires the assignment of both a thermal conductivity and porosity field. The thermal conductivity is assigned to the solid portion of each element, and the porosity is therefore used to determine the fraction of the element that is assigned the thermal conductivity of fresh water at the appropriate temperature and pressure. Unfortunately, there has been

very little research into the thermal conductivities of the solid portion of rocks. Instead, there have been studies which list thermal conductivities for typical rock-types. These values are given without regard to porosity, water content, or pressure. Pressure is important because an important factor in the thermal conductivity of the solid matrix is the nature of the grain to grain contacts, which may vary as a function of pressure. Of course, the mineral composition of the sediment grains is also quite important because different minerals have widely different conductivities. Therefore, the thermal conductivity of the solid matrix should be dependant upon rock type. Because of the lack of research in this field, the only guide to thermal conductivities at present are the values for typical rock types. The porosity field modifies these values.

Limitations of Scale. The mesh and all the output related to the mesh have been printed at a vertical exaggeration of forty. At true scale, the actual cross section represented by the mesh is very wide relative to its depth as shown in Fig. 2.3. Since the parameters in the design of the model are assigned on a per element basis, and the width of each element is at least several kilometers, many details of the cross section were omitted. A finer mesh could have been used to increase resolution, but this would have caused longer run times and an increased complexity in the set up and calibration of the model. Of course, an important consideration in the design of the mesh was the quality of the data.

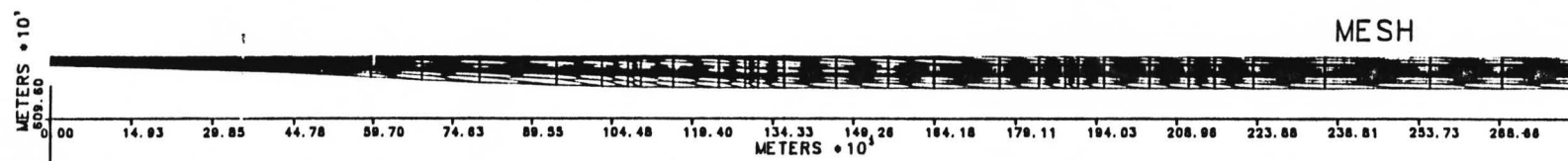


Figure 2.3: The finite element mesh shown at true scale.

Generally, most of the loss of detail was forfeited in the shallow meteoric portions of the basin. The deeper areas are apparently simpler lithologically although there are fewer well-logs available to document this fact. Because of scale, faults were treated as zones rather than as discrete discontinuities, and these zones thus reflect an averaged behavior. Modeled fluid velocities may therefore be less than the actual velocities that may exist in discrete portions of the fault zone. Likewise, much of the anisotropy caused by the small scale interlayering of muds and sands was averaged by assigning anisotropic permeabilities to the grid elements.

2.3 DESIGN

2.3.1 CALIBRATION

Calibrating adjusts the model to accurately simulate the actual system. It involves repeatedly running the model, checking the output, making revisions, and then rerunning it. In an ideally calibrated model, all of the independent variables are known. The model is adjusted so that these are all accurately reproduced by the model, and the dependent variables are then computed. Such ideal cases are rarely realized.

The primary calibration tool for this investigation was the temperature data. Other constraints were the stratigraphy and the top of geopressure plus a general knowledge of the shape of the pressure-versus-depth profile. In addition were crude estimates on the magnitude of overall vertical flow from the deep basin (section 2.4.2).

The modeling procedure was as follows: 1) Four fields were assigned for vertical and horizontal intrinsic permeability, thermal conductivity, and porosity. For each field a code map was constructed where each element is represented by a coded number which is assigned a value for the particular parameter. 2) Flow rates were assigned to constant flux nodes, keeping the total flux as low as possible. 3) These parameters were then run with the program.

The output from the model includes plotted isotherms and hydraulic heads. The isothermal plot was superimposed on the isothermal plots compiled from temperature data and any discrepancies were noted. The same was done for the head plot. Adjustments were made to the various design aspects on a priority of permeability fields first, then if necessary flow rates, followed by conductivities and finally porosities. Initially, several parameters were adjusted between each run. As the model came closer to meeting the calibration criteria, only one or two changes were made per run.

2.3.2 THE MESH

The mesh (Fig 2.2) corresponds to a cross-sectional width of 176 miles (283 km) and a depth of 20,000 ft (6096 meters). In this program, four-sided polygons are each divided into two triangular elements. All rows and columns of nodes are continuous. A general rule in designing the elements is to avoid any elements with obtuse angles. For mathematical reasons, obtuse angles will lead to erroneous values and convergence problems. Convergence problems may also result from

extreme differences in assigned parameter values in adjacent elements. As a further consideration, to compensate for some of the scale problems mentioned earlier, nodes were concentrated in areas of greatest interest, such as in the growth fault zones.

2.3.3 THE CODE MAPS

Porosity Field. The data used to code for porosity was mostly taken from Loucks and others (1979), Bonham (1980) and Fyfe and others (1978). In general, porosity was treated as a function of depth, ranging from 35% at the surface to 5% at the greatest depths. Additionally, this decline with depth was suspended in the geopressured transition zone. Changes made to the porosity field proved to have very little effect on the modeling results. Porosity is only used in the conduction terms of the heat transfer equation (3), and its low sensitivity seems to indicate that saturated porosity is perhaps not as significant a factor in basinal thermal patterns as has been suggested by Lewis and Rose (1970).

Thermal Conductivity. The thermal conductivity values were assigned using as a guide a table from Reiter and Tovar (1982) which lists thermal conductivities for various rock types. The conductivities were assigned to the mesh according to the rock types shown on the type-section. Somewhat lower conductivities were found to be necessary for the deep marine shales than the values listed in the table for shales, but this is justifiable in that much of the marine shale occurring under geopressures has been found to be more like a "dense

clay" than lithified rock (Weaver & Beck, 1971). The values chosen were consequently intermediate between shale and clay. Additionally, the deepest marine shales were assigned higher conductivities because they are presumably more dense and consequently more conductive.

Permeability. The data presented in Loucks and others (1979) was used for permeability values. This data consisted mostly of whole core analysis plotted against depth. A major draw-back with this type of measurement is that it ignores large scale permeabilities created by faults and fractures. These are perhaps more important in determining hydrologic characteristics of a large area. Additionally, permeability is not only a function of depth but also of rock type, pressure and chemical reactions. The whole core analyses are relative to atmospheric pressure. Nonetheless, the range of values for each depth was quite great, which provided room for adjustments in the calibration process. Permeability was the most sensitive of the input parameters.

A matter of some speculation is the contribution to permeability made by hydraulic fracturing, also known as microfracturing. According to Fyfe and others (1978), microfractures are theorized to originate in flaws, which grow when pore-fluid pressure exceeds the least principal stress by an amount equal to the tensile strength of the rock. The stresses on rock can be resolved into three vectors, one for each dimension, x,y, and z. These are called the principal stresses, the smallest of which is called the least principal stress. In undeformed rock these stresses are usually compressional. Pore pressures tend to oppose compressional stresses equally in all

directions by forcing pore spaces open. When pore pressures exceed the least compressional force by an amount greater than the matrix can tolerate (the tensile strength), the rock ruptures, causing a fracture. Fracturing will occur in the direction perpendicular to the least principal stress. In a compacting basin the least principal stress is usually horizontal, while the greatest principal stress, resulting from the weight of sediments above, is vertical. Microfractures are usually oriented vertically (Jaeger & Cook, 1979). Stronger rocks are more susceptible because they tend to have greater stress differentials (the difference between least and greatest principal stresses). The strongest rocks should be those in the deeper basin, owing to their degree of consolidation. On the other hand, stress differentials decrease with depth. How these opposing factors control the occurrence of hydrofracturing is currently unstudied, but faulted areas may be particularly susceptible owing to an abundance of suitably oriented flaws.

2.4 RESULTS

The investigation was broken down into three simulations, each testing a basic type of flow regime. The first considered the effects of evenly dispersed advection while the second concentrated the advecting fluids along faults. The third was a special case of the second, adding a speculative region of high thermal and hydraulic conductivity in order to more closely meet the constraints.

Several criteria were employed to calibrate the model. They were fluid flux, fluid pressure, and temperature. The degree to which model output matches the calibrating criteria is a measure of how well the design of the model depicts the actual system. The same criteria were used for each simulation, and the following discussion systematically compares them to the model's output. The implications of the results are covered in the discussion section 2.5.

2.4.1 SIMULATION ONE - CONDUCTION DOMINATED

The object of this simulation was to test the hypothesis that thermal patterns are basically controlled by the movement of dispersed (as opposed to fault-concentrated) compactional fluids coupled with geopressure-dominated thermal conductivities.

Flux Nodes. Fluid flux nodes were needed to simulate the release of compaction fluids, thus establishing a geopressured zone. These fluxes were evenly distributed across the base of the section. Because fluid fluxes must be assigned to nodes, which are point sources, the basal elements were given artificially high horizontal permeabilities. If this were not done, all nodes, being separated by several kilometers, would produce high pressure plumes. By assigning high horizontal permeabilities to the basal elements, a permeable avenue was created which permitted the simulation of a diffuse upward flow that would be typical of a hydraulically homogeneous compacting basin. Fig. 2.4 shows the flux distribution along the basal boundary. During calibration, very low fluxes were initially assigned. After

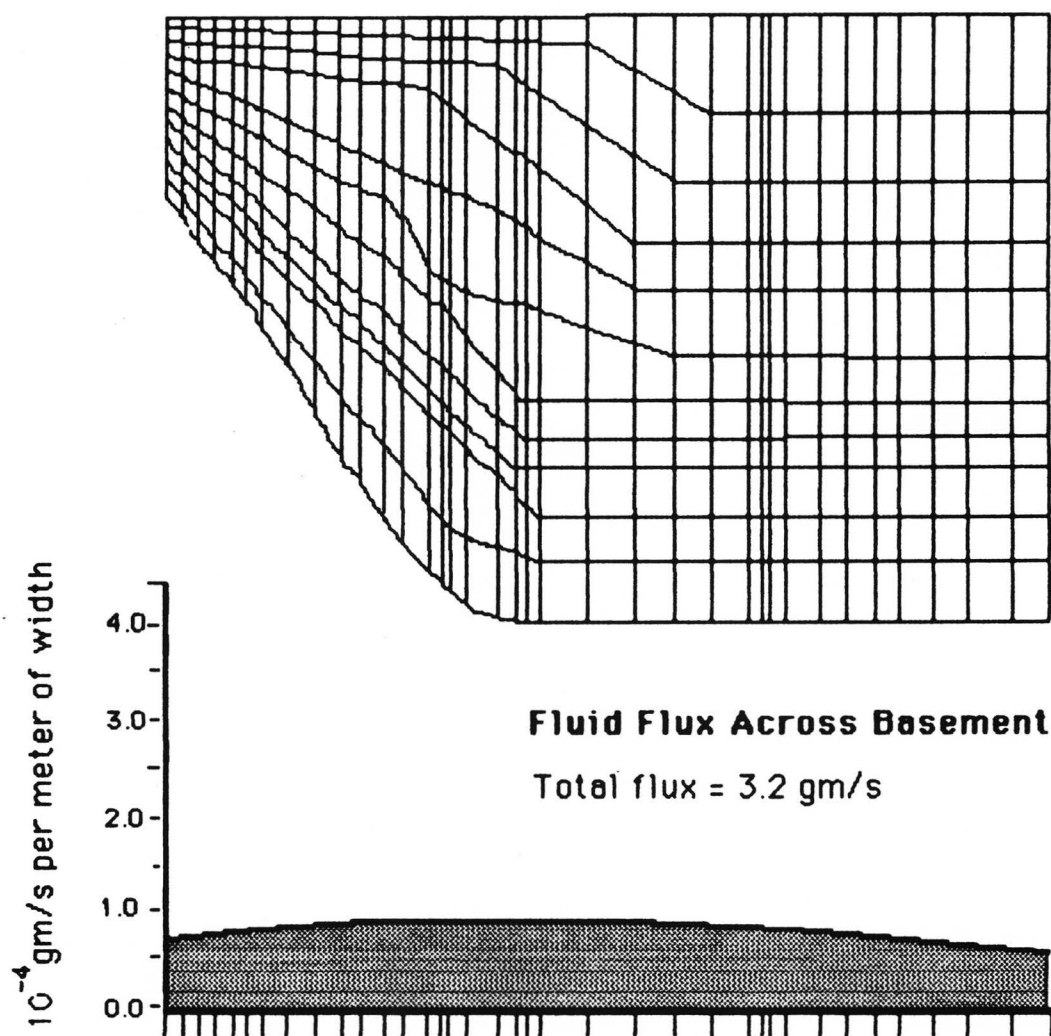


Figure 2.4: Fluid flux assigned to the basement nodes of the mesh for Simulation One.

permeabilities and thermal conductivities were optimized, remaining calibration discrepancies were addressed by conservatively incrementing the fluxes. This procedure was repeated until a "best fit" was obtained. The overall average was 8.91×10^{-8} kg/s for each meter across the section; the total flux for the entire section being 3.2×10^{-3} kg/s.

Permeability Codes. Figs. 2.5 through 2.8 show the coded mesh for horizontal and vertical permeability, porosity and thermal conductivity. The permeability fields were kept simple by assuming faults to have no influence. The geopressured zone was assigned the lowest permeabilities, with a rapid increase in permeability at the transition to normal pressures. The southeastern or right portion of the mesh corresponds to the sands of the Frio Formation and consequently higher permeabilities.

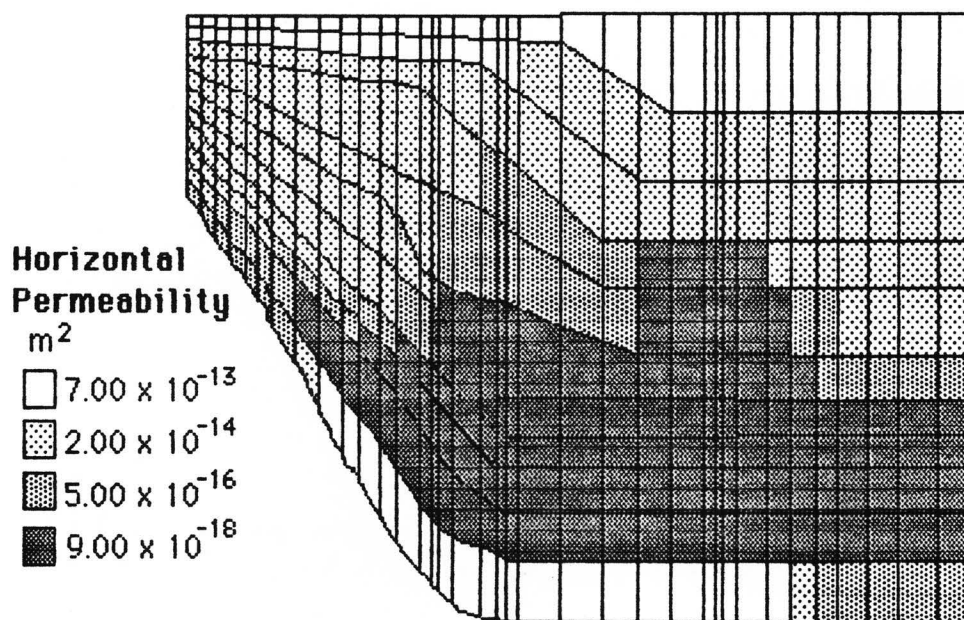


Figure 2.5

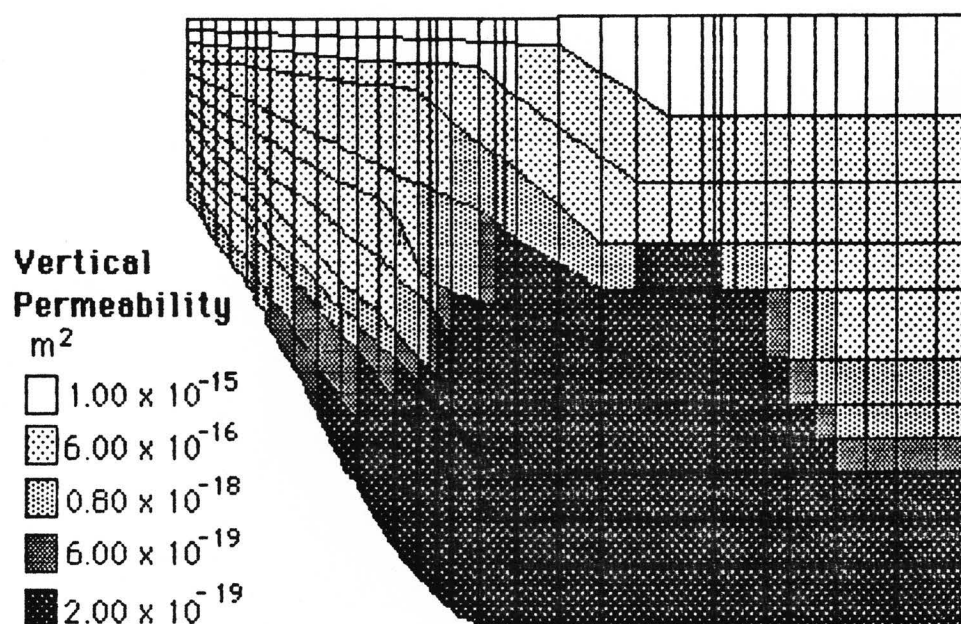


Figure 2.6

Porosity Code. The porosity mesh was also simply designed. Porosity decreases from a maximum of 35% at the surface to 5% at 20,000 feet (6096 m). Porosity is held constant at 21% across the pressure transition to geopressuring. Fig. 2.7 shows the coding for the porosity mesh.

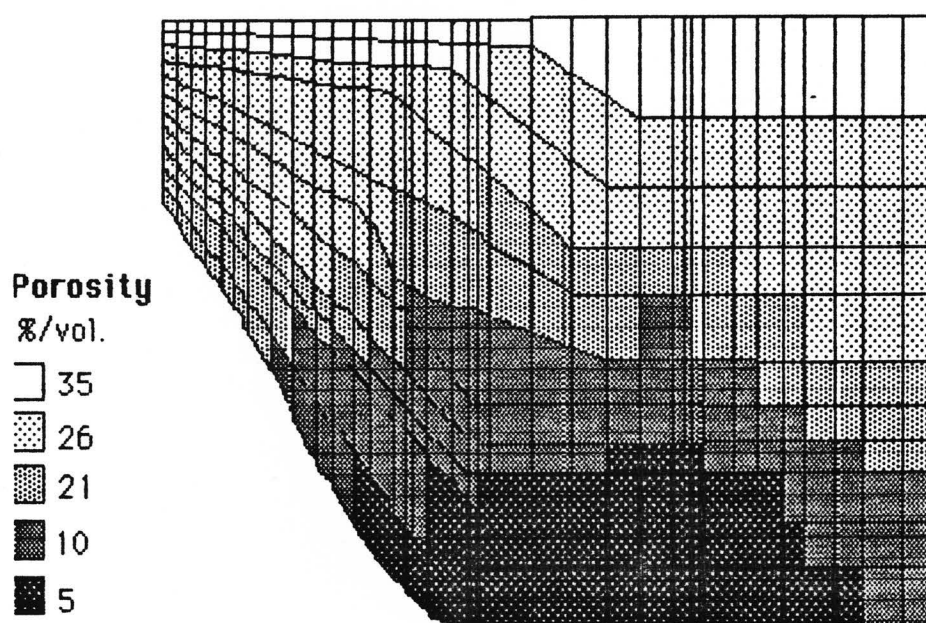


Figure 2.7

Thermal Conductivity Code. The Thermal Conductivity mesh is comparatively complex (Fig. 2.8), particularly in the portions above the top of geopressure, because the lithology is most variable there. Sandy units typically have higher conductivities than clays and shales (Reiter and Tovar, 1982). Thus, higher conductivity units were assigned in the Frio formation on the right portion of the mesh. The lowest conductivities were assigned to the shallow geopressed section because the least consolidated thickly bedded clays were presumed to occur there. The more unconsolidated the sediments are, the more poorly they conduct heat. An isolated high conductivity unit was placed in the deep geopressed Wilcox to correspond with a large sand body which is indicated on the type-section.

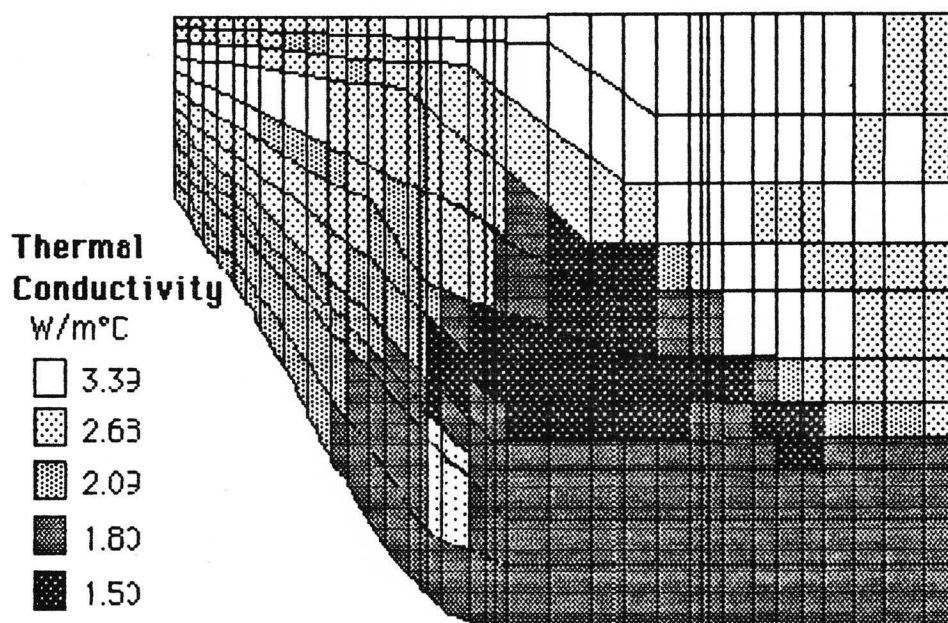


Figure 2.8

2.4.2 SIMULATION ONE-RESULTS

Fluid Flux. As mentioned above, the average fluid flux assigned across the geopressed section was 8.91×10^{-8} kg/ms for each meter across the section. An estimate for comparison was provided by Lynton Land (personnel communication, 1985). The estimate is based on the chloride concentration of Gulf Coast rivers and streams whose watersheds do not drain Pennsylvannian or Permian evaporites. Possible sources of chloride in surface waters are aerosols (mostly in the form of rain), halite from evaporites, and discharge of deep-basinal brines. The average value for river-borne chloride was 8 grams of chloride per square meter of land surface per year. Of this, only 1.2 grams can be accounted for by aerosols. The 6.8 gram deficit which remains is assumed to be due to the contribution from deep-basinal discharge.

Figuring the flux assigned in the model provides 2.81 kg or about 2.8 liters per square meter per year, this implies a chloride concentration in the upwelling deep-basinal fluids of 243 parts per thousand. This is well within the realm of halite saturation which is about 400 parts per thousand at 100°C, and is also very similar to concentrations reported from brines in the Edwards Formation of the basin. This estimate thus provides an indication that the flux rates chosen are quite reasonable.

On the other hand, a volumetric calculation yields a very different result. Using the same 2.8 kg/ms flow rate, an estimate of the initial height of a one meter square column of compacting sediment needed to produce this flow rate was obtained. Assuming a period of 40 million years and an initial porosity of 35%, a column of sediment 320,000 meters high would have to be compacted to 0% porosity to maintain this flow rate. Since the depth of the basin is probably no more than 13,000 meters, the flow rate appears from this calculation to be more than an order a magnitude too high, indicating that simple compaction is inadequate to account for the postulated flow rate. Bruce (1984) claims that the smectite-illite clay transition can supply as much as three times more fluid than compaction in the depth range of 7000-11000 feet (2134-3353 m). However, this is only for a 5% volume reduction from 14% to 9% porosity. Though potentially important, the transition alone does not account for all the fluid needed, nor does it account for fluids sourced from depths greater than 11000 feet (3353m).

Fluid Pressure. Fig. 2.9 shows the pressure contours generated by the model. The contours are meters of excess hydraulic head relative

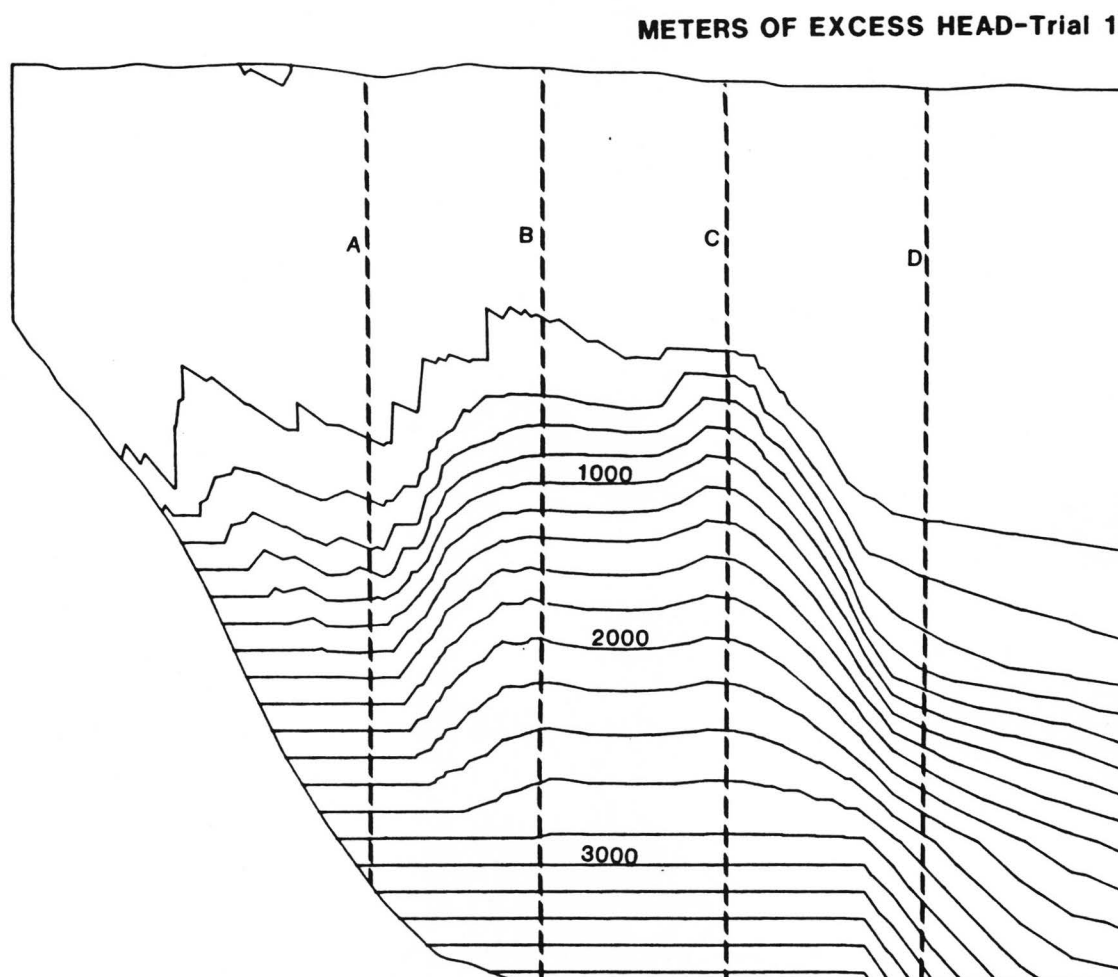


Figure 2.9: Hydraulic head generated by the model for Simulation One. Dashed lines show location of pressure profiles in Figure 2.10.

to the highest topographic elevation of the section. The contour interval is 200 meters. Since the contours are excess heads, they mark the geopressured zone. The first several contours should therefore correspond to the top of geopressing as indicated by the hatchured line in the cross section of Fig. 1.2. Fig. 2.10 shows pressure profiles taken along four transects of the section indicated on Fig. 2.9. Here the agreement with other data (shown on the figure) is not good. The pressures modeled at each of the four transects are within reason but are generally low compared to the generalized pressure trends, which are taken from Neogene deposits of South Texas (line 1) and recent to Miocene in Louisiana (line 2). But missing from the modeled pressures is the characteristic shape of the pressure profile, namely the inflection in the line caused by the transition from increasing pressure gradients at shallow geopressures to a constant pressure gradient at continued depths. The profiles are instead relatively linear, with little variation in the pressure gradient. The pressures therefore can not be considered well calibrated.

Temperature. Fig. 2.11 shows the isotherms produced by the model (dotted lines) superimposed on those from the BHT data (solid line). The gray shaded areas are the error ranges for the BHT data. These were taken from the contoured variances (from Kriging) shown in Figs. 1.6-1.10d. The match of the 200°C isotherm is quite good, but discrepancies increase at higher temperatures, and are especially large between the 350°C and 400°C isotherms. The largest discrepancies

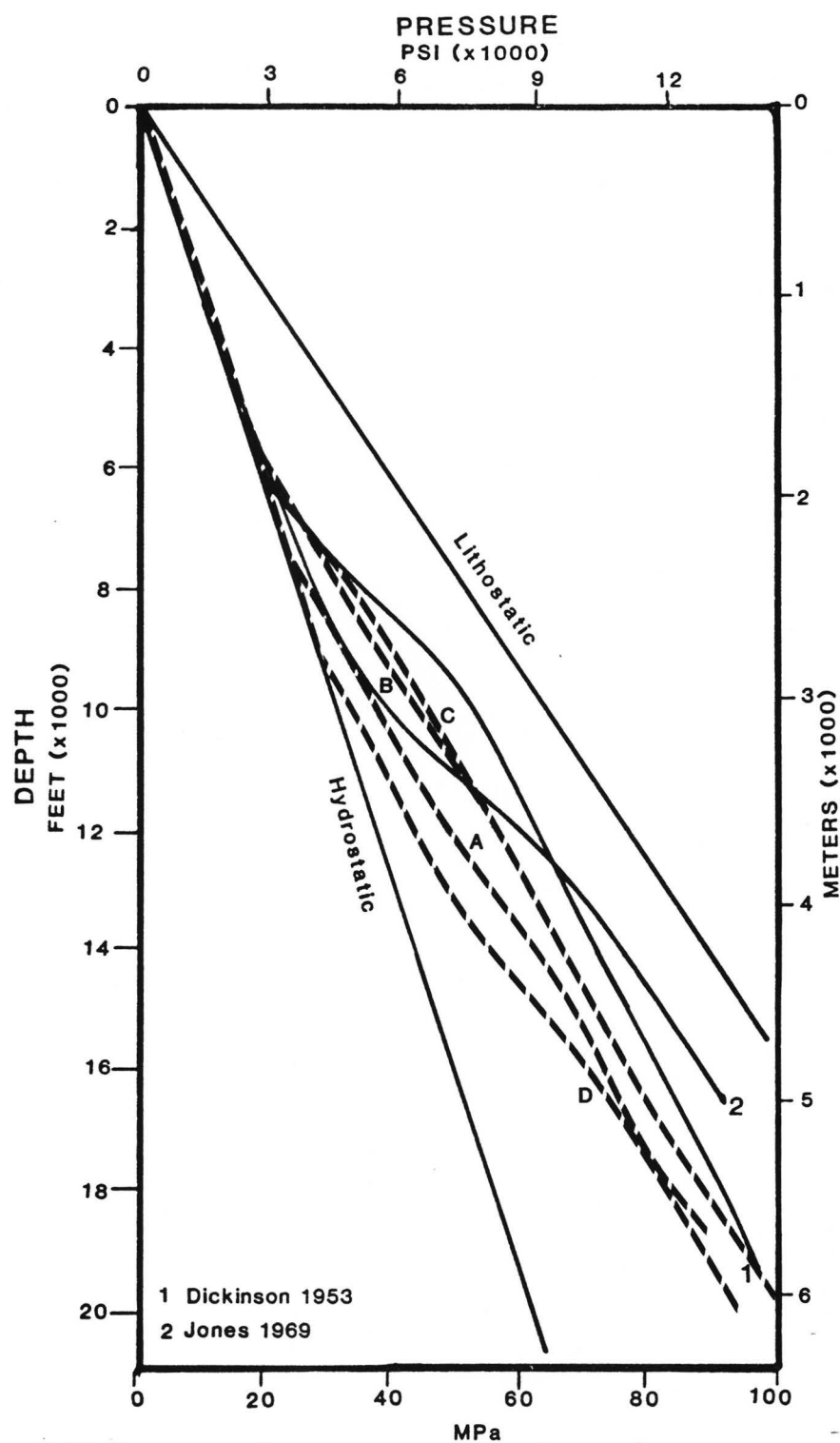


Figure 2.10: Pressure profiles for Simulation One. Dashed lines show modeled pressure profiles keyed in Figure 2.9. Solid lines 1 & 2 show generalized profiles for Gulf Coast.

occur at the Wilcox growth fault zone, where modeled temperatures are much lower than the data indicate.

Four temperature profiles were constructed from selected transects of Fig. 2.11 and these were superimposed on the temperature profiles from the corresponding subareas presented in Figs. 1.18, 1.14, 1.20, and 1.23. The advantage of this comparison is the ability to view the actual BHT measurements in conjunction with the modeled temperatures. Figs. 2.12 through 2.15 show these profiles. Once again, the matches are generally poor. In all but the group nine area, in fact, modeled temperature gradients are below average, whereas the data show that gradients in all areas are above average.

The Flow Regime. Fig. 2.16 is of the fluid velocity vector field. The arrows in the figure are vectors, meaning that their length is proportional to the fluid velocity at the point of the arrow's origin, and the direction of the arrow indicates the flow direction. The direction is corrected to account for the vertical exaggeration of the figure. The vectors are placed at element centers. Since many of the flow rates are extremely low, they appear only as the head of the arrow. Wide differences may still exist between these velocities, though not apparent from the figure. Appendix F lists the flow rates for each element. Readily apparent from the figure is the break from geopressured to meteroic regimes. The meterioc system is characterized by much higher flow rates as well as more variable, topographically controlled directions.

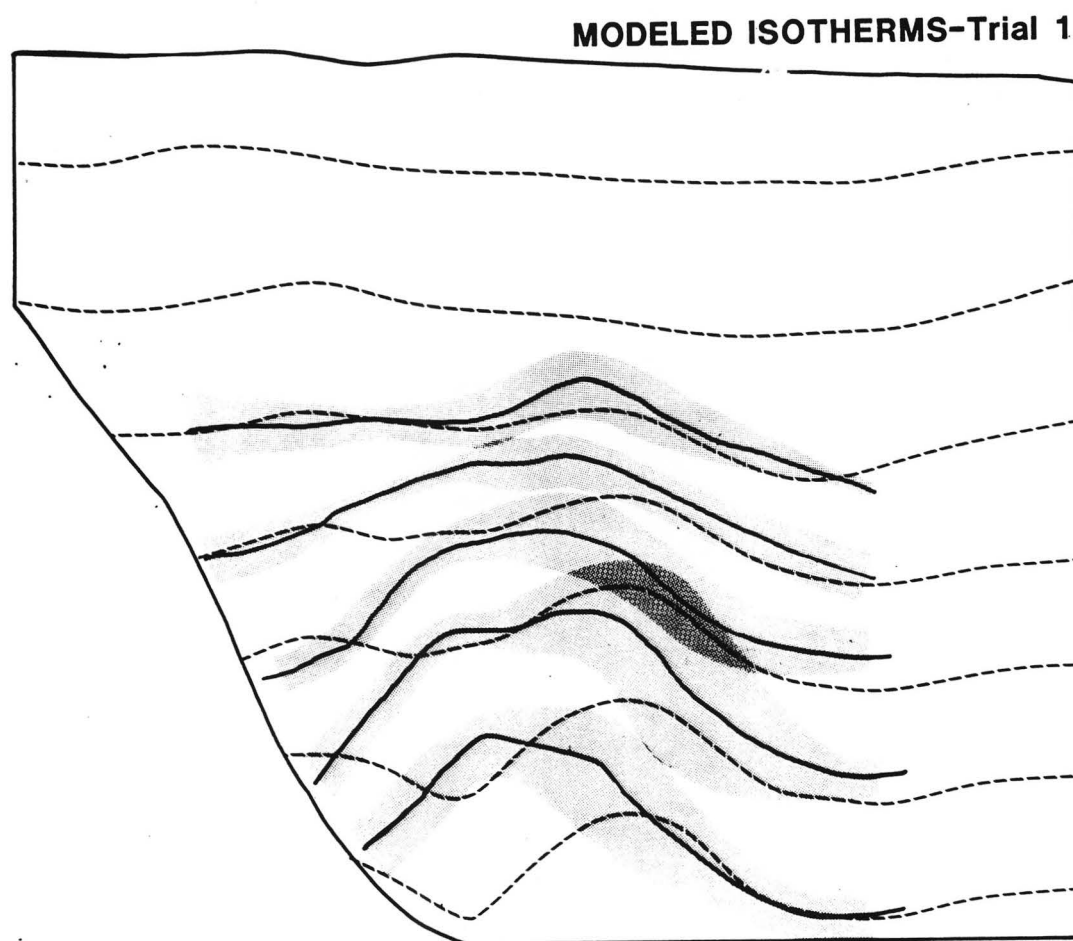


Figure 2.11: Modeled isotherms for Simulation One shown by the dashed lines; solid lines indicate isotherms from kriged BHT data with the grey shaded areas showing their standard deviations.

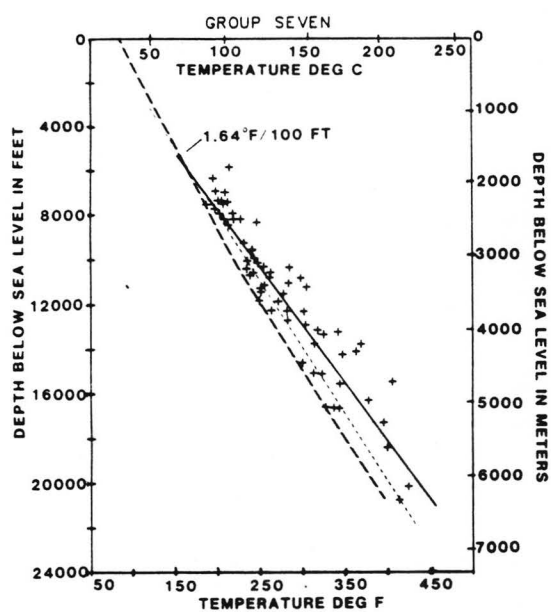


Figure 2.12

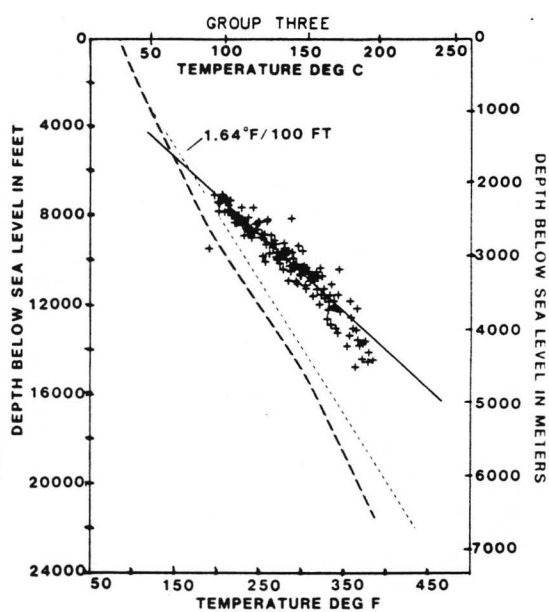


Figure 2.13

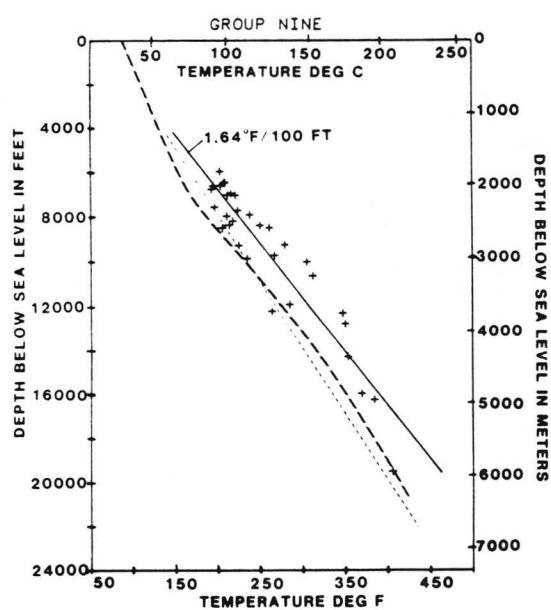


Figure 2.14

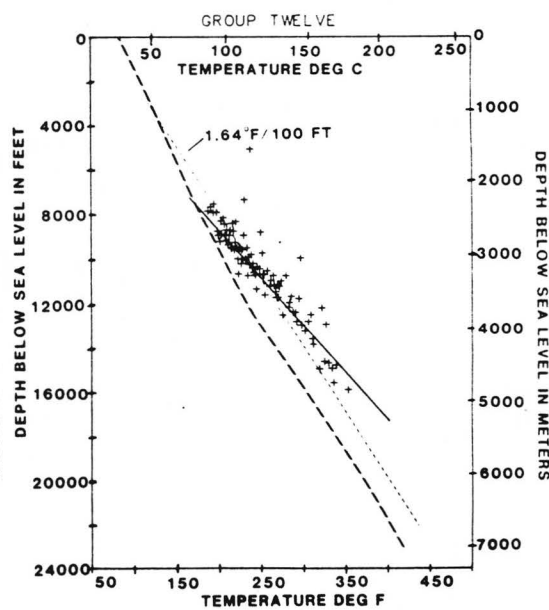


Figure 2.15

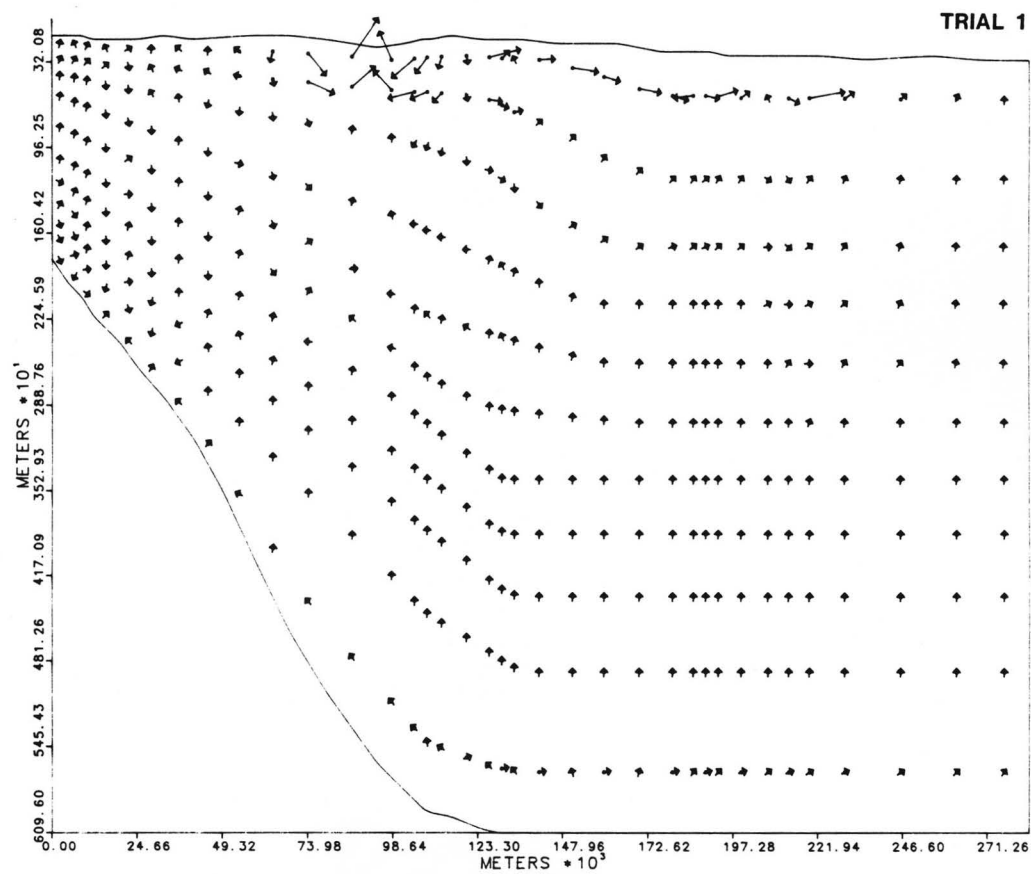


Figure 2.16: Simulation One flow velocity vectors.

2.4.3 SIMULATION TWO - CONCENTRATED FLOW

Flux Nodes. The second simulation employed areas of concentrated vertical flow to model the basin. The same overall flux used in the first simulation was maintained in Simulation Two, but it was redistributed in such a way as to concentrate most of the flux along the Wilcox growth-fault zone. Fig. 2.17 shows this distribution.

Permeability Codes. Figs. 2.18 and 2.19 show the permeability codes. These are much more complex than in the previous simulation. Relatively high horizontal permeabilities were assigned to the basal row of elements to avoid pressure plumes around each node, but these values were much lower than those used previously in order to avoid widely dispersed fluxes across the basement.(Fig. 2.4). Elsewhere, horizontal permeability was decreased with depth, and in response to the new flux conditions, the horizontal permeability field was refined in the upper geopressured section. Specifically, values were increased in order to disperse the concentrated flux below the pressure transition, a condition necessary to calibrate the isotherms. The justification for this alteration is that although vertical permeability in the marine shales is certainly very restricted, horizontal permeability can increase near the top of geopressure because of high porosity and a relatively unconsolidated, laminar structure.

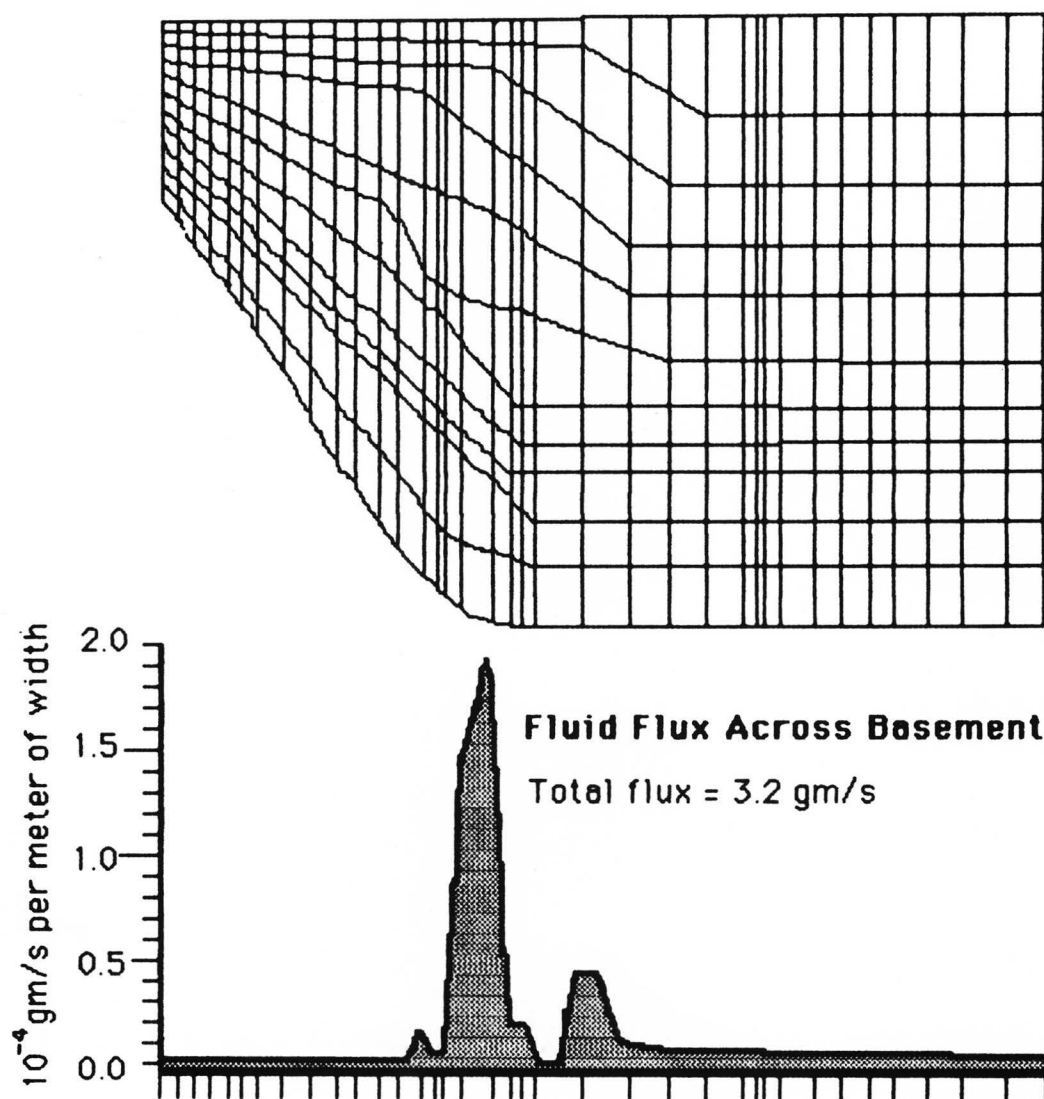


Figure 2.17: Fluid flux assigned to the basement nodes of the mesh for Simulation Two.

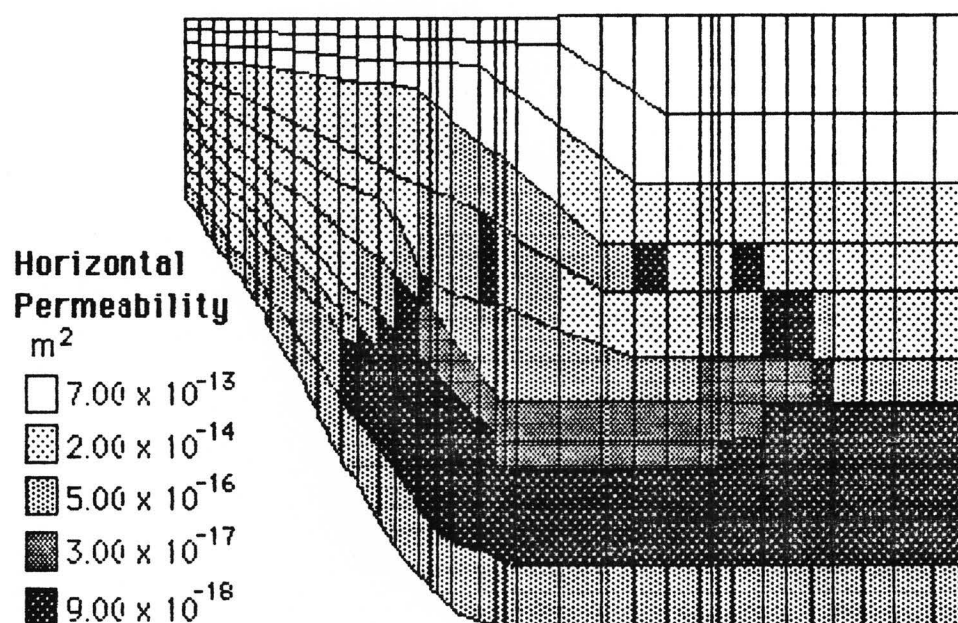


Figure 2.18

The vertical permeability mesh has a different configuration. Its values were assigned as a function of depth except where vertical conduits for flow were hypothesized. Two portions were modeled as such, the deeper parts of the Wilcox growth faults, and the geopressed portion of faults in the Catahoula Fm.

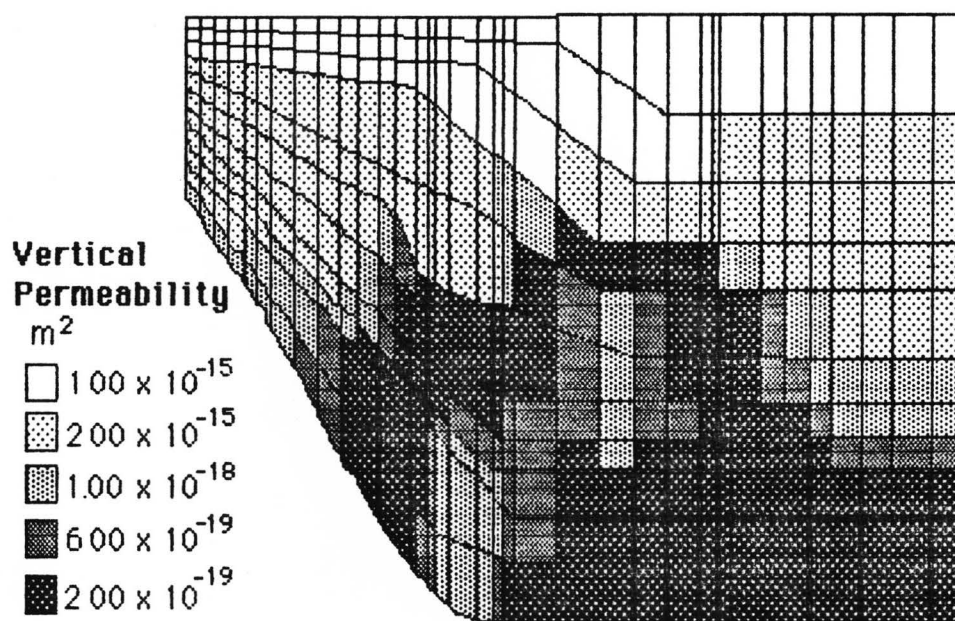


Figure 2.19

Porosity & Thermal Conductivity Codes. The porosity field is unchanged from the first simulation (Fig 2.7). Slight modifications were made to the thermal-conductivity mesh (Fig 2.20). The conductive sand body in the deep basin was removed, and the elements of least conductivity were redistributed, adding more in the Wilcox growth-fault zone. These refinements were needed to calibrate the model given the new fluid flux conditions, being careful to keep values within ranges supported by the available data.

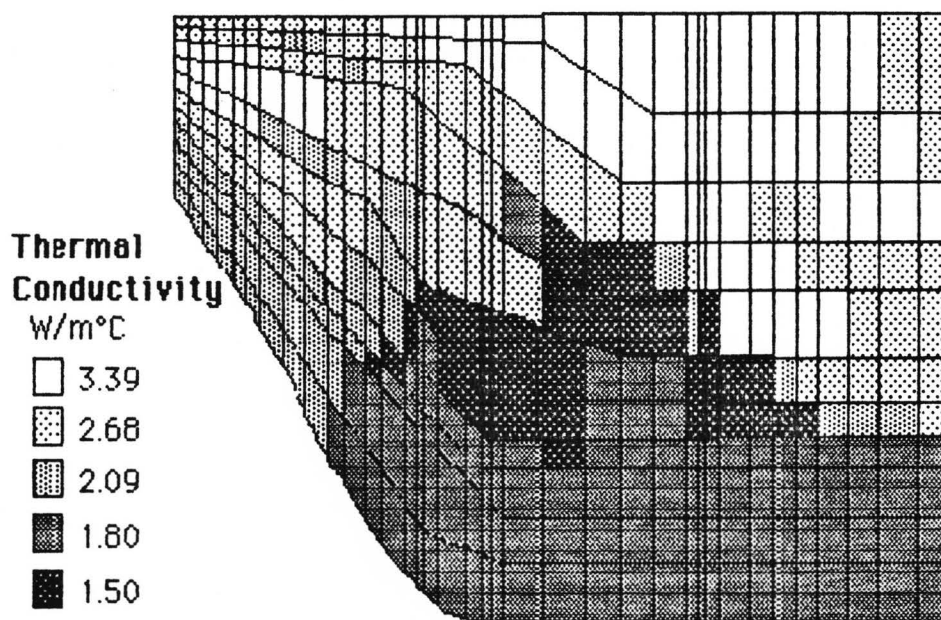


Figure 2.20

2.4.4 SIMULATION TWO-RESULTS.

Fluid Pressure. Fig. 2.21 shows the pressure contours for Simulation Two. Unlike the previous simulation, a plume indicative of concentrated upwelling is apparent in this figure. The pressure vs. depth profiles derived from this diagram and given in Fig. 2.22 show much better agreement with the two other data sources shown. Line B from the growth fault zone shows near perfect agreement with Dickinson's data, reproducing the characteristic inflection point. The fact that the actual top of geopressure (indicated by the first deviation away from the hydrostatic line) does not match is not significant because Dickinson's data is from a different location, and the depth to the top of geopressuring is variable across the basin. Line D,

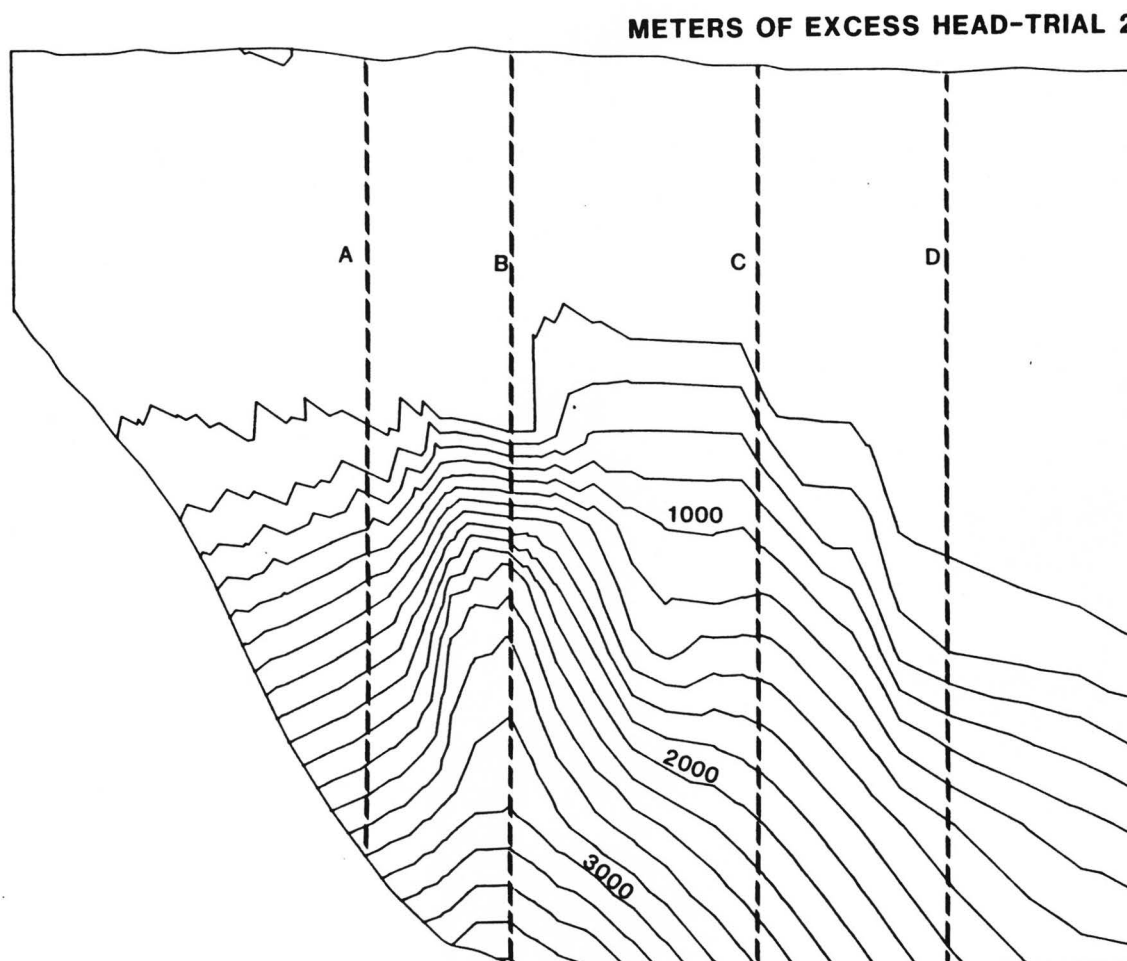


Figure 2.21: Hydraulic head generated by the model for Simulation Two. Dashed lines show location of pressure profiles in Figure 2.22.

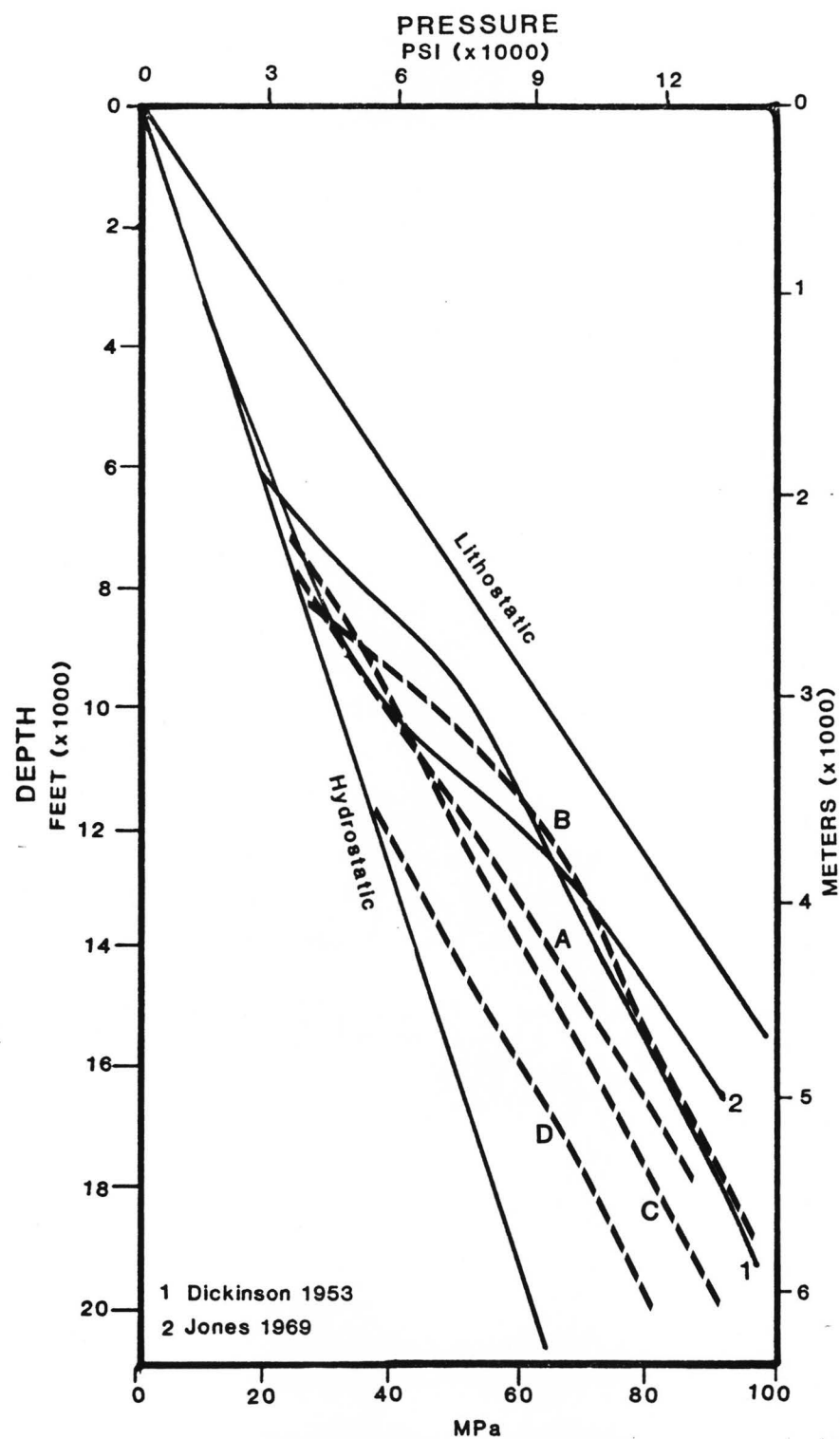


Figure 2.22: Pressure profiles for Simulation Two. Dashed lines show modeled pressure- profiles keyed in Figure 2.21. Solid lines 1 & 2 show generalized profiles for Gulf Coast.

corresponding to the Frio Formation, shows lower overpressures, as would be expected in an area of higher sand content and permeability.

Temperature. The temperature contours modeled in Simulation Two are shown in Fig. 2.23. The match with the data is better here as well. The modeled isotherms stay within the error margins in most places. Some discrepancies appear along the 350 and 400°C isotherms just southeast (right) of the Wilcox growth-fault zone, where temperatures are once again somewhat cool relative to the data. This trend is also evident in the temperature-vs.-depth profiles of Figs. 2.24 through 2.27, where the modeled temperatures fall consistently below the general trend indicated by the BHT measurements. No manner of fluid flux could lessen these discrepancies without disrupting the rest of the cross-section.

The Flow Regime. Fig. 2.28 differs little in appearance from Fig. 2.16. In each, the direction and magnitude of fluid movement appear quite similar. However, as previously mentioned, the flow rates are so low in the geopressured zone that differences within this zone do not show on the diagram. Appendix F lists the velocities from this figure. The highest values for the meteoric zone are on the order of 10^{-8} m/s. In the geopressured zone where fluids are concentrated, velocities are about 10^{-11} m/s, and drop off elsewhere in the geopressured zone to as low as 10^{-14} m/s.

2.45 SIMULATION THREE - MIXED CONDITIONS

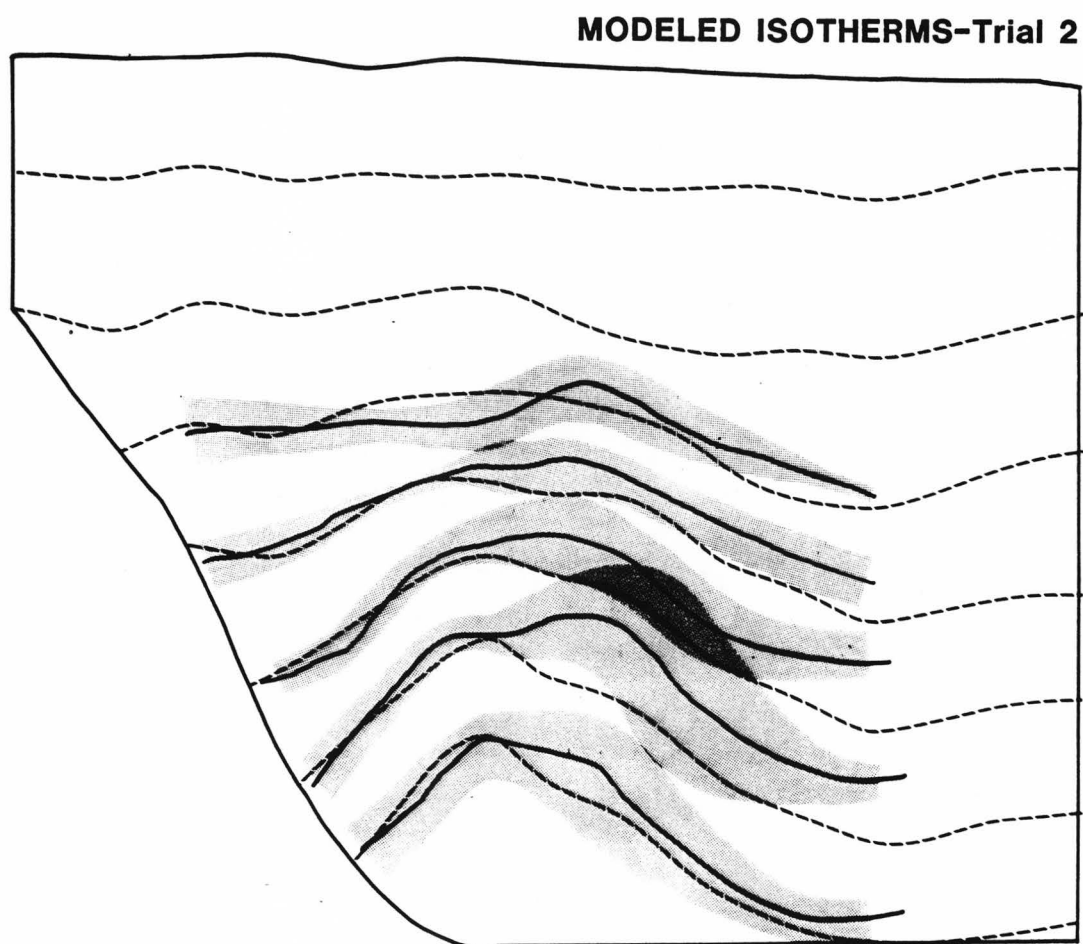


Figure 2.23: Modeled isotherms for Simulation Two shown by the dashed lines; solid lines indicate isotherms from kriged BHT data with the grey shaded areas showing their standard deviations.

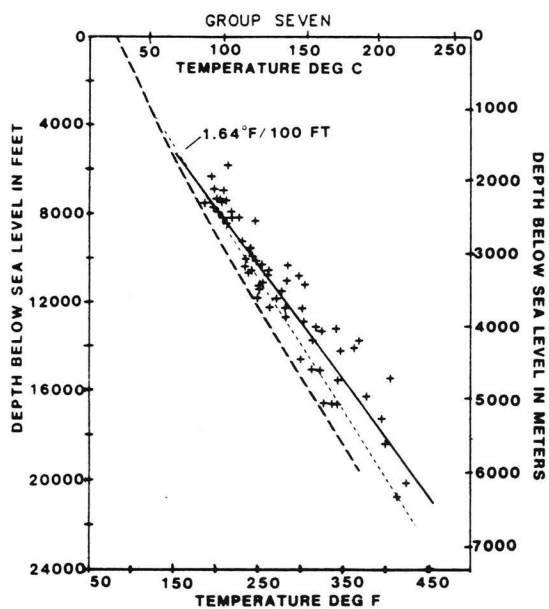


Figure 2.24

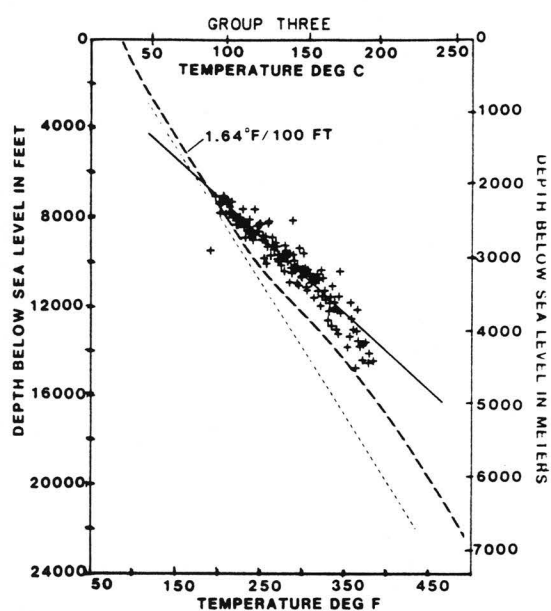


Figure 2.25

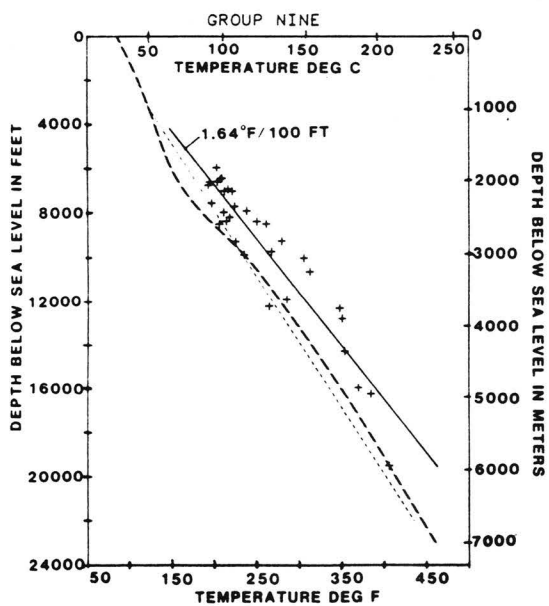


Figure 2.26

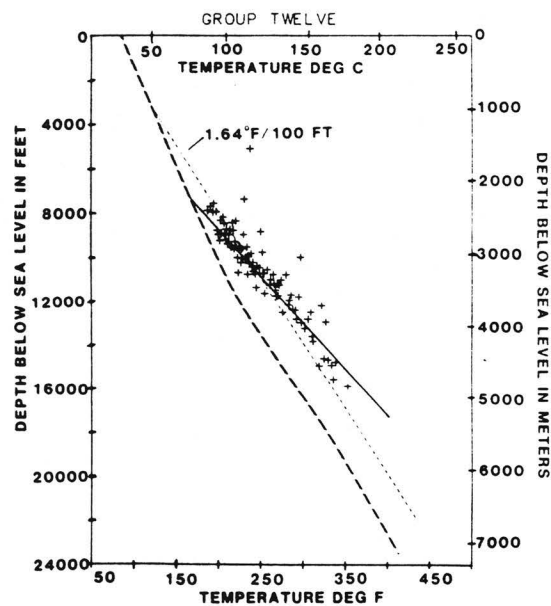


Figure 2.27

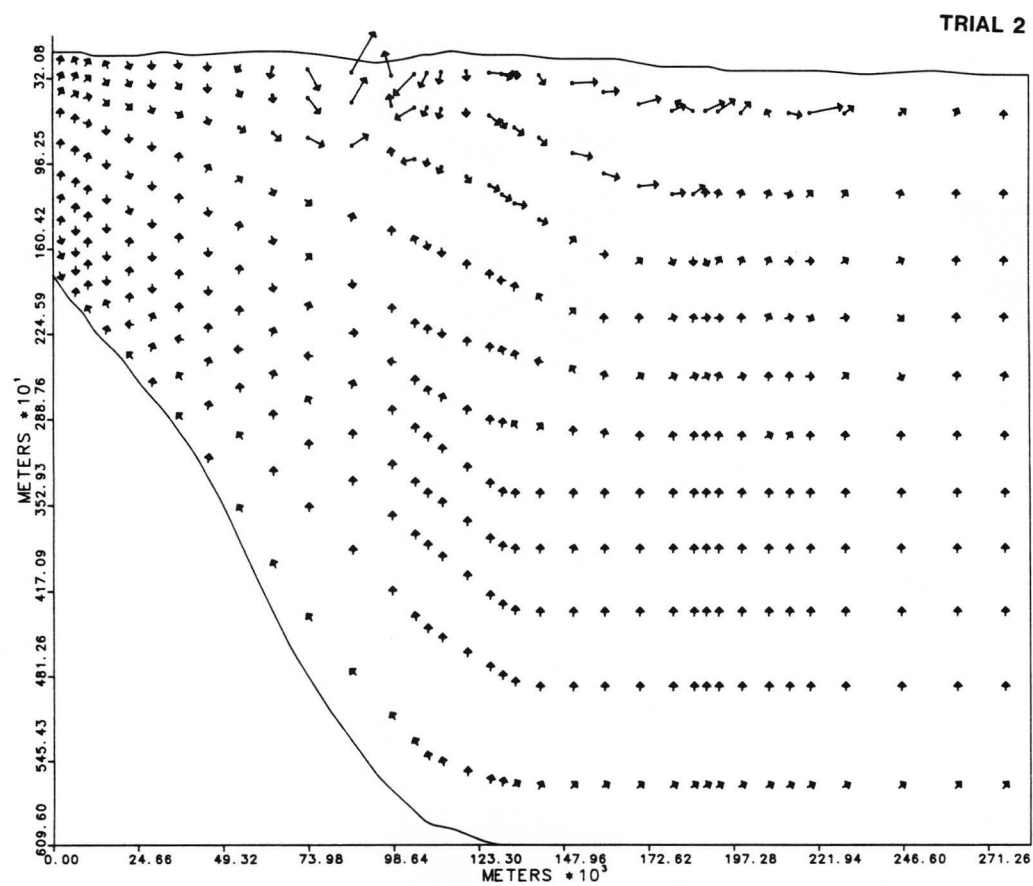


Figure 2.28: Simulation Two flow velocity vectors.

Simulation Three evolved from Simulation Two after successive refinements. The suite of code maps are shown in figs. 2.29 through 2.32. Porosity and fluid fluxes are essentially unchanged and the permeabilities are also very similar, excepting the addition of a speculative zone of high thermal and hydraulic conductivity located southeast (right) of the Wilcox growth-faults, which was added to simulate the isotherms more closely than in Simulation Two. The differences in the code maps between Simulation Two and Three are centered around the addition of this body. In conjunction, the depth of influence of the Carrizo-associated fault was decreased as indicated in Fig. 2.30.

2.4.6 SIMULATION THREE - RESULTS

Fluid Pressure. The fluid pressures are quite similar in this simulation to those from Simulation Two as illustrated in Fig. 2.33. The pressure profiles (Fig. 2.34) are consequently also quite similar and agree well with the available data.

Temperature. Simulation Three produced temperature isotherms which correspond very closely to the data, as seen in Fig. 2.35. The major discrepancies which appeared in Fig. 2.23 from Simulation Two were largely eliminated in this simulation. An even better indication of the success of this simulation are the Temperature-vs.-Depth profiles in Figs. 2.36 through 2.39. The correlation of the model to the data is nearly perfect for all four profiles.

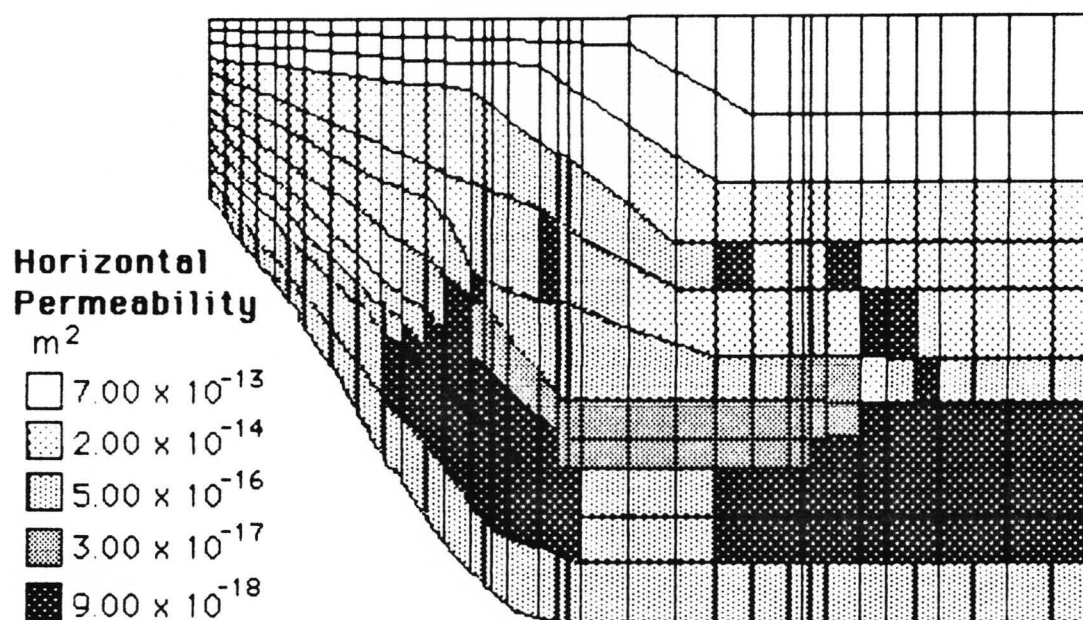


Figure 2.29

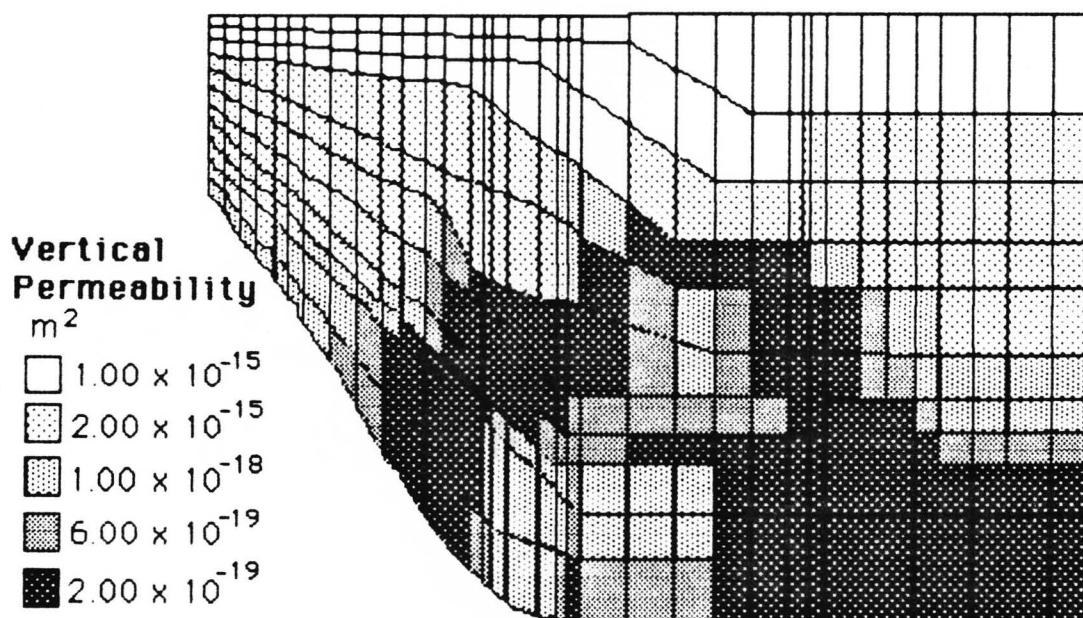


Figure 2.30

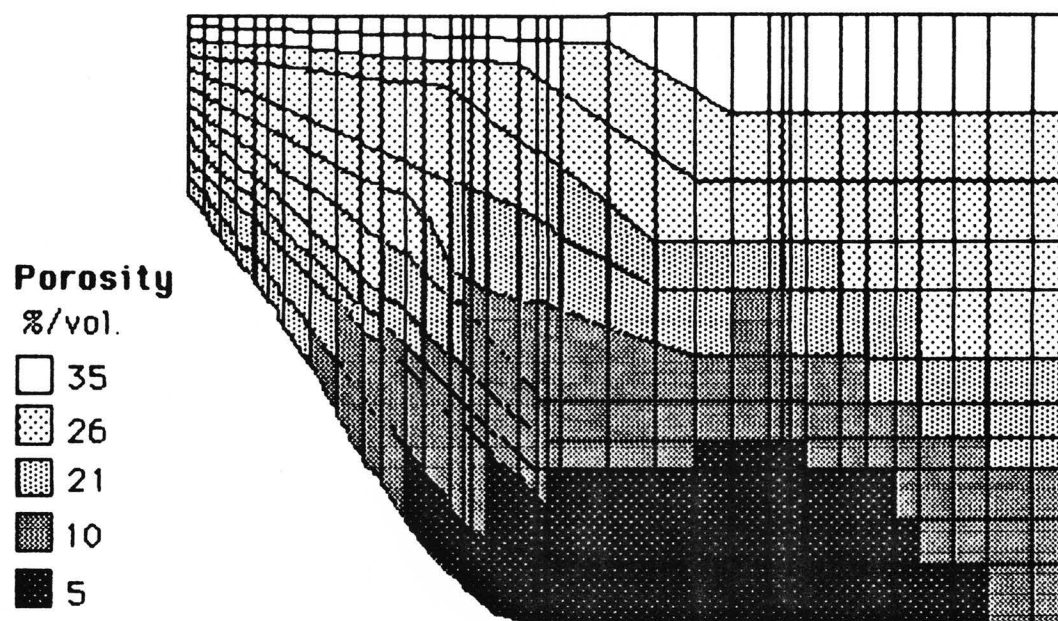


Figure 2.31

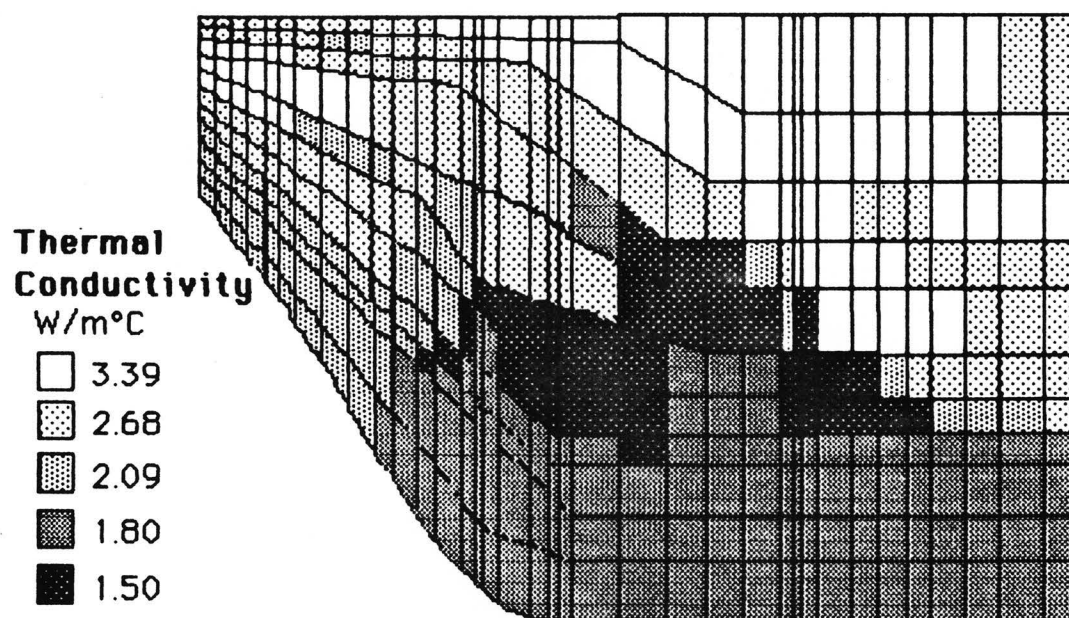


Figure 2.32

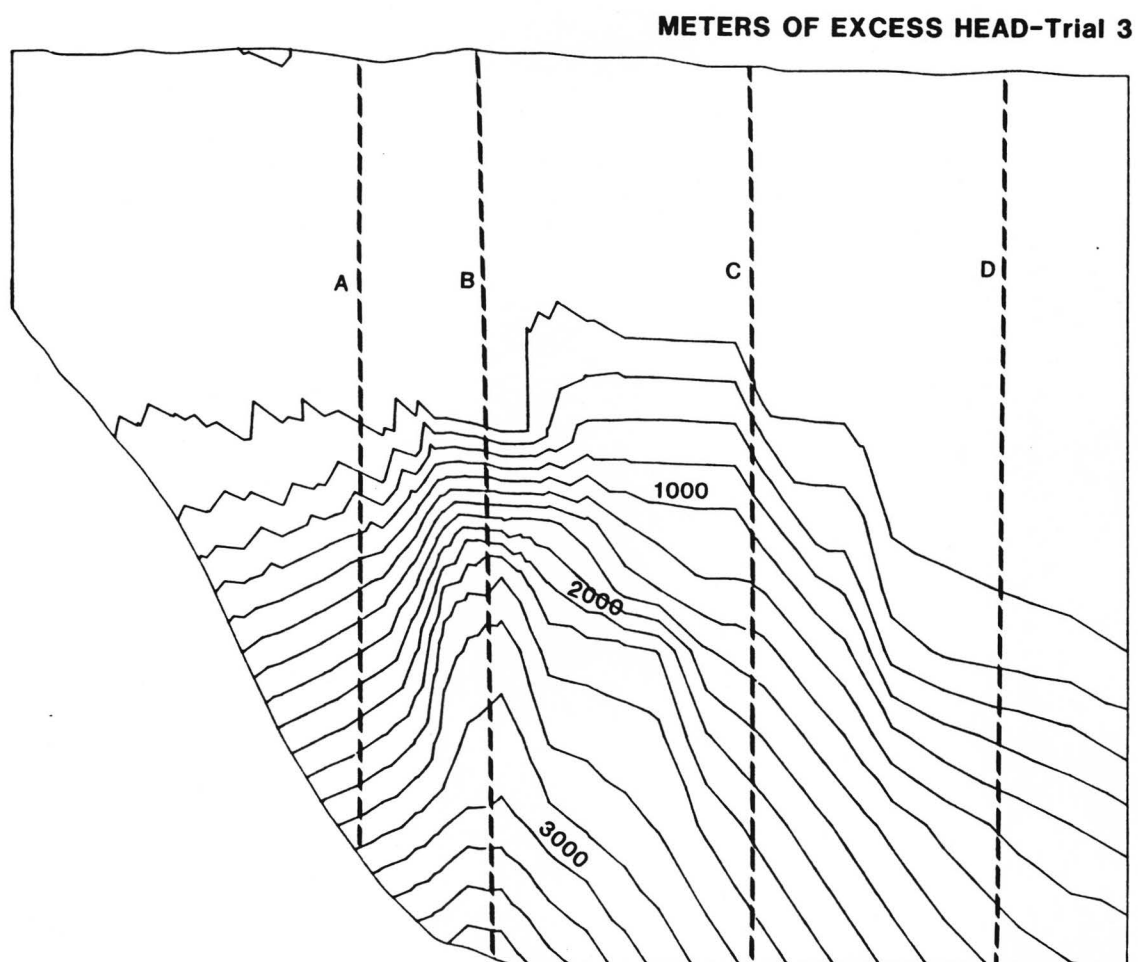


Figure 2.33: Hydraulic head generated by the model for Simulation Three. Dashed lines show location of pressure profiles in Figure 2.34.

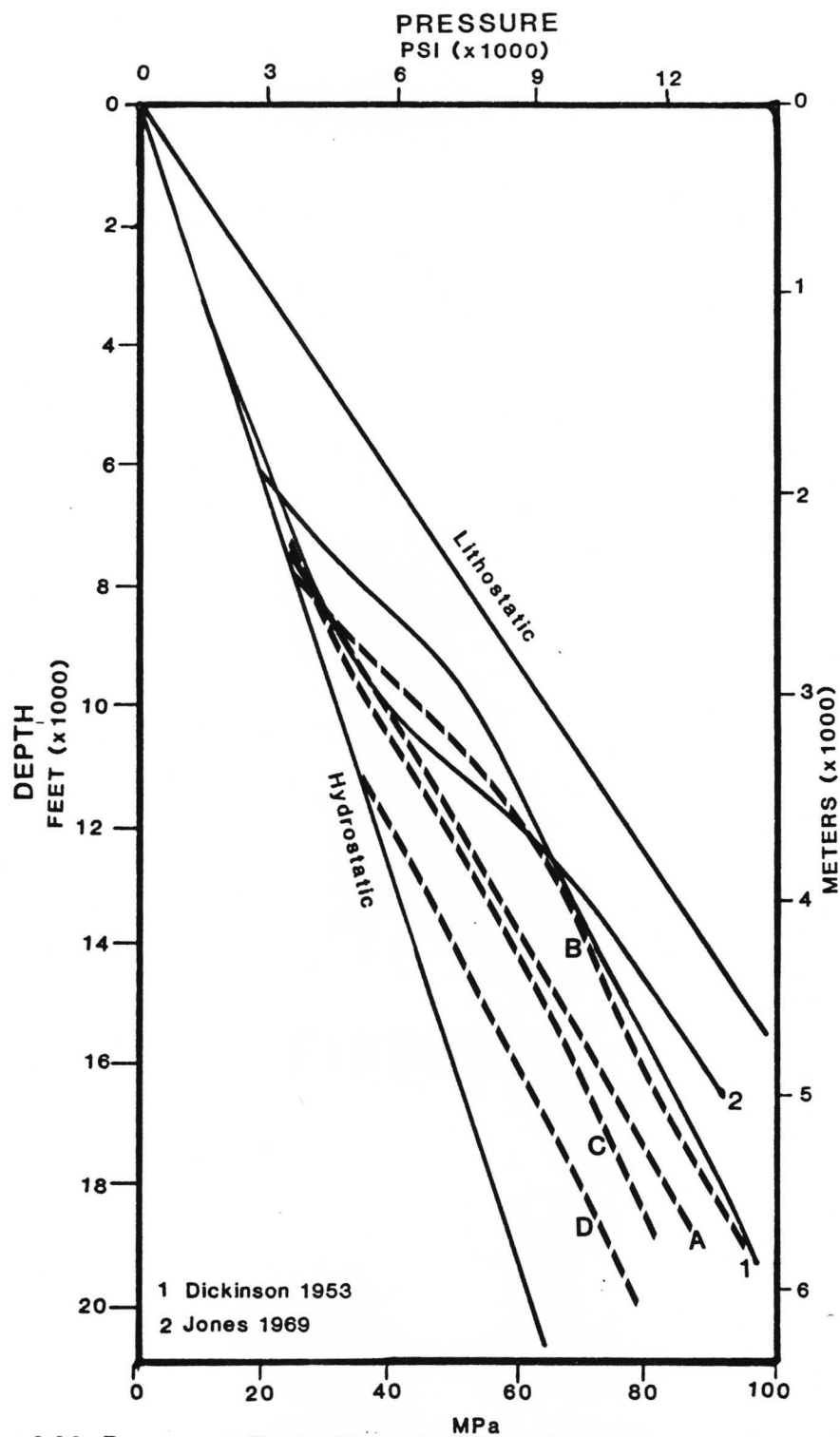


Figure 2.34: Pressure profiles for Simulation Three. Dashed lines show modeled pressure- profiles keyed in Figure 2.33. Solid lines 1 & 2 show generalized profiles for Gulf Coast.

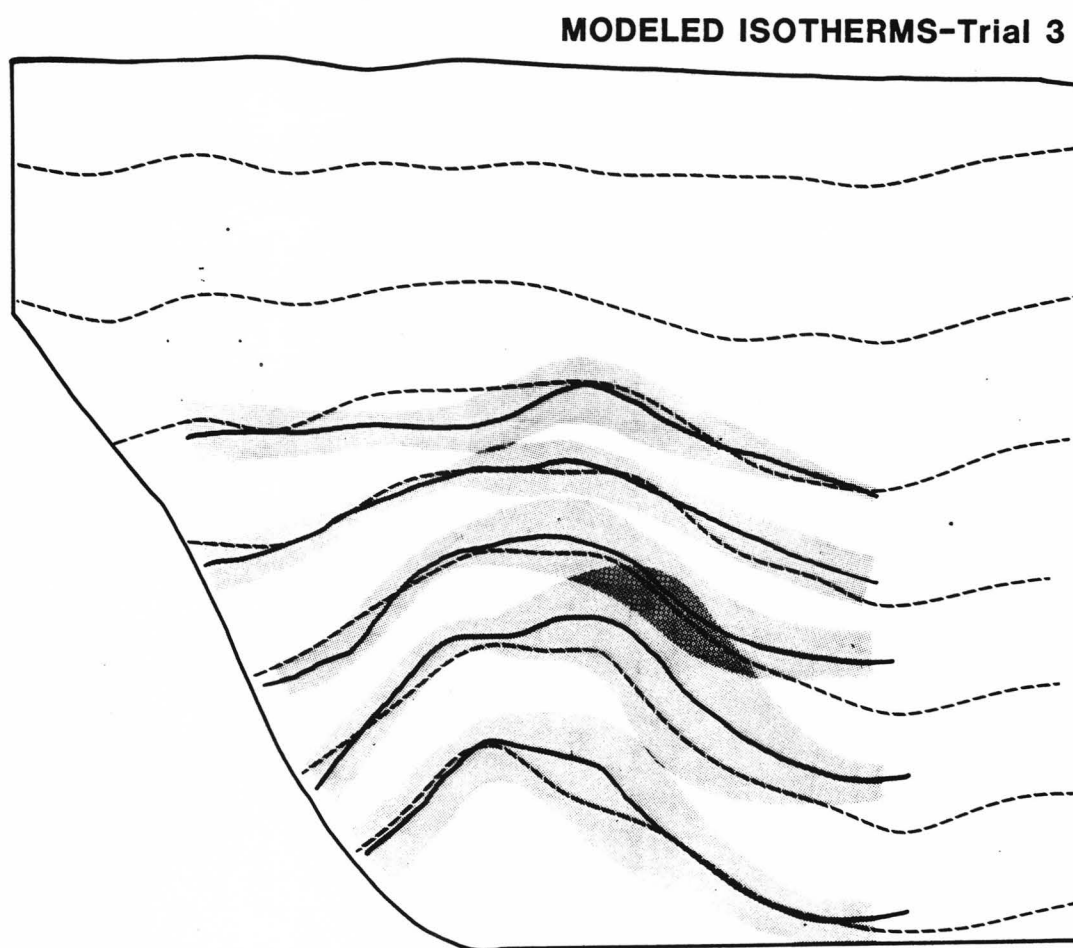


Figure 2.35: Modeled isotherms for Simulation Three shown by the dashed lines; solid lines indicate isotherms from kriged BHT data with the grey shaded areas showing their standard deviations.

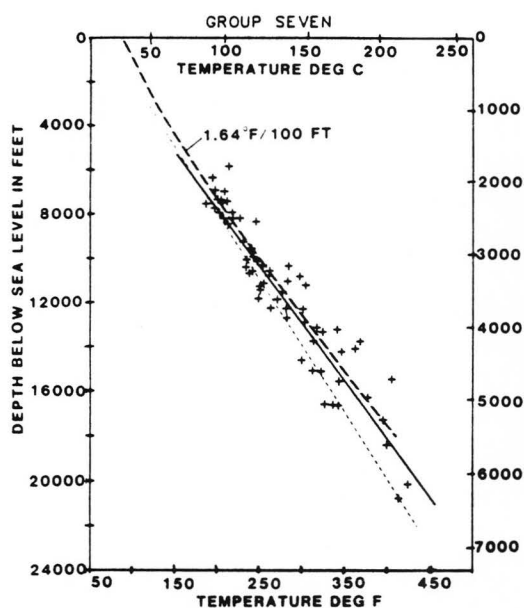


Figure 2.36

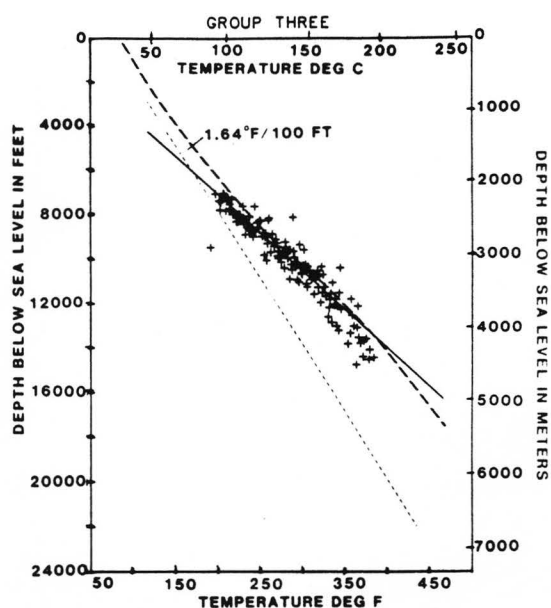


Figure 2.37

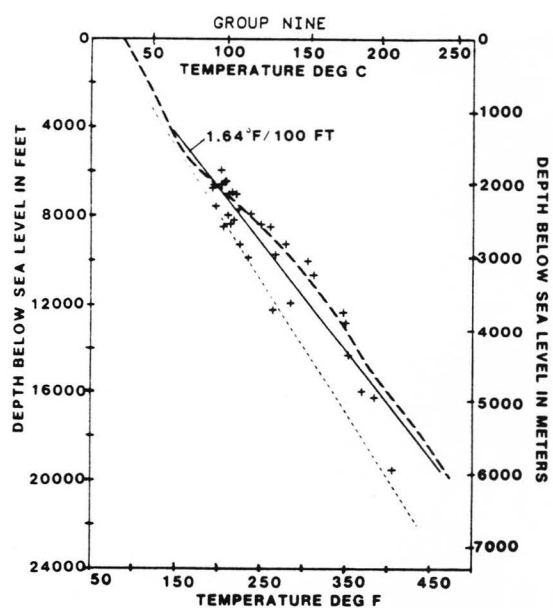


Figure 2.38

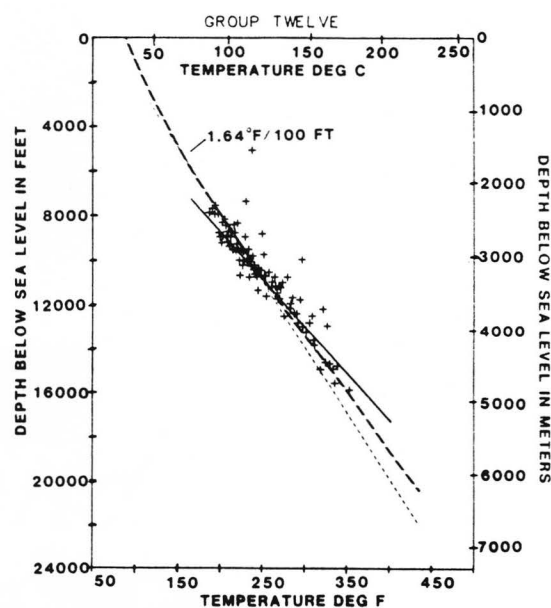


Figure 2.39

The Flow Regime. Once again, the fluid velocity vector plot (Fig. 2.40) is nearly identical to the other simulations, and is especially similar to simulation two, as would be expected, since conditions governing the basic flow regime were essentially unchanged between simulations Two and Three. Appendix F lists the velocities for each element.

2.5 DISCUSSION

As a result of the three simulations, I conclude that dispersed advection coupled with geopressure-dominated conduction is an inappropriate model for the Gulf Coast Basin of South Texas. Simulation One tested this hypothesis and failed to simulate either the basinal temperatures or the pressure data. Markedly better results were obtained in Simulation Two where concentrated flow was introduced in the growth-faulted zone associated with the Wilcox Group. The third simulation evolved from refinements to Simulation Two and achieved a close match to the calibrating criteria. The major adjustment was the addition of a hypothetical area of high thermal and hydraulic conductivity placed just coastward of the Wilcox growth faults.

2.5.1 IMPLICATIONS OF THE MODEL

Although the digital model cannot provide a unique solution, the simulations do provide insights into the operative geological processes.

The flux of fluids were all assigned along the basal boundary at a depth of 20,000 feet (6096 m) because the sensitivity analysis showed

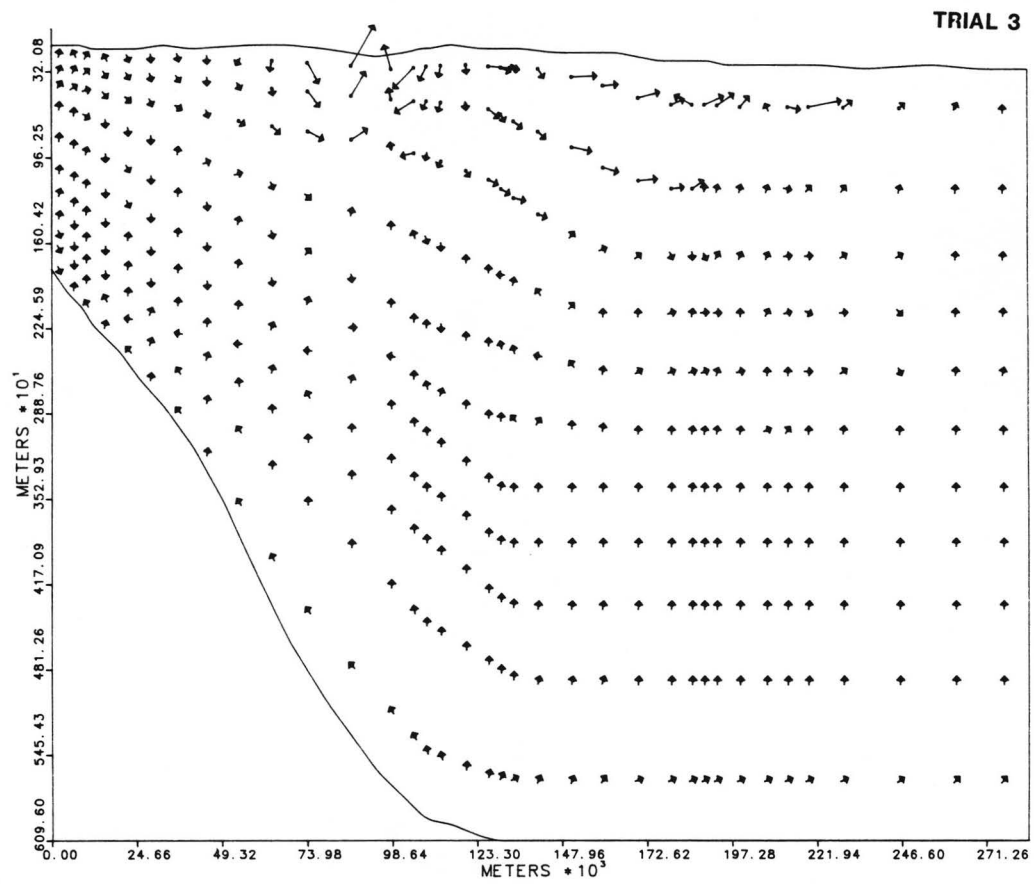


Figure 2.40: Simulation Three flow velocity vectors.

that this was the best placement of the flux nodes. Compactional fluids should be supplied diffusely throughout most of the section, predominantly by consolidating clays and shales, which are abundant above as well as below the basal boundary. The model, however, implies that most compactional fluids are from sediments below the basal boundary at 20,000 feet (6096 m), contradicting a study by Magara (1976) which places the maximum cumulative volume of expelled fluids at about 11,000 feet (3353 m). (Magara calculated this assuming fluid release resulting only from the steady consolidation of a typical Gulf Coast sediment-pile.) Furthermore, not only were the fluxes sourced from the basal boundary, but in Simulations Two and Three, the fluxes were also concentrated along the boundary under the growth faults, implying that not only are the fluids supplied from great depths, but that they have been focussed at these depths.

The treatment of fault permeabilities opens other avenues of speculation. The deepest portion of the Wilcox growth fault zone was assigned higher vertical permeability ($1.0 \times 10^{-18} \text{ m}^2$) than the surrounding rock. At intermediate depths still within geopressure, the higher permeabilities were discontinued. In a test simulation without this termination, fluids escaped the geopressured zone, leaving a large temperature plume penetrating the meteoric zone, with little geopressuring left at depth. The configuration presented thus supports the premise that faults leak fluid, but do not do so along their entire length. How faults might become sealed at intermediate depths is open to speculation. Perhaps the shales become more consolidated and more

brittle at depth, healing the faults more easily at intermediate depths, but leaving the faults open further below. There might also be greater fault movement in the deep basin. Finally, there is an unassessed potential for hydraulic fracturing. The fault zone may be particularly susceptible to microfracturing because of the existing faults and associated structural flaws. Furthermore, since the occurrence of microfracturing is dependent on high fluid pressure, suitable conditions may exist at greater depths. In short, it is not difficult to formulate several ways in which the configuration of the growth fault zone modeled in Simulation Two and Three might be produced in nature.

Shallower faults in the Catahoula Formation were modeled in a similar way, with higher permeabilities terminated just below the top of geopressure. Uranium deposits in the meteoric portions of these faults, similar to those studied by Galloway (1982), indicate that pulses of fluid have periodically moved up these faults. Therefore, probably during periods of fault movement, higher permeability did extend up through the top of geopressure. This was apparently a temporary situation. The faults may have become plugged near the geopressure boundary, as configured. Causes could be rapid cementation or the fast healing of poorly consolidated sediments, especially aided by high differential stresses at these shallow depths. W. E. Galloway (personal communication, 1985) states that there is little indication, at least at shallow depths, of extensive plugging due to cementation.

The most puzzling aspect of Simulation Three is the speculation of a zone of high hydraulic and thermal conductivity. Despite

that fact that the temperature contours most closely linked to this hypothesis also exhibit the largest standard deviations, comparison with the data points plotted on the temperature-vs-depth profiles argue persuasively that the temperature trends are real and that Simulation Three offers a possible solution to the temperature distribution. This region, if it exists, could best be explained by the presence of salt diapirs, the salt itself being virtually impermeable but of high thermal conductivity, while the diapirism can lead to overlying faulting and good vertical permeability. S. Seni and W. Mullican (personal communication, 1985) confirm that salt is present in this portion of the basin at the depths corresponding to the hypothetical zone. This salt may be in the form of salt pillows at various stages of maturity, and perhaps also diapiric structures. Several shallower salt structures in the form of salt domes have been identified in the region, including the Moca Dome, the Palangana Dome, and the Gyp Hill Dome, although W. E. Galloway (personal communication, 1985) suggests that the overall presence of salt in this region may not be very extensive.

Thus far, we have assumed that the model is useful in providing insights into the processes within the study area. However, unknowns preclude a unique solution. One of these is the problem of fluid flux. The estimate using chloride concentrations shows that the fluxes modeled are reasonable for present day conditions. However, these values are not volumetrically possible if they have been maintained over the history of the basin. They are in fact, more than an order of magnitude too large for this. Other flow estimates have been made based on various

observations such as the volume of cementation in the basin (Land, 1984, Table 2). These estimates have consistently been above flow rates that are volumetrically reasonable. My simulation adds one more high estimate to the pile.

A host of alternatives can be invoked to explain this discrepancy. One possibility is that the temperature data are seriously skewed. This is unlikely because the data base is large and the consistency of the data is good. Recall from Part One that the correlation coefficients for the thermal profiles were all above 90%. The steady-state assumption of the model is another possible contributing source of error. The flow rates may indeed be correct, but are transient, and have not persisted for a significant geologic period of time. Finally, error may have been introduced in the method employed to simulate compactional fluid flux. In reality, the volume of fluid moving out of the basin as a result of consolidation must be about equal to the volume loss of the consolidating sediments. Since the model does not compact, this inverse relationship does not occur. It is not known how significant this limitation is, but for the purpose of this investigation it was assumed to be insignificant.

If the flow rates are correct, then perhaps there are other sources for the fluids besides compaction. A source from outside the modeled portion of the basin or from the Cretaceous sediments cannot be ruled out, nor can the possible occurrence of some type of free convection, in which fluids flow in recirculating cells. Because fluids recirculate, much more flow can be generated for a given volume of fluid.

Convection cells might occur on a small scale, such as within a local sediment package or sand-body, or they might also exist on a much grander scale, where faults function as the vertical portion of the cell, and horizontal bedding planes and cross-bedding flow function to complete the cycle. Descending legs of the cells may be located off-shore and also northwest of the cross section, in the shallower Cretaceous.

A limitation of the model is the inability to account for the dependence of fluid density on dissolved chemical species. The model uses pure water as the fluid. But the deep-basinal brines may be more dense than pure water and would therefore require greater fluxes or other forces to concentrate them into upwelling plumes. However, if fluids released by consolidating clays are fresher than these brines, (Kharaka and others, 1980; Price, 1976; Russell, 1971), then these fluids would also be less dense than the brines. Consequently, upon release, if mixing were limited, these fluids would be buoyed upward, creating upwelling plumes of fluids where conditions were appropriate, and requiring less fluid flux from below to drive the flow. In this way, a sufficient upward flow could be sustained to account for the temperature distribution while requiring smaller overall fluid fluxes. Such an explanation would unfortunately unbalance the chloride concentration calculation.

Clearly, much work remains in deciphering the hydraulic and thermal regimes of the Gulf Coast Basin. This research has hopefully provided a spring-board for further research and insight into the complex

nature of this system.

CONCLUSIONS

Having compiled and assembled sufficient data to characterize the thermal distributions in a portion of the Tertiary rocks of the South Texas Gulf Coast Basin, it is evident that regional-scale processes are at work modifying the heat flux above the Cretaceous boundary. The most prominent feature of the temperature distributions is an arc-shaped region of high temperatures which correspond to the growth fault zone of the Wilcox Group. The anomaly begins to emerge at 10,000 feet (3048m) and becomes increasingly pronounced toward 20,000 feet (6096m), beyond which depth there are no data. Although high porosities associated with geopressures have been linked to lower thermal conductivities and higher geothermal gradients, this mechanism does not explain why the temperatures and gradients are highest in the faults zones, and lower in adjacent areas with similar geopressures. Instead, the fault zones themselves seem to be important, perhaps serving as conduits for vertical flow. The configuration of the thermal

ridge implies upwelling from below the 20,000 foot (6096m) extent of data.

This largely empirical method of deciphering the cause for the observed temperatures is limited because of the complexity of the problem. It is complex because many factors can influence temperature, including the nature of the heat source, thermal properties of the sediments, the thermal properties of the fluids which saturate the sediments, the hydraulic properties of the entire system through which the fluids move and advect heat, and the chemical behavior and interaction of the sediments and the fluids. Numeric modeling provides a way of accounting for these factors in a coordinated and quantitative way.

The modeling effort provided additional support for the hypothesis that upwelling fluids are the major cause of the thermal anomaly in the Wilcox trend. For example, it was determined that the optimum placement of flux nodes was along the basal boundary of the model, with the flux concentrated along the Wilcox fault zone. This configuration supports the initial inference that upwelling occurs from below 20,000 feet (6096m), implying in addition, that the flow is focussed at these depths. Permeabilities configured for the model allow flow pathways along the growth fault zone, but only within the deeper portions of the zone. The model calls for a sealing of the fault zone at intermediate depths nearing the top of geopressure. Speculative plugging mechanisms include cementation, differential healing of faults, differential fault movement and displacement, and depth constrained

microfacturing. A puzzling aspect of the model is the call for a zone of high hydraulic and thermal conductivity coastward of the deep Wilcox growth faults. This region, if it exists, could be explained by the presence of salt diapirs, the salt itself being virtually impermeable but of high thermal conductivity, while diapirism can lead to faulting and good vertical permeability.

The model leaves unsolved a problem common to numerous other studies involving fluid movement in the basin. Namely, that the fluid fluxes called for are beyond what can be supplied from compactional processes alone. Studies of dissolution and cementation call for fluxes that are even greater. For this reason, the occurrence of free convection has been suggested. It may occur on both a local and regional scale. Regionally, growth faults may be integrated into the circulation of convection cells, a circulation which would presumably be confined to the compactional system since fluids could not freely convect downward across the geopressure transition.

There is clearly much work still needed to decipher the hydraulic and thermal regimes of the Gulf Coast Basin. This research has hopefully provided a spring-board for further research and insight into the complex nature of this system.

APPENDIX A – DATA

This appendix contains first a complete listing of the names of the well-data base, followed by the file of all the data which is associated with each well. They can be cross referenced using the well number.

Township Well

Range	No.	Well Name
11S 21E	26	Harkins,*B-1 Fimbel
11S 21E 1	27	Assoc O&G et al,*1 Van Dohlen
11S 21E 3	28	Hewitt,*1 Baker
11S 21E 9	29	Plymouth,*1 Thompson
11S 21E 1	30	Hewitt,*1 Hall
11S 21E 1	31	Assic O&G et al, *1
Cunningham		
11S 21E 1	32	Texas O&G,*1 "A" Richter
11S 22E 3	33	Texas O&G,*1 Earp
11S 22E 1	34	Shell,*1-A Bego
12S 21E	35	Plymouth,*1 Thompson
9S 17E 1	36	Standard,*1 Rzeppa
9S 17E 2	37	Standard,*1 Moczygemba
9S 17E 3	38	Mortimer et al,*1 Dzuik
9S 17E 2	39	Coastal States,*1 Kowalik
9S 17E 2	40	Standard,*1 Pollok
9S 18E 7	41	Standard,*1 Pace
9S 19E 7	42	Tex East. Trans.,*1 Garbe
9S 19E	43	Tex East. Trans.,*1 Mugge
9S 19E	44	Tex East. Trans.,*2 Voelkel
9S 19E 7	45	Tex East. Trans.,*1 Woskow
9S 19E 8	46	Standard,*1 Tipton
9S 19E	47	Mustang,*1 Yanta
9S 19E	48	Occidental,*1 Osterloh et al.
9S 19E 6	49	Hunt Est.,*1 Schuenemann
9S 19E 8	50	*1 Flenniken
19S 20E	51	Southland,*1 Responder
19S 20E	52	Ranger,*1 Clark
19S 20E	53	Tidewater,*1 Musselman
19S 20E 5	54	Hamon,*1 Roeder
19S 20E 8	55	Hanson et al.,*1 Altman
19S 20E 1	56	Southland,*1 Gips
19S 20E 7	57	McCulloch,*1 Franke
19S 20E 7	58	Monsanto et al.,*1 Green
19S 20E 8	59	Monsanto et al,*1 Kolodzeg
19S 20E 7	60	Monsanto et al.,*1 Estrella
19S 20E 5	61	Tidewater,*1 Korth
19S 20E 1	62	Harper-Smith,*2 Gips
19S 20E 2	63	Hamon & Ehman,*1 Kleberg
19S 20E 8	64	Murphy Baxter,*1 Hilgartner
19S 20E 3	65	Untex,*1 Warwas
19S 20E 7	66	McCulloch,*1 Hurst et al.
10S 17E 9	67	Delta Am, *1 Sahm
10S 17E 7	68	General Crude,*1 Gunewald
10S 18E 9	69	General Crude,*1 McDowell
10S 18E 7	70	Atlantic Ref.,*1 Pullin et al.
10S 18E 6	71	General Crude,*1 Wessendorff
10S 18E	72	Bruns,*1 Tips

Township Well

Range	No.	Well Name
10S 18E 3	73	Harrison,*1 Hysaw et al.
10S 18E 5	74	Hunt,*1 Huckman
10S 19E 9	75	Shell,*1-R Atkinson
10S 19E 8	76	Hunt,*1 Zavesky
10S 19E 2	77	Argo,*1-A Dittmer Est
10S 19E 5	78	Cities,*1 Stanchos "A"
10S 19E 6	79	Cities,*1-A Janssen
10S 19E 5	80	Cities,*1-A Pargmann
10S 19E 1	81	Kirk & Neeb,*1 Mueller Est.
10S 19E 1	82	Cities,*1-A Gaus
10S 19E 7	83	Nat. Explor. Co.,*1 Effenberger
10S 20E 8	84	Continental,*42 Pettus
10S 20E 6	85	Lone Star,*1 Jank
10S 20E 3	86	Getty,*11 Nordheim
10S 20E 3	87	Southland,*1 Fuhrken
9S 13E 4	88	Continental,*1 Taylor & Brow
9S 13E 7	89	Corthay Land,*1 Brite
9S 14E 1	90	Exxon,*1 Heinen
9S 14E 3	91	Sorelle & Sorelle,*1 Heinen
9S 14E 7	92	Smith,*1 Westland
9S 16E 6	93	Seaboard,*2 Gabrysch
9S 16E 2	94	Coloma,*2 Dzuik
9S 16E	95	Seaboard-Hamon,*1 Szalwinski
9S 16E 8	96	Viking,*1 Clark
9S 16E	97	Renwar et al.,*1 Franklin
10S 15E 6	98	Burnwell,*1 Pfiel
10S 15E 7	99	Harkins,*1 Zunker
10S 15E 4	100	Continental,*1 Stienle
10S 16E	101	Filon,*1 Sturcken
10S 16E 8	102	Hundt,*1 Hardt
11S 13E 9	103	Dougherty,*1 Henry
11S 15E 1	104	Steward,*1 Mcilvaine et al
11S 15E 8	105	Buzzini,*1 Seale
12S 13E 8	106	Continental,*1 Horton
12S 13E 6	107	Tenneco,*1 Jamber
12S 14E 4	108	Chiles,*1 Morril
12S 14E 8	109	Tex East.Trans,*1-A Atkinson
12S 15E 4	110	Glasscock,*1 Terrel
12S 15E 7	111	Tenneco,*1 Schulz
12S 16E 2	112	Pickens,*1 Meyer
12S 16E 8	113	Vreeland,*1 Probst
12S 16E 8	114	Burns,*1 Stidde
12S 16E 1	115	Shell,*2 Turner
12S 16E 1	116	Shell,*1 Overby
12S 16E 7	117	Texas O&G,*1-C Smith
12S 16E 1	118	Shell,*1 Leppard
12S 16E 1	119	Shell,*1 Currer
12S 16E 2	120	Shell,*2 O'Neal

Township Well

Range	No.	Well Name
12S 16E 1	121	Shell,*1 Turner
12S 16E 8	122	Shell,*1 Braune
12S 11E 1	123	Pan Amer,*1 Franklin
12S 12E 6	124	Colo Oil,*1 Roark
12S 12E 1	125	Humble,*B-1 Gubbels
12S 12E	126	Supron,*1 Henry
9S 9E	127	Magnolia,*1 McKinley
9S 12E 9	128	Skelly,*1 Winkler
10S 9E 5	129	Pan Amer,*1 Oppenheimer
10s 11E 4	130	Gulf,*1 Reese
13S 14E 3	131	Standard,*1 Calliham
13S 14E 1	132	Standard,*1 Isaacks
13S 15E 7	133	Kilroy,*2 McClelland
13S 16E 9	134	Western Nat.,*1 Harris
13S 16E 8	135	Carri,*1 Smith
13S 16E 2	136	Peet,*1 Goebel
13S 16E 2	137	Coloma,*1-A Dove
13S 16E 1	138	Hanson et al,*1 Maguglin
13S 16E 8	139	Coastal St,*1 Morrison
13S 16E 1	140	Texas O&G,*1 Maguglin
13S 16E 1	141	Sinclair,*3 Dilworth
13S 16E 9	142	Shoreline,*1 Bartlett
13S 16E 4	143	Texas O&G,*10 McClelland
13S 16E 8	144	Love et al,*1 McMurray
13S 16E 8	145	Jones et al,*1 Martin
13S 16E	146	Newman,*1 Edwards
13S 16E 9	147	Huisache,*1 Allen
13S 16E 7	148	Texas O&G,*1-A Dunn
13S 16E 7	149	Continental,*1 McKinney
13S 16E 4	150	Neumin,*1 Hinton et al
13S 16E 7	151	Nat. Explo,*1 McKinney et al
14S 14E 8	152	Meeker,*1 Lebman
14S 14E 9	153	Kilroy,*1 Hays-Ezzell
14S 14E 4	154	Texam,*1 Hays-Ezzell
14S 14E 1	155	Kilroy,*1 Herring
14S 14E 1	156	Kilroy,*2 Herring
14S 14E 9	157	Kilroy,*2 Hays-Ezzell
14S 14E 6	158	Tenneco,*1 Stephens
14S 15E 7	159	Getty,*1 Christenson
14S 15E 1	160	Pen Oreille,*3 Nelson Tr
14S 15E 7	161	Limes,*1 Houdmann
14S 15E 6	162	Hanson,*1 Prosen
14S 15E 8	163	LaGloria,*1 Bush
14S 15E 8	164	Kilroy,*1 Beirwith
14S 15E 9	165	Atlantic Rich,*12 Lyne
14S 15E 6	166	So Petro,*1 West
14S 15E 9	167	Warren,*1 Whitley-Johnson
14S 15E 3	168	Inter. Amer.,*2 Herring

Township Well

Range	No.	Well Name
14S 15E 3	169	Venus,*1 Schreiner
14S 15E 9	170	Mesa,*1 Johnson
14S 15E 1	171	Coastal St.,*1 Lukas et al
14S 15E 5	172	Atlantic Rich,*2 Richter
14S 15E 1	173	Hanover,*1 Fair
14S 15E 5	174	Hamon-Carl,*1 Katzfey
14S 15E 5	175	Hamon-Carl,*1 Ybaney
14S 15E 6	176	Argo,*1 Braslau
14S 15E 6	177	Hughes,*1-A W Townsite
14S 15E 6	178	Crow-Muslow,*1 Smith
14S 16E 3	179	Pan Amer,*1 Scogin
14S 16E 4	180	Hughes-Hughes,*2 Kendall
14S 16E	181	Hanson,*1 Perkins
14S 16E 8	182	Atlantic,*1 Riser
14S 16E 7	183	Atlantic,*1 Baker
14S 16E 6	184	Brown,*1 Hayes
14S 16E 7	185	Smith,*1 Cravens
14S 16E 9	186	Abercrombie,*1 West et al
14S 16E 8	187	Pan Am,*1 Randall
14S 16E 3	188	Gulf,*1 Lee
14S 16E 4	189	Kilroy,*1-A Crocker
15S 13E 7	190	Argo,*1-A Edrington Est
15S 13E 9	191	Appell,*1 Atkinson
15S 13E 5	192	Steward-Quigley,*1 Rhode
15S 13E 1	193	Mobil,*1 Wheeler
15S 13E 5	194	Southland et al,*1 Coron
15S 13E 5	195	Peet,*1 Rhode
15S 14E 6	196	Atlantic,*1 El Paso Nat
15S 14E 5	197	Hamon,*1 Dougherty
15S 14E 7	198	Coastal St,*1 Lehmberg
15S 14E 2	199	Hurt MPS et al,*1 Riser
15S 14E 4	200	Arco,Rhode Sect.18
15S 14E 9	201	Rutherford,*4-A Baker
15S 14E 4	202	Atlantic Rich,*3 Rhodes
15S 14E 6	203	Aninoil,*1 Ausa et al
15S 14E 8	204	Rutherford,*2-A Baker et al
15S 14E 1	205	Davis,*2 Lyne
15S 14E 2	206	Kilroy,*1 Pettit et al
15S 14E 1	207	Davis,*3 Lyons
15S 15E 2	208	Hanson-Hurt,*1 Sparkman
15S 15E 4	209	Sohio,*A-2 Slick
15S 15E 6	210	Huber-Manco,*1 Tullis
15S 15E 5	211	Standard,*1 Burns
15S 15E 2	212	Stanolind,*1 Sparkman
15S 15E 5	213	Tidewater,*1 Burns
15S 15E 8	214	Placid,*1 Patterson
15S 15E 8	215	Tenneco,*1 Jones
15S 15E 6	216	Cities,*1 Bailey

Township Well

Range	No.	Well Name
15S 15E 2	217	Atlantic Ref,*1 Burns
15S 15E 9	218	Cities,*1-B Hendrick
15S 15E 4	219	Coastal St,*1 Ferrell
15S 15E 2	220	Coastal St.,*2-A Lennox
15S 15E 4	221	Hunt Sands,*1 Dolan
15S 15E 2	222	McMoran,*1 Burns
15S 15E 9	223	Hamill,*1 McClure
15S 15E 9	224	Monsanto,*1 McCullough
15S 15E 3	225	Austral,*1 Lyne
15S 15E 6	226	Hilliard,*1 Bailey
15S 15E 9	227	Katz,*C-1 Slick
15S 15E 4	228	Constantin,*1 Dolan
15S 15E 4	229	Atlantic Ref,*2 Lyne
15S 15E 3	230	SRG,*1 Lyne
15S 16E 2	231	Lone Star,*1 Watson
15S 16E 4	232	Tidewater,]-A Hall Est
15S 16E 4	237	Highland,*1 Cocker
15S 16E 3	234	Hanson-McCormick,*1 Johnson
15S 16E 8	235	Austral,*1-A Hinnant
15S 16E 6	236	Anschutz,*1 Parker
15S 16E 3	237	Daniel,*1 Peter et al
11S 17E	238	MGF,*1 Zuniga
11S 17E	239	Apache et al,B-1 Tips
11S 17E 5	240	Shell,*1 Ruhmann
11S 17E 6	241	Spartan,*1 Tips
11S 18E 5	242	Union,*2 Burnell Spielhagen
11S 18E 7	243	Union,*84 Ray
11S 18E 7	244	Union,N Pettus Ray7
11S 18E 9	245	Cox-Hering,*1 Copeland
11S 18E 3	246	Southland,*1 Mueller
11S 19E 4	247	Miller-Fox,*1 Lott
11S 19E 7	248	Amerada,*3 McFadden
11S 19E 7	249	Tex East Prod,*1 Hall
11S 19E 1	250	Haring-Alcoa,*1 Powell et al
11S 19E 1	251	Samedan,*1 Berckenhoff
11S 19E 7	252	Amerada,*3 McFadden
11S 19E 4	253	Miller-Fox,*1 Lott
11S 19E 8	254	Viking,*1 Ray Est
11S 19E 7	255	Brown,*1 Pereira
11S 19E 7	256	Coastal St,*1 McKinney
11S 19E 5	257	Fly,*1 Key Estate
11S 20E 6	258	Abel & Bancroft,*2 Neese
11S 20E 6	259	Abel & Bancroft,*2 Ramsey
11S 20E 1	260	McCormack,*25 Albrecht
11S 20E 4	261	Coastal St,*1 Ussery
11S 20E 6	262	Humble,*1 Neese
11S 20E 6	263	Normanly,*1 Nesse
11S 20E 4	264	General Crude,*1 Pettus

Township Well

Range	No.	Well Name
11S 20E 9	265	Cattle Land,*1 Gantt
11S 20E 9	266	Relco,*1 Gantt
12S 17E 3	267	Shell,*1 Alvarado et al
12S 17E 3	268	Shell,*1 Ford
12S 17E 7	269	Mosbacker-Century,*1 Lytle
12S 17E 9	270	South Tex,*1 McKinley
12S 17E 3	271	Shell,*1 Tomasek et al
12S 17E 3	272	Shell,*1 ONeal
12S 17E 7	273	Gasoline,*C-1 Holzmark
12S 18E 8	274	Atlantic Rich,*2 Dougherty Est
12S 18E 7	275	Coastal St,*1-A Scott
12S 18E 2	276	Harris,*2 West Tuleta-Hall
12S 18E 5	277	Skelly,*2 Hill
12S 18E 1	278	Hughes et al,*2 McKinney
12S 18E 1	279	Tex O & G,*1-A McKinney
12S 18E 2	280	Ellsworth-Stapp,*1 McKinney
12S 18E 6	281	Skelly,*1 Zook
12S 18E 2	282	Fly,*2 Ray
12S 18E 1	283	Hughes et al,*3 McKinney
12S 18E 6	284	Hanson et al,*1 Harris
12S 19E 3	285	Haring et al,*1 Ray
12S 19E 8	286	Coastal St,*1 Farish
12S 19E 4	287	Miller,*1 Farish
12S 19E 4	288	Tidewater,*1 Kidd
12S 19E 7	289	El Paso et al,*1 Hardison
12S 19E 4	290	Kissinger,*1 Kidd
12S 19E 8	291	Coastal St,*1 Farish
12S 19E 1	292	Coastal St,*3 McCord
12S 19E 6	293	Coastal St,*1 McCord
12S 19E 6	294	Coastal St,*4 McCord
12S 19E 2	295	Forest,*1 First VNTI Bank
12S 19E 3	296	Forest,*1 Ray
12S 19E 4	297	Magnolia,*2 Irby
13S 17E 3	298	Chorchran Zoch,*1 Lippard
13S 17E 6	299	Carl,*1 Gillette
13S 17E 6	300	Carle,*1 Turnbow
13S 17E 1	301	Hughes et al,*1 Pouloit "A"
13S 17E 2	302	Tex Co.,*1 Ragsdale
13S 17E 1	303	Hughes et al,*1 Macdon et al
13S 17E 1	304	Viking et al,*1 Williams
13S 17E 3	305	Cities,*B-1 Beck
13S 17E 7	306	N Central et al,*1 Bomar
13S 17E 2	307	Hamon,*1 Ragsdale
13S 17E 8	308	Argo,*1 Huegler
13S 17E 8	309	Colorado,*1 Choate
13S 17E 5	310	Colorado,*1 Tindal
13S 18E 2	311	Carri-Leahy,*1 Littlejohn
13S 18E 4	312	Tidewater,*1 Taylor

Township Well

Range	No.	Well Name
13S 18E 1	313	Coastal,*1 Scott
13S 18E 3	314	Dougherty et al,*1 Johnson
13S 18E 8	315	Sunray D-X,*1 May
13S 18E 8	316	Tamarack,*1 Humberson
13S 18E 3	317	Texaco,*1 Littlejohn
13S 18E 4	318	Tidewater,*1 Taylor
13S 18E 4	319	Tidewater,*1 Algea
13S 18E 9	320	Mitchell et al,*B-1 Robertson
13S 18E 2	321	Dougherty,*2 Perez
13S 18E 4	322	White Shield,*1 Martin
13S 18E 4	323	McCormick,*1 Littlejohn
13S 18E 4	324	McCormick,*1 Kearns
14S 17E 3	325	Shenandoah,*1 Good
14S 17E 2	326	Pend Oreille,*1 Womack
14S 17E 7	327	Humble,*B-9 Thompson
16S 17E 9	328	Max Pray,*3 Banker
16S 20E 8	329	Morgan,*3 Dodson Van Way
16S 20E 8	330	Morgan,*1 Johnson
13S 21E 9	331	Pure Oile,*1-B O'Brien
14S 22E 8	332	Kelly-Brock,*1 Snyder
14S 24E 7	333	Cities,*D-1 Talton
15S 21E 6	334	Cities,*1-A Rooke
15S 22E 9	335	4-B Trust,*1 Rooke
15S 23E 8	336	Sunray et al,*1 Hartman
15S 23E 7	337	Storm,St. Tr. *53,*1
15S 23E 3	338	Continental,*34 St Charles
15S 23E 7	339	Prairie,St Tr 12 *1
16S 21E 5	340	Superior,*27 Welder
16S 21E 9	341	Smith,*1 Cage et al
16S 21E 8	342	Atlantic Rich,*1 Fite
16S 22E 7	343	Cities,*1 Elzner
16S 22E 7	344	Fox,*1 Wendland
16S 22E 7	345	Fox,*1 Netek
16S 22E 8	346	Texaco,*3 Ritchie
16S 22E 1	347	Gloria,*A-1 Welder
16S 22E 6	348	Hurt,*1 Ritchie
16S 22E 9	349	Cherryville,p Sweatt
16S 23E 7	350	Hamon,*1 Banker Mort
16S 23E 2	351	Cities,St Tr 86 *1
16S 23E 8	352	Hunt Trust,*65-1 Banker M
16S 23E 1	353	Aluminum Co,St Tr 84 *2
16S 23E 6	354	Shenandoah,St Tr 120 *1
16S 23E 5	355	Amerada,*1 Bankers Mort
16S 23E 9	356	Graham,et al, Bible *2
16S 23E 6	357	Midwest,St Tr 122,*1
16S 23E 9	358	Amer Petrofina,*1 Ray et al
16S 23E	359	State Tract 128 *1
16S 24E 6	360	Getty-Mission,St Tr 125 *1

Township Well

Range	No.	Well Name
16S 24E 8	361	Amerada,St Tr 191 *1
16S 24E 2	362	Phillips,*1 Pano
16S 24E 4	363	Chevron,St Tr 125 *1
16S 24E 1	369	Getty,St Tr 94 *1
16S 24E 1	370	Getty,St Tr 119 *1
16S 24E 6	371	Getty,St Tr 118, *1
13S 9E 5	372	Gulf,*1 Naylor & Jones
13S 11E 6	373	Exxon,*13 Dilworth
13S 12E 9	374	Standard,*1 Henry
13S 12E 7	375	Pan Am,*1 McClaugherty
14S 9E 6	376	Phillips,*1 La Salle
14S 9E 3	377	Pan Am,*1 Foerster
14S 10E 1	378	Texas O & G,C-1 Washburn
14S 10E 3	379	Sutton,So Tex Syn *1
14S 10E 1	380	Pan Am,*1-A Alamo Ntl Bank
14S 11E 6	381	Amerada,*1 Grines
15S 11E 4	382	Phillips,*1-A Nueces
15S 12E 6	383	Ridley Wheeler,*1 Rives
15S 12E 7	384	Tex Co,*1 Atkinson
16S 9E 9	385	Smyth,*1 Withers
16S 10E 3	386	Texaco,*1 Canaies
16S 10E 8	387	Gulf,*1 Friedrichs
16S 11E 1	388	Mid Century,*1 Hubbend
16S 12E 9	389	Atlantic,*1 Hagist
16S 12E 8	390	Marion,*1 Heirs
16S 12E 6	391	Rutherford,*1 Pursch
16S 12E 1	392	Cox,*1 Atkinson
16S 12E 9	393	Atlantic Rich,*44 Foster
16S 12E 9	394	Arco,*2 Hagist
17S 5E 5	395	Burns Corp,*1148 Matrix
17S 7E 1	396	Southwest Petro,*1 Olmitos
17S 7E 8	397	Mobil,*5 Callaghan
17S 7E 3	398	Haynes-UT,*1 Olmitos
17S 7E 8	399	Southern Un,*4 Williams et al
18S 5E 8	400	Lamar Hunt,*1 Benavides
18S 5E 3	401	Lamar Hunt,*1 Hachar
18S 5E 3	402	MGF,*2-C Hachar
18S 6E 7	403	Mobil,*1 Gretchen et al
18S 7E 7	404	Gulf,*1 Hirsch
18S 7E 5	405	Sunray Mid,*1 Lincoln
18S 8E 8	406	Hunt,*1 Walker
17S 9E 9	407	Dougherty et al,*1 Nicholson
17S 9E 6	408	Appell,*1 Adami
17S 10E 5	409	Rowe,*1 Hendrickson
17S 10E 1	410	Blair-Vreeland,*1 Harlson
17S 11E 9	411	Mobil,*9 DCRC Sec 80 -
17S 11E 9	412	Magnolia,*79 DCRC
17S 11E 7	413	Marine,*1 Hahl Weiderkehr

Township Well

Range	No.	Well Name
17S 11E 6	414	Horizon,*1 Lundell 364
17S 12E	415	Atlantic Rich,*1 Arco et al
17S 12E 1	416	Texaco,Hugo State *1
17S 12E 1	417	Argo,*1 Gorman
17S 12E 6	418	Ramada,Welder Heirs 2-217
18S 9E 5	419	McDermott,*A-1 Mclean
18S 9E 7	420	The Tex Co,*B-1 Moos
18S 9E 8	421	Conoco,*2 Hughes est
18S 12E 5	422	Shell,A-3 Weatherby
18S 12E 5	423	Shell,*2 Lloyd
18S 12E 5	424	Shell,*1-B Hubberd
18S 12E 6	425	Shell,*1-A Weatherby
18S 12E 6	426	Shell,*2 Weatherby
18S 12E 6	427	Shell,*1 Hubberd
18S 12E 6	428	Socony Mobil,*1 Weatherby
18S 12E 5	429	Shell,*1 Penwell
18S 12E 5	430	Shell,*1-B Hubberd
18S 12E 6	431	Shell,*3Wenecke
19S 5E 4	432	Lamar Hunt,*1 Reuthinger
19S 5E 3	433	General Crude,*1 Ruethinger
19S 5E 3	434	General Crude,*1 Killam-Hurd
19S 7E 9	435	Mobil,B-1 Garza
19S 7E 8	436	North Cent,*1 Webb Co Schl
19S 8E 8	437	Delange et al,*1 Walker
19S 8E 1	438	Forest,*1 Russel-Winch
19S 8E 5	439	Good Hope,*1 Carr
20S 6E 2	440	Ginther et al,*1-A Killam
20S 6E 6	441	Ginther et al,*1 Killam
20S 7E 4	442	Ginther et al,*1 Benavides
20S 7E 3	443	Ginther et al,*1 Benavides
20S 8E 7	444	Bright-Schiff,*1-24 Atlantic
20S 8E 9	445	Good Hope,*3 Frost Ntl Bank
17S 13E 6	446	Pan Am,*1-E Farmers Life
17S 13E 6	447	El Paso,*1-B Yates
17S 13E 5	448	Davis,*1 Bravo Land
17S 13E 3	449	Texas Crude,*1 Farmers Life
17S 13E 1	450	El Paso,*2-B Farmers
17S 15E 7	451	Kidd,*1 Trejo
18S 13E 3	452	Cherryville,*1 Reyes
18S 15E 2	453	Esunas et al,*1 Hawkins
18S 15E 6	454	Howell,*1 Moos
17S 19E 9	455	Northern Pump,*1 Ocker
17S 20E 7	456	Sinclair,*1 Doney Jr
17S 20E 5	457	Renwar,*1 Peeks
17S 20E 4	458	Union,*1 Parker
17S 20E 7	459	Hamon,*B-1 Odem
17S 20E 1	460	Trans Cont,*1 Taft Morrow
17S 20E 6	461	Diamond Sham,*1 Cleo Pianta

Township Well

Range	No.	Well Name
17S 20E 8	462	Southern Min,*1 Griffith
18S 18E 6	463	Southern Min,*1 Stone et al
18S 18E 2	464	Bell et al,*1 Franks Est
18S 18E 8	465	Hughes et al,*4 Shulze
18S 18E 4	466	Amer Petrofina,*1 Kaiser
18S 18E 1	467	Gillring,*2 Dix
18S 18E 6	468	Dominion,*1 Charba
18S 19E 8	469	Chicago et al,*1 Lewis
18S 20E 4	470	Barnsdall,*1 Merritt
18S 20E 2	471	Gulf,*1 McGregor
19S 9E 5	472	Sundance,*1 Frost
19S 9E 3	473	Forest,*4 Winch-State
19S 10E 4	474	Houston,*1 Billings
19S 10E 5	475	Rowe,*1 Peal Est
19S 11E 9	476	Eason,*1 Peters Est
19S 11E 9	477	Eason-Harper,*1-160 Peters
19S 12E 1	478	Aluminum,*1 Hoffman et al
19S 12E 1	479	Maguire,*1 Braslau
20S 10E 8	480	Mobil,*1 Dinn
20S 10E 5	481	Hamon,*1 Leal
20S 10E 6	482	Morgan,*3-B Richardson
20S 10E 4	483	Shell,*1 Benavides
20S 11E 8	484	Argo,*A-21 Bennett
20S 11E 3	485	Amerada,*1 Gruy
19S 14E 5	486	Aluminum,*1 Oliveira
19S 15E 1	487	Delta Gulf,*1 Kalinec
20S 13E 7	488	Hiawatha,*B-1
20S 13E 9	489	Morris et al,*1 Southland
19S 18E 9	490	Hawley,*1 Treybig
19S 18E 3	491	*3 Ingram
19S 18E 5	492	Outline,*1 Hoepfner
19S 19E 6	493	Atlantic Ref,*1 Womack
19S 20E 5	494	Midwest,*1 Walton
19S 20E 4	495	Hamon,*1 Hoepfner
19S 20E 5	496	Southern Un,*1 Bevly
19S 20E 2	497	Hoover,*1 Baker
19S 20E 6	498	Sabine,*1 Doughty
19S 20E 3	499	Fisher-Davidson,*1 Klosterman
19S 20E 6	500	Inexco,*1 Bevly
19S 20E	501	Coastal St,*1 Cudding Airport
19S 20E 9	502	Anderman et al,*1 Ocker
19S 20E 3	503	May,*1 Merritt
20S 17E 1	504	Champlin,*127 Wardner
20S 20E 3	505	Texaco,*55 Lobby
20S 20E 7	506	Humble,*39 King Rch
21S 5E	507	Gulf,*A-2 Martin
21S 6E 5	508	Chevron,*1 Martin
21S 7E 9	509	Gulf,*A-10 BMT

Township Well

Range	No.	Well Name
21S 7E 9	510	Gulf,*A-11 BMT
21S 7E 9	511	Gulf,*A-2 BMT
21S 7E	512	Northern Natural,*1 BMT
21S 8E 9	513	Tex East,*1 Brunie Est
21S 8E 8	514	Skelly,*6 Martin
21S 8E 1	515	Gulf,*1 Villarreal
22S 5E 4	516	Consol O&G et al,*2 Zachry
22S 6E 9	517	Chevron,*1 Briones
22S 6E 7	518	Tex O&G,*4 Hein
22S 6E 1	519	Moore et al,*G-1 Hubbard
22S 6E 5	520	Good Hope,*2 Bruni Rch
22S 6E 8	521	Chevron,*1 McCann
22S 6E 7	522	HNG,*1 Hein Est
22S 6E 6	523	Tex O&G,*E-1 BMT
22S 6E 7	524	El Pason,*1 Hein Trust
22S 7E 3	525	Gulf,*A-11 BMT
22S 7E 2	526	Good Hope,*9 BMT
22S 7E 2	527	Good Hope,*1 Saldivar
22S 7E 3	528	Gulf,*A-16 BMT
22S 7E 3	529	Gulf,*A-4 BMT
22S 7E 8	530	Gas Prod,*1 Bartera
22S 7E 2	531	Mobil,*1 Bonugli
22S 8E 3	532	Coastal,*1 Martin et al
22S 8E 7	533	Superior,*1 McGrath
21S 9E 5	534	Pauley,*1 Laurel Fee
21S 9E	535	Hughes et al,*1-A Laurel
21S 10E 4	536	McCulloch,*1 Cuellar et al
21S 10E 2	537	Brown,*1 Benavides
21S 10E 5	538	Getty,*1 Benavides
22S 9E 2	539	Coastal St,*1 Puig
22S 9E 3	540	Atlantic Ref,*1 Puig
22S 9E 4	541	Pickens,*1 Brunie
22S 9E 4	542	Hamon,*1 Ramirez
22S 9E 9	543	Atlantic Ref,*A-1 Hinnant
22S 9E 2	544	Atlantic Ref,*1 Garcia
22S 9E 8	545	Mosbacher,*1 Fulbright
22S 9E 7	546	Austral,*2 Mclean
22S 9E 7	547	Atlantic Ref,*3 Mclean
22S 9E 7	548	Atlantic Ref,*1 Stroman
22S 9E 7	549	Atlantic Rich,*2 Stroman
22S 9E 7	550	Austal,*1 Mclean
22S 12E	551	Lone Star,*1 Miller
21S 14E 4	552	Jocelyn-varn,*1 Palacios
21S 16E 8	553	Sun,*117 Canales
22S 13E 1	554	Harkins,*1 Leal
22S 13E 7	555	Tex. Co,*1 Denman
22S 16E 5	556	Mobil,B-3 Blucher
22S 16E	557	Sun,*9 Sullivan

Township Well

Range	No.	Well Name
22S 16E 8	558	Sun,*1 Wright
22S 16E 6	559	Sun,*21 Canales
21S 17E 9	560	Morgan,*1 Hill
21S 17E 1	561	Southern Min,*1 Werneck
21S 17E 9	562	Exxon,*A-34 King Rch
21S 17E 6	563	Meeker Haas,*1 Burris
21S 17E 7	564	Meeker Haas,*1 Conner
21S 17E 3	565	Humble,*ME-1 Bornegos
21S 18E 7	566	Lone Star,*1 Muil
21S 18E 3	567	Fly,*1 Fisher
21S 18E 5	568	Sun,*1 Dietz
21S 19E 9	569	Mokeen,*1 May
21S 19E 1	570	Exxon,*15 King Rch Madero
21S 19E 9	571	Mokeen,*1 Salazar
21S 19E 9	573	Viking et al,*1 Muil
21S 20E 3	574	Exxon,*303 Alazon
21S 20E 5	575	Exxon,*308 Alazon
21S 20E 3	576	Exxon,*306 Alazon
21S 20E 5	577	Humble,*245 Alazon
21S 20E 9	578	Humble,*267 Alazon
21S 20E 5	579	Humble,*266 Alazon
22S 17E 3	580	Humble,*31 King Rch Canelo
22S 18E 7	581	Arkans Conroe,*1 Kaufer
22S 18E 4	582	Exxon,*6 Rincon
22S 18E 7	583	Arkansax McKeen,*1 Womack
22S 18E 1	584	Hill et al,*1 Brookshire
22S 18E 6	585	Cities,*1 Poteet
22S 18E 9	586	Humble,*4 Rincon
22S 18E 4	587	Exxon,*5 Rincon
22S 19E 8	588	Humble,*B-23 East
22S 19E 8	589	Sun,*1 Laguna Olmos
22S 19E 4	590	Forest,*1 Alvarado
22S 19E 9	591	Davis Chiles,*1 Koch
22S 19E 8	592	Humble,*B-25 East
22S 19E 6	593	Carre et al,St Tr 70 *1
22S 19E 3	594	Cities,*1 Yaklin
22S 19E 5	595	Cosden,*1 Hubert
22S 19E 7	596	Carrll Hawn,St Tr 68 *1
22S 19E 2	597	Pan Am,Cayo Del Grullo 84 *1
22S 19E 3	598	McMormick,*1 King Rch
22S 19E 4	599	Arkansas,*1 Huff
22S 19E 2	600	Morgan,*1 Matilda Limp
22S 19E 9	601	Tex O&G,*1-B Martin
20S 21E 6	602	King Ranch G-38
20S 21E 2	603	Marion *1 Heirs
20S 21E 1	604	Mobil *B-9 Chapman
20S 22E 1	605	Humble, ST Tr 173 *1
20S 22E 6	606	McMoran, ST Tr 171 *1

Township	Well	
Range	No.	Well Name
20S 22E 2	607	Hill, #1 Richardson
20S 22E 8	608	Am Petrofina, St Tr 169 #2
19S 21E 1	609	Nor Am, #1 Vision
19S 21E 4	610	Socony Mobil, #1 Russel
19S 21E 9	611	Morgan Ent, #1 Zdansky
19S 21E 4	612	Getty, #1 Bevly
19S 21E 9	613	Clinton, Sims 1-V Trust
19S 21E 7	614	Davis, #1 Gwynn
19S 21E 3	615	Nor Am, #1 Peterson Prop
19S 21E 2	616	Hamon, #1 Pulliam
19S 21E 2	617	Hamon, #1 Sikora
19S 21E 7	618	Heard BDK, Guaranty Nat'l Bk
19S 22E 4	619	Mitchell et al, #1 Nugent
19S 22E 4	620	Atlantic Ref, #1 Pearse
19S 22E 8	621	Coastal St, #1 Dunn et al
19S 22E 2	622	Anadarko, St Tr 181 #1
19S 22E 2	623	Cities, St Tr 96 #2
19S 22E 9	624	Marion, #1 Peterson
19S 22E 1	625	Cherryville, St Tr 81 #1
19S 22E 3	626	State Tract 88, #1
19S 22E 4	627	McMoran, St Tr 37 #1
19S 22E 3	628	McMoran, ST Tr 97 #2
19S 22E	629	McMoran, St Tr 97 #1
19S 22E 4	630	Coastal St, St Tr 83 #1
17S 21E 6	631	Phillips, #1 Flinn
17S 21E 7	632	Galaxy, #1 Moore
17S 21E 7	633	Houston Natrl, #1 Campbell
17S 21E 7	634	Galaxy, #1 Cities
17S 21E 8	635	Marathon, #2 Kellog
17S 21E 1	636	Hiawatha, #1 Green est
17S 22E 7	637	Am Petrofina, #1 Green Est
17S 22E 1	638	TerO&G, Ray "b" #1-A
17S 22E 9	639	Republic Natrl, #1 Floerke
17S 22E 4	640	Lawbar, Hunt Dugat #2
17S 22E 1	641	Hamon, #3 Harvey
17S 22E 9	642	Stark #1
17S 23E 1	643	Pennzoil, #1 Grant
17S 23E 9	644	Midwest, #3 McCampbell
17S 23E 3	645	Pan Am, #1 Banker's Mort
17S 23E 4	646	Hamon Sinclair, #1 Guettler
17S 23E 5	647	Getty, #1 McCampbell
17S 23E 8	648	McCampbell #1-A
17S 23E 1	649	Bass, Atlantic-Porterfield
17S 23E 2	650	Tenneco, #1 Barry Est
17S 24E 9	651	Mobil, St Tr 284 #1
17S 24E 9	652	Mobil, St Tr 310 #1
17S 24E 9	653	Getty, St Tr 275 #1
17S 24E 9	654	Tamarack, St Tr 274 #1

Township	Well	
Range	No.	Well Name
17S 24E 2	655	Skelly, #1 Schnitz
21S 22E 2	656	Stnrd, #1 Dunn McCampbell
21S 22E 9	657	Humbel, St Tr 196 #1
21S 21E 6	658	Shell, St Tr 206 #1
21S 21E 8	659	Samedan et al, St Tr 6 #1
21S 22E 2	660	Cherryville, St Tr 978-S #1
21S 22E 2	661	Sun, #1 Dunn-McCampbell
21S 22E 4	662	Sun, #3-A Dunn-McCampbell
21S 22E 6	663	Humble, #3 King Ranch et al
16S 25E 4	664	Hanson, St Tr 129 #1
16S 25E 3	665	Sunray, St Tr 96 #1
16S 25E 4	666	Getty, St Tr 142 #1
16S 25E 3	667	St Tr 77 #1
16S 25E 4	668	St Tr 143 #1
16S 25E 9	669	Exxon, Nine Mi Pt, #12
16S 25E 3	670	Getty, St Tr 95 #1
16S 25E 9	671	Phillips, St Tr 163 #1
15S 25E 4	672	Union, #15 Tatton
15S 25E 4	673	Union, #16 Tatton
15S 25E 6	674	Continental, #46 St Charles
12S 25E 5	675	Tenneo, #1 Schultz
9S 21E 4	676	Hanson, #1 Gips est
9S 21E 4	678	Skelly #1 Menn
9S 21E 5	679	Texaco, #2 Broughton
9S 21E 3	680	Musselman, #1 Danysh
9S 21E 5	681	Texaco, #1 Feller
9S 21E 9	682	Lone Star, #1 Gips
9S 21E 6	683	Brazos, #1 Sievers
9S 21E 9	684	Atlantic, #1 Kerlick
9S 21E 5	685	Houchins #2
9S 21E 1	686	Austral, #1 Schroeter
9S 21E 2	687	Monsanto Et Al, #1 Boldt
9S 21E 2	688	Quintana, #1 Jochen
9S 21E 6	689	Hamilton, #1 Harter
9S 21E 5	690	Monsanto, #3 Kulawik
9S 21E 1	691	Mosbacher, #1 Spies
10S 21E 1	692	Normandy, #1 Fromme
10S 21E 5	693	Cummings et al, #1 Urban
10S 21E 7	694	Houston, #1 Fromme
10S 21E 5	695	Burns, #1 Norrel
110S 21E 3	696	Lone Star, #1 Dohlmann
10S 21E 2	697	Mitchell, #1 Dohlmann
10S 21E 3	698	Lone Star, #1 Haynes
10S 21E 3	699	Humble, #2 McMillan
10S 21E 6	700	Kirkwood, #1 Hoff
7S 16E 3	701	Sohio, #1 Southern
8S 13E 6	702	Petro Tex, #1 Garcia
9S 12E 9	703	Skelly #1 Winkler

Township Well

Range	No.	Well Name
10S 10E 3	704	Stanolind,*1 Garcia
10S 11E 4	705	Gulf,*1 Morris Reese
15S 7E 2	706	Shell,*1 Martin
15S 7E 6	707	W Artesia Trans, *1 Koehne
15S 8E 8	708	Flamingo Vent.et al, *1 Coquat
16S 5E 1	709	Stampede,*3 Burkholder
16S 5E 3	710	Stampede,*6 Dawson
16S 5E 6	711	Mengden,*1 Carr
16S 6E 5	712	Mengden,*1 Schletze
17S 6E 6	713	Royal O&G,*1 Dellano et al
16S 7E 1	714	General Crude,*1 Miller
14S 7E 1	715	Bass,*1 Johnson
14S 8E 5	716	Stanolind,*1 Cooke
14S 8E 1	717	Suttons,*1 Hillje
14S 8E 4	718	Pan Am,*C-1 Cooke
7S 18E 8	719	Sun,*1 Patton
8S 17E 6	720	Morgan,*1 Pawelek
8S 19E 8	721	Geochem Sur,*1 Tam Jr
8S 19E 3	722	Wiatrek *1
8S 21E 4	723	Shell,*1 Brown
8S 21E 6	724	Superior,*1 Blackwell
8S 21E 1	725	Gulf,*1 Mueller
8S 21E 7	726	CommonWealth,*1 Richards
8S 21E 9	727	Shell, *1 Roehl
8S 21E 7	728	Howell,*1 Gansow
8S 21E 7	729	Am Petrofina,*1 Boldt
8S 21E 5	730	Highland,*1 Wood Jr
8S 21E 4	731	Esparanza,*1 Gohmert
8S 21E 5	732	Atlantic,*1 Smith
8S 21E 9	733	Shell,*1 Roehl
16S 13E 9	734	Harkins,Ragsdale Sec 120*1
16S 13E 7	735	Siegfried,*1 Lowe
16S 13E 8	736	Mobil,Labbe Rnch 413*1
16S 13E 9	737	Inland Ocian,*1 Roos
16S 13E 7	738	Petro-Lewis*1 Bindewald
16S 13E 6	739	Sunray,*1 Amer Nat Ins
16S 13E 6	740	Texaco,*1 Hughes
16S 13E 6	741	Sun Dx,*4 Stella Penn
16S 13E 4	742	Harkins,*1-112 Murphy est
16S 13E 1	743	Atlantic *1 Rhodes 54
16S 13E 5	744	Texaco,*27 Gouyer
16S 14E 1	745	Kilroy et al,*1 Dunn
16S 14E 3	746	Atlantic,*1 Glasscock
16S 14E 9	747	*1 Eubanks
16S 14E 3	748	Centura,*1 Roos
16S 14E 9	749	Austral,*1 Baker
16S 14E 3	750	Argo,*1 Dearman
16S 14E 3	751	Atlantic,*4 Baker

Township Well

Range	No.	Well Name
16S 14E 1	752	Skelly,Weil "A" *1
16S 14E 2	753	Mitchell,*1 Brookshire
16S 14E 1	754	Tenneco,*1 Patterson
16S 14E 6	755	Kilroy, et al,*1 Passmore
16S 14E 3	756	Hamon,*1 Hefner
16S 14E 3	757	Energy Soures,*1 Hinnant
16S 14E 2	758	Sanchez-O'Brien,*1 Jones
16S 16E 3	759	Humble, McCaslin *12
16S 16E 1	760	Tex O&G,Hinnant *5
16S 16E 2	761	Hughes & Hughes,Wilson*1
16S 16E 7	762	Tidewater/Skelly,NoSH-22

ID	UTM COORDINATES	DEPTH/TEMP PAIRS....
121	3163369.9	593195.3 14110 335 0 0 0 0 0 0 0 0 0 0 0 0
122	3155095.9	500172.7 6988 212 0 0 0 0 0 0 0 0 0 0 0
123	3163081.6	533221.1 10385 234 13960 303 15097 322 16601 336 0 0 0 0
124	3158748.1	547124.7 10844 262 0 0 0 0 0 0 0 0 0 0
125	3163773.5	548354.6 7346 207 8793 213 9689 243 10306 254 10823 261 0
126	3153383.6	539984.7 10272 262 0 0 0 0 0 0 0 0 0 0
127	3199115.0	501743.1 9612 246 11020 238 11434 251 0 0 0 0
128	3195146.6	537211.5 9839 208 14594 299 15061 312 0 0 0 0 0 0
129	3188466.9	508119.5 9242 231 0 0 0 0 0 0 0 0 0 0
130	3185210.9	524491.0 7572 218 8657 246 0 0 0 0 0 0 0 0
131	3150797.8	563501.2 7314 201 11510 277 11951 271 12024 271 12194 301 12361 303
132	3152597.5	569174.7 7284 198 12247 263 12395 301 12573 313 13610 362 0
133	3139142.1	584843.4 7978 228 8281 232 0 0 0 0 0 0 0 0
134	3139950.8	588502.0 8296 211 8575 224 0 0 0 0 0 0 0 0
135	3141647.2	590931.0 9008 230 0 0 0 0 0 0 0 0 0 0
136	3149363.6	593391.1 8542 229 0 0 0 0 0 0 0 0 0 0
137	3148839.7	593392.7 7309 216 0 0 0 0 0 0 0 0 0 0
138	3150280.4	596291.4 8459 216 10183 246 11144 287 0 0 0 0 0 0
139	3140528.1	592861.1 7929 234 9001 240 9392 239 0 0 0 0 0 0
140	3148686.6	597314.0 8262 249 8699 255 0 0 0 0 0 0 0 0
141	3151050.5	596387.0 7448 217 0 0 0 0 0 0 0 0 0 0
142	3140647.0	587636.9 8412 226 0 0 0 0 0 0 0 0 0 0
144	3140132.1	590538.2 7803 227 8793 240 0 0 0 0 0 0 0 0
145	3140862.2	591664.4 7499 212 9040 241 0 0 0 0 0 0 0 0
146	3144243.1	597229.4 9631 249 10042 292 0 0 0 0 0 0 0 0
147	3142527.6	589652.9 8905 258 0 0 0 0 0 0 0 0 0 0
148	3143046.2	595494.0 10649 312 11607 341 13828 348 14610 358 0 0 0 0
149	3141000.7	597411.9 9743 249 11363 267 13194 332 13818 339 0 0 0 0
150	3146522.4	587726.6 7623 223 0 0 0 0 0 0 0 0 0 0
151	3140843.5	596295.0 10367 262 11200 328 0 0 0 0 0 0 0 0
152	3126331.7	567243.4 7115 206 0 0 0 0 0 0 0 0 0 0
153	3128140.0	564643.3 8195 225 0 0 0 0 0 0 0 0 0 0
154	3133455.6	562901.7 7220 212 8224 209 10264 257 13228 342 14763 363 0
155	3135647.5	572847.9 8200 222 0 0 0 0 0 0 0 0 0 0
156	3135385.8	572160.2 8163 220 0 0 0 0 0 0 0 0 0 0
157	3128109.9	564656.3 8168 237 0 0 0 0 0 0 0 0 0 0
158	3132043.2	572473.2 8313 226 0 0 0 0 0 0 0 0 0 0
159	3126826.1	584014.8 6808 197 10231 297 11172 310 0 0 0 0 0 0
160	3134992.5	582259.4 8337 222 11732 291 12730 330 0 0 0 0 0 0
161	3128587.4	583341.1 8157 237 10619 270 12241 338 13701 356 0 0 0 0
162	3130907.2	582591.9 7774 215 9165 269 9679 291 0 0 0 0 0 0
163	3127711.7	578983.7 7535 209 8904 264 9107 268 9750 281 0 0 0 0
164	3128471.0	580992.5 7584 247 8114 288 0 0 0 0 0 0 0 0
165	3127715.6	576345.6 7225 214 9209 269 9710 277 0 0 0 0 0 0
166	3130874.3	584591.0 7530 223 8035 230 9232 266 0 0 0 0 0 0
167	3128963.2	577458.3 7294 197 8222 233 8706 240 0 0 0 0 0 0
168	3134164.2	575061.7 7793 222 0 0 0 0 0 0 0 0 0 0
169	3136248.2	576938.3 6935 206 8383 227 0 0 0 0 0 0 0 0

ID	UTM COORDINATES	DEPTH/TEMP PAIRS....
170	3128374.3	577534.6 7217 211 9568 271 0 0 0 0 0 0 0 0 0 0
171	3136935.1	583853.4 8307 224 0 0 0 0 0 0 0 0 0 0 0
172	3131543.1	579425.5 7721 213 9201 238 0 0 0 0 0 0 0 0 0
173	3134555.8	583263.9 7854 219 0 0 0 0 0 0 0 0 0 0 0
174	3130318.9	581065.9 7813 223 8350 247 9010 253 9692 269 0 0 0 0
175	3130020.6	580610.7 7602 226 9236 265 0 0 0 0 0 0 0 0
176	3132026.0	583671.4 7610 217 8235 237 8344 279 9334 301 9834 301 0 0
177	3133528.1	585892.2 8158 231 8824 296 9281 262 9708 270 0 0 0 0
178	3134038.8	585153.9 7750 227 9325 285 0 0 0 0 0 0 0 0
179	3137327.9	588613.4 9170 272 9202 270 0 0 0 0 0 0 0 0
180	3132687.3	585889.9 8467 223 9625 281 0 0 0 0 0 0 0 0
181	3130928.1	590828.1 8080 234 9882 291 10370 308 0 0 0 0 0
182	3128101.7	592937.1 9263 240 9765 279 10733 271 10829 273 0 0 0 0
183	3127521.1	594725.1 8885 232 10393 278 11654 293 12197 331 0 0 0 0
184	3133351.5	597514.9 8552 236 9847 285 0 0 0 0 0 0 0 0
185	3128191.4	594731.9 7770 213 8996 234 9933 291 0 0 0 0 0
186	3129395.5	587907.4 7852 216 9002 280 9543 271 10201 282 10714 305 0
187	3128738.2	591653.2 8686 248 9786 291 10800 309 0 0 0 0 0
188	3134929.8	589626.6 8777 230 9925 273 10895 284 11472 297 0 0 0 0
189	3131496.1	587395.5 8523 219 9803 254 10153 292 0 0 0 0 0
190	3113723.8	558600.1 8005 229 0 0 0 0 0 0 0 0 0 0
191	3112932.6	552200.9 8120 220 0 0 0 0 0 0 0 0 0 0
192	3116198.6	557027.8 7062 216 0 0 0 0 0 0 0 0 0 0
193	3124752.3	560702.7 7665 214 0 0 0 0 0 0 0 0 0 0
194	3118704.8	555693.2 7745 250 0 0 0 0 0 0 0 0 0 0
195	3117076.0	558237.3 7321 203 0 0 0 0 0 0 0 0 0 0
196	3116647.0	571197.3 7710 214 11072 303 11790 305 13100 364 14410 371 0
197	3116048.5	566962.0 7100 207 7849 219 0 0 0 0 0 0 0 0
198	3112544.2	572801.3 7722 209 10621 310 0 0 0 0 0 0 0 0
199	3122639.1	565364.9 7776 260 8253 259 0 0 0 0 0 0 0 0
200	3116392.0	562039.4 7739 221 10924 297 13574 366 0 0 0 0 0
201	311448.6	564566.5 6777 201 7759 217 8526 260 10067 298 10560 299 11059
314		
202	3116097.3	562943.2 8020 226 0 0 0 0 0 0 0 0 0 0
203	3116684.6	572782.6 10123 292 0 0 0 0 0 0 0 0 0 0
204	3113610.9	565973.0 7487 201 8987 221 9966 294 0 0 0 0 0
205	1.0	1.0 -513 -1 -513 -1 -513 -1 0 0 0 0 0
206	3122999.8	568247.6 7234 229 0 0 0 0 0 0 0 0 0
207	3122449.0	573189.4 7359 212 9976 297 0 0 0 0 0 0 0
208	3122323.9	577891.4 8527 249 10234 295 11838 333 0 0 0 0 0
209	3116540.8	576039.5 8653 230 10308 282 11014 299 0 0 0 0 0
210	3118718.1	582834.9 8550 235 11045 291 0 0 0 0 0 0 0
211	3117459.3	578224.5 8623 225 11165 298 12207 323 14093 379 14445 384 0
212	3123077.4	577290.4 8669 227 9639 289 10269 303 0 0 0 0 0
213	3119044.7	580894.2 8758 242 9809 280 10486 287 10968 294 0 0 0
214	3113098.3	580689.5 8754 240 10389 282 11545 321 12620 341 13596 375 0
215	3111570.4	581275.7 9775 256 10827 316 12483 330 0 0 0 0 0
216	3119349.5	584905.3 8587 234 10053 258 11098 312 0 0 0 0 0
217	3122556.9	580860.0 10208 293 11716 295 12394 304 0 0 0 0 0

ID	UTM COORDINATES	DEPTH/TEMP PAIRS....
218	3111722.5	574469.8 9595 281 10737 323 12540 358 0 0 0 0 0 0
219	3115964.5	576694.6 8665 230 10365 302 11865 337 0 0 0 0 0 0
220	3123498.9	581489.9 7151 214 7624 222 11505 333 0 0 0 0 0 0
221	3118823.7	575187.6 8491 231 11620 323 0 0 0 0 0
222	3121734.6	581065.0 8607 253 11671 321 12867 334 0 0 0 0 0 0
223	3112711.0	574611.6 8870 255 10311 306 10991 319 0 0 0 0 0 0
224	3112013.4	574078.2 11293 333 0 0 0 0 0 0 0 0 0
225	3120805.1	575099.4 8579 250 9182 256 9950 290 0 0 0 0 0 0
226	3117684.2	584778.7 9771 236 12410 327 13692 372 0 0 0 0 0 0
227	3114903.9	574101.7 8752 242 10290 300 10817 304 11825 332 0 0 0 0
228	3117630.5	575632.1 8746 244 10442 299 0 0 0 0 0 0 0
229	3119500.6	573849.1 8527 256 8830 252 9086 258 9628 276 10128 274 10236
		286
230	3124110.9	575267.3 7303 205 9160 247 9766 277 0 0 0 0 0 0
231	3123191.6	590376.0 7691 219 0 0 0 0 0 0 0 0 0
232	3120092.1	587298.4 7511 203 9589 266 10111 272 11212 295 0 0 0 0
233	3118948.7	586626.7 7967 230 10678 317 0 0 0 0 0 0 0
234	3123402.3	589783.5 7843 201 11283 318 0 0 0 0 0 0 0
235	3114763.1	592816.7 8901 238 11587 313 12318 337 13031 340 13509 366
		14533 378
236	3116634.8	595418.1 7839 225 9751 252 12653 321 13829 353 0 0 0 0
237	3122715.0	587481.6 7783 202 10618 312 11418 327 0 0 0 0 0 0
238	3171969.5	606718.5 13391 332 0 0 0 0 0 0 0 0 0
239	3174374.9	607268.1 8002 244 8186 254 0 0 0 0 0 0 0
240	3172611.0	604689.5 12915 284 13484 361 0 0 0 0 0 0 0
241	3173029.5	609132.6 7610 226 0 0 0 0 0 0 0 0 0
242	3174143.1	617140.3 9076 271 0 0 0 0 0 0 0 0 0
243	3168825.1	618871.3 7663 226 0 0 0 0 0 0 0 0 0
244	3167146.7	618689.3 8259 243 8587 268 0 0 0 0 0 0 0
245	3167308.5	612416.8 7177 203 0 0 0 0 0 0 0 0 0
246	3180316.9	610716.6 8688 270 13277 309 0 0 0 0 0 0 0
247	3175205.6	622980.1 6975 201 0 0 0 0 0 0 0 0 0
248	3170599.3	633974.7 10739 250 0 0 0 0 0 0 0 0 0
249	3169351.5	632847.8 7652 217 0 0 0 0 0 0 0 0 0
250	3176634.3	633737.9 8917 236 10323 288 0 0 0 0 0 0 0
251	3180798.7	632688.3 8805 223 9660 270 9975 278 0 0 0 0 0 0
252	3169914.5	633003.9 8148 221 10739 250 0 0 0 0 0 0 0
253	3175205.5	622974.0 6965 201 0 0 0 0 0 0 0 0 0
254	3167344.2	626985.6 8939 249 0 0 0 0 0 0 0 0 0
255	3170092.2	632257.8 9017 270 0 0 0 0 0 0 0 0 0
256	3169106.0	630694.4 9130 247 0 0 0 0 0 0 0 0 0
257	3173858.0	628606.2 10161 300 0 0 0 0 0 0 0 0 0
258	3174070.8	645232.0 9452 269 0 0 0 0 0 0 0 0 0
259	3175231.5	642990.5 8873 286 0 0 0 0 0 0 0 0 0
260	3178118.2	643294.5 9033 239 10693 302 0 0 0 0 0 0 0
261	3176470.3	636208.2 8430 233 0 0 0 0 0 0 0 0 0
262	3174397.7	645093.0 7767 212 8823 274 9272 280 9575 282 9734 281 10244
		287
263	3173657.1	645683.0 9768 269 0 0 0 0 0 0 0 0 0

ID	UTM COORDINATES	DEPTH/TEMP PAIRS....
317	3149574.2	614327.4 9986 250 0 0 0 0 0 0 0 0 0 0 0 0
318	3146383.1	613988.9 9085 242 10820 293 11879 348 0 0 0 0 0 0 0
319	3146005.8	612378.6 9026 232 10093 274 10913 292 10959 293 11647 327 11925 339
320	3141573.4	610270.0 9722 261 10704 300 0 0 0 0 0 0 0 0
321	3149385.5	614797.1 10127 282 11497 331 11922 348 0 0 0 0 0 0
322	3147293.8	6119566.5 9936 264 10189 308 0 0 0 0 0 0 0
323	3147226.9	610919.3 10116 268 0 0 0 0 0 0 0 0 0
324	3146183.6	613554.2 10010 262 0 0 0 0 0 0 0 0 0
325	3136184.7	600220.3 10129 262 0 0 0 0 0 0 0 0 0
326	3138850.3	602307.3 10348 282 0 0 0 0 0 0 0 0 0
327	3127470.3	644451.6 5883 190 0 0 0 0 0 0 0 0 0
328	3097920.8	601046.0 7400 207 0 0 0 0 0 0 0 0 0
329	3101577.1	642244.8 7850 194 0 0 0 0 0 0 0 0 0
330	3101505.4	641039.2 7924 198 0 0 0 0 0 0 0 0 0
331	3139980.6	648027.4 11887 285 14274 353 16213 384 0 0 0 0 0 0
332	3126250.5	665199.7 8961 210 0 0 0 0 0 0 0 0 0
333	3127752.4	695443.8 8400 199 9078 206 9589 216 0 0 0 0 0 0
334	3117404.9	657458.1 7971 201 0 0 0 0 0 0 0 0 0
335	3113962.2	661286.5 8480 202 0 0 0 0 0 0 0 0 0
336	3114345.1	676840.7 8031 198 8971 208 0 0 0 0 0 0 0
337	3112958.8	681620.4 9000 219 0 0 0 0 0 0 0 0 0
338	3123604.8	685364.9 8707 215 9792 226 11108 259 11937 240 12576 249 13109 297
339	3117032.9	693444.8 8666 214 0 0 0 0 0 0 0 0 0
340	3107094.2	653178.5 8570 234 11468 279 13481 296 15457 328 16827 364 17142 378
341	3098328.5	650707.1 8454 215 0 0 0 0 0 0 0 0 0
342	3098431.1	653946.4 9391 240 11112 276 0 0 0 0 0 0 0
343	3099146.7	667568.0 7718 194 9958 219 0 0 0 0 0 0 0
344	3098610.6	670395.2 10167 230 0 0 0 0 0 0 0 0 0
345	3098610.4	669979.6 9980 229 0 0 0 0 0 0 0 0 0
346	3100108.7	664222.4 9321 225 9694 231 0 0 0 0 0 0 0
347	3107454.3	669618.5 8989 210 9183 216 0 0 0 0 0 0 0
348	3105751.1	668517.5 7680 209 8757 215 8984 219 0 0 0 0 0 0
349	3099900.5	662382.1 9193 215 9475 220 0 0 0 0 0 0 0
350	3101154.7	681013.6 9135 215 11500 273 12140 283 0 0 0 0 0 0
351	3108768.3	678948.1 9232 206 0 0 0 0 0 0 0 0 0
352	3102449.5	677759.4 9380 215 10485 250 0 0 0 0 0 0 0
353	3109638.2	680320.6 9644 222 9982 229 0 0 0 0 0 0 0
354	3104569.3	679733.7 10009 211 10116 215 10328 251 0 0 0 0 0 0
355	3103330.9	678288.4 9042 208 10251 221 11996 281 0 0 0 0 0 0
356	3103137.0	673779.3 8411 222 9057 237 9469 241 0 0 0 0 0 0
357	3104648.5	681407.2 8807 211 9027 216 10666 260 0 0 0 0 0 0
358	3100010.8	672988.7 9990 249 0 0 0 0 0 0 0 0 0
359	3103698.2	681505.7 8752 225 10567 245 0 0 0 0 0 0 0
360	3106582.8	694751.5 9060 204 9814 215 0 0 0 0 0 0 0
361	3099161.6	691145.6 8422 201 9481 215 11021 237 0 0 0 0 0 0
362	3111898.4	690485.5 8995 215 10878 246 12479 265 0 0 0 0 0 0

ID	UTM COORDINATES	DEPTH/TEMP PAIRS....
363	3104760.8	684086.2 9034 206 12324 268 0 0 0 0 0 0 0 0 0
364	3109754.5	695764.4 9069 210 11379 268 0 0 0 0 0 0 0 0 0
365	1.0	1.0 -21 -1 -21 -1 0 0 0 0 0 0 0 0
366	1.0	1.0 -21 -1 -21 -1 0 0 0 0 0 0 0 0
367	1.0	1.0 -21 -1 -21 -1 0 0 0 0 0 0 0 0
368	1.0	1.0 -21 -1 -21 -1 0 0 0 0 0 0 0 0
369	1.0	1.0 -21 -1 -21 -1 0 0 0 0 0 0 0 0
370	3108382.2	695037.9 9056 219 10447 256 10918 270 0 0 0 0 0 0 0
371	3108589.9	696040.8 9178 220 11808 288 0 0 0 0 0 0 0 0
372	3144552.5	505813.6 7920 203 10578 242 0 0 0 0 0 0 0 0
373	3144945.4	533340.2 11617 297 13326 324 14217 346 0 0 0 0 0 0 0
374	3141493.9	537278.6 11125 255 11261 250 0 0 0 0 0 0 0 0
375	3140475.1	547726.4 11733 303 0 0 0 0 0 0 0 0 0 0
376	3132261.9	510950.0 9539 241 11586 271 0 0 0 0 0 0 0 0
377	3134328.4	503489.9 8795 297 11210 304 12356 368 13220 341 16267 377 17262 395
378	3135182.6	522216.9 16572 326 18374 399 20124 424 0 0 0 0 0 0 0
379	3138194.4	513461.3 10344 238 0 0 0 0 0 0 0 0 0 0
381	3133944.8	532769.1 13557 303 14561 362 0 0 0 0 0 0 0 0
382	3119135.0	524606.3 7327 205 12273 281 14915 313 17372 343 20746 413 23937 452
383	3117932.8	546185.9 8156 227 0 0 0 0 0 0 0 0 0 0
384	3113012.3	545421.5 7601 207 0 0 0 0 0 0 0 0 0 0
385	3100124.8	500166.5 6130 198 0 0 0 0 0 0 0 0 0 0
386	3107292.1	513950.2 7712 208 16628 342 0 0 0 0 0 0 0 0
387	3100721.1	519488.4 9031 240 15136 368 15448 405 0 0 0 0 0 0 0
388	3107094.8	533128.4 7703 197 0 0 0 0 0 0 0 0 0 0
389	3101094.4	539628.7 9049 217 9549 279 0 0 0 0 0 0 0 0
390	3097179.5	542113.2 7522 217 0 0 0 0 0 0 0 0 0 0
391	3103091.0	549024.6 8354 246 10406 302 10907 314 11385 327 0 0 0 0
392	3107344.0	547252.9 8178 224 8994 241 0 0 0 0 0 0 0 0
393	3099061.9	539422.0 7868 214 9870 290 10968 295 0 0 0 0 0 0 0
394	3101425.7	540140.1 9572 280 0 0 0 0 0 0 0 0 0 0
395	3088406.3	458107.4 7921 219 0 0 0 0 0 0 0 0 0 0
396	3096639.9	484730.0 8639 230 9023 232 9405 237 0 0 0 0 0 0 0
397	3087825.8	483311.3 6567 195 0 0 0 0 0 0 0 0 0 0
398	3093120.7	492019.0 6770 196 0 0 0 0 0 0 0 0 0 0
399	3084194.8	494285.1 7482 223 0 0 0 0 0 0 0 0 0 0
400	3072431.2	455681.2 12486 274 14279 369 0 0 0 0 0 0 0 0
401	3081822.9	454480.3 11786 277 13555 330 14628 334 0 0 0 0 0 0 0
402	3081615.2	454869.0 8465 232 0 0 0 0 0 0 0 0 0 0
403	3071451.8	474529.7 6459 195 6837 201 0 0 0 0 0 0 0 0
404	3072069.5	486303.7 7441 214 0 0 0 0 0 0 0 0 0 0
405	3076104.9	487322.3 8248 212 9032 224 0 0 0 0 0 0 0 0
406	3072280.9	492281.6 9662 256 13300 322 14090 342 17360 403 18520 437 20055 459
407	3087274.1	500151.9 7144 206 0 0 0 0 0 0 0 0 0 0
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409	3092037.4	517017.1 6637 208 0 0 0 0 0 0 0 0

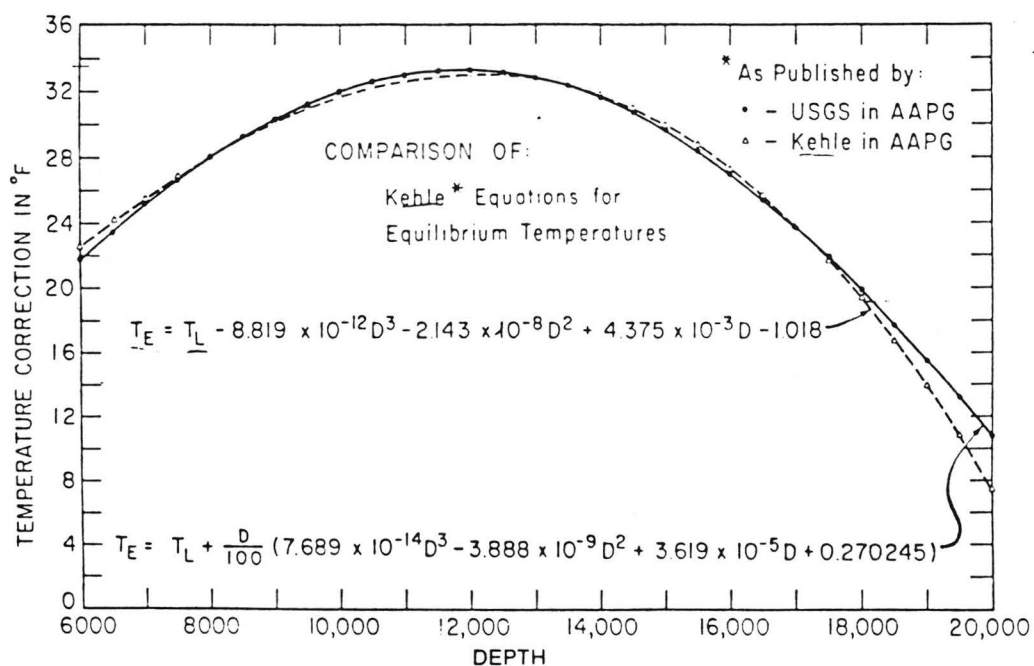
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411	3085741.3	528048.9 7717 219 10467 298 0 0 0 0 0 0 0 0 0
412	3084824.1	528273.5 7489 202 8959 251 0 0 0 0 0 0 0 0 0
413	3085355.0	535497.7 7348 197 7510 206 9569 241 9700 248 0 0 0 0 0
414	3091738.1	533730.9 8111 225 8655 258 0 0 0 0 0 0 0 0 0
415	3096796.1	548057.4 8520 248 10065 297 0 0 0 0 0 0 0 0 0
416	3093238.3	546260.2 8485 221 0 0 0 0 0 0 0 0 0 0 0
417	3095001.9	548141.4 8857 269 0 0 0 0 0 0 0 0 0 0 0
418	3091719.8	545661.1 8316 260 0 0 0 0 0 0 0 0 0 0 0
419	3076905.3	507740.8 7125 199 0 0 0 0 0 0 0 0 0 0 0
420	3070891.9	510284.1 7111 215 7395 219 0 0 0 0 0 0 0 0 0
421	3071623.8	504621.8 8013 209 10515 287 0 0 0 0 0 0 0 0 0
422	3077077.3	544807.7 8099 247 9655 280 13645 376 0 0 0 0 0 0 0
423	3076493.1	542791.5 7990 247 9894 298 13294 376 0 0 0 0 0 0 0
424	3077622.3	544396.7 8411 250 10414 326 11061 337 12387 367 0 0 0 0 0
425	3078001.9	545787.6 8299 249 11147 317 11948 372 13400 401 14323 414 15415 439
426	3076318.6	546096.0 8466 252 9466 267 0 0 0 0 0 0 0 0 0
427	3075906.6	545380.6 8442 244 8899 277 9447 295 0 0 0 0 0 0 0
428	3075803.9	546482.4 8602 276 9065 281 9361 287 0 0 0 0 0 0 0
429	3078026.8	542903.8 9119 230 11069 337 12349 368 12881 383 13274 383 0
430	3077689.2	544390.2 8425 250 10398 326 11045 337 12373 367 0 0 0 0 0
431	3074181.6	545835.3 7654 229 10020 306 14631 409 0 0 0 0 0 0 0
432	3063536.9	453624.7 12536 288 0 0 0 0 0 0 0 0 0 0 0
433	3064689.3	454767.7 8936 232 0 0 0 0 0 0 0 0 0 0 0
434	3058498.9	455074.0 9152 225 0 0 0 0 0 0 0 0 0 0 0
435	3057797.1	476666.1 9212 260 0 0 0 0 0 0 0 0 0 0 0
436	3058535.3	482881.0 7650 232 8748 256 0 0 0 0 0 0 0 0 0
437	3059710.4	492140.5 7326 201 0 0 0 0 0 0 0 0 0 0 0
438	3066942.8	499776.6 7486 216 9285 277 9832 291 0 0 0 0 0 0 0
439	3064491.4	494565.9 7428 207 10462 237 10467 266 0 0 0 0 0 0 0
440	3052507.1	470702.0 7328 210 9070 245 10647 266 11808 277 13140 314 0
441	3050633.5	472270.1 7379 214 10421 268 12412 308 0 0 0 0 0 0 0
442	3049128.9	476948.7 7023 220 0 0 0 0 0 0 0 0 0 0 0
443	3054017.5	477829.2 7927 221 0 0 0 0 0 0 0 0 0 0 0
444	3045008.0	499617.8 6783 216 0 0 0 0 0 0 0 0 0 0 0
445	3044392.8	491212.7 7106 211 10517 303 0 0 0 0 0 0 0 0 0
446	3089863.8	557919.5 7919 214 9469 275 0 0 0 0 0 0 0 0 0
447	3092013.1	558844.8 8756 250 9487 296 9800 302 0 0 0 0 0 0 0
448	3089345.3	554733.2 9014 236 9868 302 0 0 0 0 0 0 0 0 0
449	3096708.7	550364.2 8771 228 0 0 0 0 0 0 0 0 0 0 0
450	3092981.1	559308.7 8050 291 9005 287 9500 291 9636 292 0 0 0 0 0
451	3087740.6	583209.5 6703 195 0 0 0 0 0 0 0 0 0 0 0
452	3081731.3	551215.5 8355 250 9230 279 9993 305 0 0 0 0 0 0 0
453	3080694.5	578229.0 6470 195 0 0 0 0 0 0 0 0 0 0 0
454	3077197.1	584977.6 7116 212 7888 238 0 0 0 0 0 0 0 0 0
455	3091223.5	629616.0 7646 199 0 0 0 0 0 0 0 0 0 0 0
456	3085955.5	647405.4 9926 223 0 0 0 0 0 0 0 0 0 0 0
457	3091378.5	639370.6 8958 210 0 0 0 0 0 0 0 0 0 0 0

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459	3085011.7	646035.2 9773 215 10303 221 0 0 0 0 0 0 0 0 0 0
460	3097412.7	646726.8 9210 216 9543 241 10106 255 0 0 0 0 0 0 0 0
461	3089382.0	643630.8 9888 211 0 0 0 0 0 0 0 0 0 0 0 0
462	3088273.2	641650.6 8643 202 9287 223 9942 229 0 0 0 0 0 0 0 0
463	3079223.7	621053.7 8454 214 8704 216 0 0 0 0 0 0 0 0 0 0
464	3080015.8	616917.2 7085 198 0 0 0 0 0 0 0 0 0 0 0 0
466	3078906.6	614409.8 7450 213 0 0 0 0 0 0 0 0 0 0 0 0
467	3079312.2	621852.5 8416 219 0 0 0 0 0 0 0 0 0 0 0 0
468	3078827.3	622840.1 8461 214 0 0 0 0 0 0 0 0 0 0 0 0
469	3073118.3	628284.7 8420 223 0 0 0 0 0 0 0 0 0 0 0 0
470	3079016.5	637902.9 8018 217 0 0 0 0 0 0 0 0 0 0 0 0
471	3082494.4	641351.2 8897 210 0 0 0 0 0 0 0 0 0 0 0 0
472	3064407.6	508085.3 7629 213 0 0 0 0 0 0 0 0 0 0 0 0
473	3067547.0	501449.2 7610 216 0 0 0 0 0 0 0 0 0 0 0 0
474	3060053.2	515116.6 7924 228 8995 257 0 0 0 0 0 0 0 0 0
475	3061145.2	520208.0 6302 235 9309 291 0 0 0 0 0 0 0 0 0
476	3056381.0	527891.8 7507 248 10487 352 0 0 0 0 0 0 0 0 0
477	3057931.8	527454.3 7424 238 10521 339 0 0 0 0 0 0 0 0 0
478	3069106.6	546437.1 6964 227 8136 241 9335 303 0 0 0 0 0 0 0 0
479	3065921.8	545596.9 7413 229 0 0 0 0 0 0 0 0 0 0 0 0
480	3045527.6	520023.2 6743 220 10526 329 11038 347 11425 359 12741 394 0
481	3046327.0	520060.8 11105 328 11275 349 0 0 0 0 0 0 0 0 0
482	3047865.6	524347.2 8912 283 10316 302 11085 317 12285 349 0 0 0 0
483	3050300.9	514701.3 9011 241 12600 352 14197 391 19126 479 21639 507 22973 526
484	3045329.0	532070.4 6986 207 0 0 0 0 0 0 0 0 0 0 0 0
485	3054160.8	526869.7 7190 226 9841 254 10694 333 0 0 0 0 0 0 0
486	3063936.3	569076.8 6906 216 0 0 0 0 0 0 0 0 0 0 0 0
487	3068482.6	586230.5 8166 218 0 0 0 0 0 0 0 0 0 0 0 0
488	3045654.1	559059.1 8475 206 0 0 0 0 0 0 0 0 0 0 0 0
489	3046499.4	551449.2 5579 197 0 0 0 0 0 0 0 0 0 0 0 0
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491	3068423.9	614560.5 7464 197 0 0 0 0 0 0 0 0 0 0 0 0
492	3064732.4	615753.8 7115 204 0 0 0 0 0 0 0 0 0 0 0 0
493	3062056.3	634241.4 8396 232 8921 265 9200 270 9432 261 9672 265 9926 267
494	3062652.9	643215.2 8642 216 10120 251 0 0 0 0 0 0 0 0 0
495	3064428.5	639537.3 8929 214 9218 226 0 0 0 0 0 0 0 0 0
496	3063390.0	642228.5 8439 209 10486 247 0 0 0 0 0 0 0 0 0
497	3069936.8	640095.0 9129 215 0 0 0 0 0 0 0 0 0 0 0 0
498	3063199.2	647708.8 10246 240 0 0 0 0 0 0 0 0 0 0 0 0
499	3068053.3	635314.7 8232 201 0 0 0 0 0 0 0 0 0 0 0 0
500	3064705.7	645924.9 9593 237 9748 237 0 0 0 0 0 0 0 0 0
501	3066391.1	646604.4 9079 206 9849 217 10121 222 0 0 0 0 0 0 0
502	3060604.9	639515.0 8453 211 0 0 0 0 0 0 0 0 0 0 0
503	3068052.4	639561.3 8542 219 0 0 0 0 0 0 0 0 0 0 0
504	3053739.8	607427.5 7560 195 8797 234 10756 265 0 0 0 0 0 0 0
505	3052752.6	638416.3 10661 253 0 0 0 0 0 0 0 0 0 0 0

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507	3029388.6	465673.9	8389	237	0	0	0	0	0	0	0	0	0	0	0
508	3028939.7	474639.7	7491	213	9465	259	10500	280	0	0	0	0	0	0	0
509	3028273.0	477020.6	10199	280	0	0	0	0	0	0	0	0	0	0	0
510	3028264.4	479219.6	9898	267	0	0	0	0	0	0	0	0	0	0	0
511	3027716.5	478896.1	7400	207	10199	290	0	0	0	0	0	0	0	0	0
512	3031721.8	481270.2	9279	276	10037	342	0	0	0	0	0	0	0	0	0
513	3029299.5	490178.9	7241	207	0	0	0	0	0	0	0	0	0	0	0
514	3031806.8	492131.0	7638	230	8888	268	0	0	0	0	0	0	0	0	0
515	3038530.8	498071.6	6288	210	7800	239	0	0	0	0	0	0	0	0	0
516	3019222.8	454643.2	7206	225	9575	289	0	0	0	0	0	0	0	0	0
517	3015066.4	467125.5	6793	190	9208	265	9602	279	0	0	0	0	0	0	0
518	3017509.4	472037.8	7056	210	10146	288	0	0	0	0	0	0	0	0	0
519	3025489.6	471462.2	9459	217	9609	219	0	0	0	0	0	0	0	0	0
520	3022140.7	470325.4	7000	197	0	0	0	0	0	0	0	0	0	0	0
521	3014602.7	467746.2	7066	206	9566	286	0	0	0	0	0	0	0	0	0
522	3016279.7	471396.8	6596	197	0	0	0	0	0	0	0	0	0	0	0
523	3020146.2	472650.0	6612	207	0	0	0	0	0	0	0	0	0	0	0
524	3016788.6	473833.9	6824	207	10241	307	0	0	0	0	0	0	0	0	0
525	3027614.9	477202.6	6829	197	10159	282	0	0	0	0	0	0	0	0	0
526	3026325.1	481190.9	7346	201	0	0	0	0	0	0	0	0	0	0	0
527	3024991.2	481134.5	9972	262	0	0	0	0	0	0	0	0	0	0	0
528	3027685.7	479249.4	7180	199	10105	288	0	0	0	0	0	0	0	0	0
529	3026895.3	477969.3	9522	261	0	0	0	0	0	0	0	0	0	0	0
530	3014229.2	480062.3	7359	212	11039	309	0	0	0	0	0	0	0	0	0
531	3023734.7	482619.4	7674	205	11982	335	0	0	0	0	0	0	0	0	0
532	3025223.2	491228.0	6641	208	0	0	0	0	0	0	0	0	0	0	0
533	3017064.0	499198.1	6790	218	7829	259	0	0	0	0	0	0	0	0	0
534	3036224.0	504753.8	6588	213	7896	253	8338	264	0	0	0	0	0	0	0
535	3032123.2	507712.1	7133	238	10521	317	0	0	0	0	0	0	0	0	0
536	3034796.2	515050.3	6686	228	9280	305	9784	328	9980	334	0	0	0	0	0
537	3041222.5	516850.1	11432	336	0	0	0	0	0	0	0	0	0	0	0
538	3034638.5	517618.3	7247	236	9641	318	9849	319	0	0	0	0	0	0	0
539	3025112.8	506679.1	7416	246	8758	287	9803	307	10746	329	0	0	0	0	0
540	3024990.3	503137.6	7107	236	9151	267	10606	301	0	0	0	0	0	0	0
541	3020651.6	502925.5	10174	279	0	0	0	0	0	0	0	0	0	0	0
542	3019361.1	501094.2	6631	214	9139	281	0	0	0	0	0	0	0	0	0
543	3014925.4	503145.8	6595	203	9209	266	10155	314	0	0	0	0	0	0	0
544	3026731.2	505473.9	7822	229	9131	292	0	0	0	0	0	0	0	0	0
545	3016039.6	505780.5	7650	219	10342	300	11137	323	11240	323	0	0	0	0	0
546	3015724.1	510155.8	6655	204	7671	241	8840	265	0	0	0	0	0	0	0
547	3014300.2	510150.0	6662	196	8538	229	9246	234	0	0	0	0	0	0	0
548	3016661.2	510295.6	6660	199	8844	270	9561	290	0	0	0	0	0	0	0
549	3016825.9	511896.0	8228	250	10770	333	11188	348	12022	368	0	0	0	0	0
550	3015718.2	511007.8	6647	193	9636	275	10503	283	0	0	0	0	0	0	0
551	3026981.9	539027.7	8466	261	9710	267	10623	312	12282	347	12761	350	0	0	0
552	3035985.7	564335.8	5935	203	6458	207	6873	212	7688	224	0	0	0	0	0
553	3028946.6	591216.8	8349	214	9269	225	9859	235	12199	264	15943	369	19492	0	0

ID	UTM COORDINATES	DEPTH/TEMP PAIRS....
554	3027843.8	559559.0 7541 197 0 0 0 0 0 0 0 0 0 0 0 0
555	3015206.0	558413.1 7943 211 0 0 0 0 0 0 0 0 0 0 0 0
556	3022447.0	591223.3 8072 210 10746 253 0 0 0 0 0 0 0 0 0
557	3014739.6	595005.8 9266 215 10110 250 12003 295 0 0 0 0 0 0 0
558	3015401.4	594523.3 8458 207 12393 291 0 0 0 0 0 0 0 0 0
559	3020855.9	595417.4 8180 206 11684 286 0 0 0 0 0 0 0 0 0
560	3028798.6	601789.1 9006 220 0 0 0 0 0 0 0 0 0 0 0
561	3040646.3	609512.0 7941 198 8337 204 8321 203 0 0 0 0 0 0 0
562	3028599.8	599466.3 8291 218 10766 280 11559 306 12187 322 0 0 0 0 0
563	3034835.6	608742.8 8428 217 0 0 0 0 0 0 0 0 0 0 0
564	3030885.5	607739.3 9133 228 0 0 0 0 0 0 0 0 0 0 0
565	3039015.9	602419.4 7687 191 9480 220 11161 270 0 0 0 0 0 0 0
566	3031685.4	622623.1 10149 221 12974 298 14929 334 15879 353 0 0 0 0 0
567	3040088.2	612386.6 8306 204 9004 214 9304 216 9437 226 0 0 0 0 0
568	3032908.6	619155.6 7951 194 8960 208 9519 215 9740 229 0 0 0 0 0
569	3029458.7	625589.5 9240 202 9318 204 9683 211 0 0 0 0 0 0 0
570	3037952.4	633180.4 10071 233 10437 240 0 0 0 0 0 0 0 0
571	3030164.1	625251.0 8900 216 9592 226 10286 237 0 0 0 0 0 0 0
572	3040552.9	632874.3 9373 218 10550 258 0 0 0 0 0 0 0 0
573	3031743.5	624073.6 9987 225 0 0 0 0 0 0 0 0 0 0 0
574	3041915.2	636254.5 9759 252 0 0 0 0 0 0 0 0 0 0 0
575	3034804.6	643162.2 8674 218 0 0 0 0 0 0 0 0 0 0 0
576	3039434.8	636067.0 9383 210 9513 219 10222 242 11947 284 0 0 0 0 0
577	3035699.6	640076.1 9172 211 0 0 0 0 0 0 0 0 0 0 0
578	3029529.5	636804.4 9846 211 9981 216 14954 319 15573 336 0 0 0 0 0
579	3034200.3	640116.8 9451 215 9944 239 13602 311 0 0 0 0 0 0 0
580	3026499.3	602746.6 8835 214 0 0 0 0 0 0 0 0 0 0 0
581	3018812.1	622245.2 7875 188 8979 208 9480 209 0 0 0 0 0 0 0
582	3018646.3	614019.4 10437 241 0 0 0 0 0 0 0 0 0 0 0
583	3018703.6	621487.7 9548 209 9982 223 10479 229 0 0 0 0 0 0 0
584	3026156.6	620003.7 9977 225 0 0 0 0 0 0 0 0 0 0 0
585	3019463.5	623170.6 10147 231 0 0 0 0 0 0 0 0 0 0 0
586	3017566.0	614867.5 10949 244 0 0 0 0 0 0 0 0 0 0 0
587	3019417.4	614414.7 9933 236 0 0 0 0 0 0 0 0 0 0 0
588	3015206.2	630837.1 9784 223 10370 246 10465 249 12249 288 14187 325 14679 330
589	3017979.3	630297.6 9537 232 10215 250 11235 272 11724 271 0 0 0 0 0
590	3019929.8	625513.6 10254 230 10767 235 10871 242 0 0 0 0 0 0 0
591	3019203.3	626878.7 9473 213 10364 224 0 0 0 0 0 0 0 0
592	3014813.3	628936.2 9706 204 9910 227 11118 261 0 0 0 0 0 0 0
593	3019345.9	635183.7 10455 232 10549 244 11187 266 11748 274 12506 288 15060 325
594	3026202.8	624308.4 9399 215 10014 221 10524 250 11031 254 0 0 0 0 0
595	3020072.3	630976.8 9985 221 11495 269 12495 309 0 0 0 0 0 0 0
596	3017363.4	634588.0 10092 228 11361 245 11915 283 13239 9
597	3028253.5	628643.8 8874 211 9337 215 10057 230 0 0 0 0 0 0 0
598	3025116.9	625187.0 10284 224 11984 269 0 0 0 0 0 0 0 0
599	3020222.1	623948.6 8776 199 9725 219 0 0 0 0 0 0 0 0
600	3027423.2	627945.8 9060 216 9902 223 11360 261 0 0 0 0 0 0 0

APPENDIX B-THE KEHLE TEMPERATURE CORRECTION SCHEME



APPENDIX C-CALCULATING ISOTHERMS

Program DBINTP

This program creates a data file (OUTPUT) for an isothermal surface. The program interactively asks for the temperature of the isothermal surface and the temperature range (\pm ILIM) of eligible data. The program reads CORCT which is the filename of the temperature corrected data base.

```
PROGRAM DBINTP(CORCT,OUTPUT,TTY,TAPES=CORCT,TAPE6=
1OUTPUT,TAPE7=TTY)
  DIMENSION ITEMP(9)
  DIMENSION IDEPTH(9)
  NC=0
  WRITE(7,5)
5  FORMAT(1X,'WHAT TEMP TO PLOT,WHAT LIMIT OF INTERPOLA
  TION')
  READ(7,*)INTEM,ILIM
  DO 100 I=1,1000
  READ(5,*,END=110) NUM,DUM1,DUM2,IDEPTH(1),ITEMP(1),
1IDEPTH(2),ITEMP(2),IDEPTH(3),
1ITEMP(3),IDEPTH(4),ITEMP(4),IDEPTH(5),ITEMP(5),IDEPTH(6
1),ITEMP(6)
  DO 30 K=1,6
  IF(ITEMP(K).EQ.INTEM) GOTO 50
  IF(ITEMP(K).LT.INTEM AND. INTEM-ITEMP(K) .LE.ILIM
1AND. (ITEMP(K+1).GT.INTEM .OR. ITEMPT(K+1).EQ.0)) GO TO 40
  IF(ITEMP(K).GT.INTEM AND. ITEMPT(K)-INTEM.LT.ILIM) GO
1TO 70
```



```
30  CONTINUE
    GOTO 100
40  IF(ITEMP(K+1).EQ.0)THEN
      ICRCO=ITEM*IDEPTH(K)/ITEMP(K)
    ELSE
      ICRCO=(IDEPTH(K+1)-IDEPTH(K))*(ITEM-ITEMP(K))/
1    (ITEMP(K+1)-ITEMP(K))+IDEPTH(K)
    END IF
50  WRITE(6,55) NUM,DUM1,DUM2,ICRCO
      NC=NC+1
55  FORMAT(1X,I5,2F13.1,I6)
60  GOTO 100
70  ICRCO=ITEM*IDEPTH(K)/ITEMP(K)
    GOTO 50
100 CONTINUE
110 WRITE(7,120)NC
120 FORMAT(1X,'THE NUMBER OF ENTRIES IS',I5)
    WRITE(6,130)NC
130 FORMAT(1X,'C** THE NUMBER OF ENTRIES IS',I5)
    STOP
    END
```


APPENDIX D- KRIGING

Author's note: Before embarking on the study of this appendix, you are hereby warned that the Kriging process is rather involved and will not be learned in a day, nor probably even in a week. You will first need: A basic knowledge of FORTRAN, data that are digitized using a cartesian system (eg. UTM), access to the programs covered herein (GAMM8B, NEWPLTB, NEWUKB, & MAPWB which are property of the UT Bureau of Economic Geology, Contact the Bureau's computer center) and familiarity with the computer on which you intend to run them.

Kriging is a data smoothing technique which employs statistical methods to accomplish smoothing. The method can transform irregularly spaced or unevenly distributed data into regularly spaced points, each point being estimated from the actual data points surrounding it. A standard deviation at each of the estimated points is also calculated. The method was originally developed for estimating mineral recovery in mining operations, but has been applied to other geostatistical problems such as preparing data for contouring in cases where the reliability of the data is not absolute. In such cases, contouring the data without some way of first smoothing out the scatter can result in an impossibly

complex and unrealistic surface. Kriging allows one to average out the scatter while providing a measure of the reliability of the resulting estimated surface at each contoured point. In my work, the kriging technique has been employed to smooth bottom-hole temperature measurements for contouring subsurface isotherms.

The kriging process involves several steps. A computer is essential for handling the multitude of calculations required. This appendix does not attempt to explicitly review all the calculations needed. Presented is a general outline of the procedures, accompanied by instructions in the use of the various computer programs that I used to accomplish them. The programs were donated to the Bureau of Economic Geology of the University of Texas by Dr. Young C. Kim of the University of Arizona College of Mines.

THE VARIOGRAM

Constructing a variogram is the first step in the kriging process. A variogram is a plot which graphically describes how the data varies on average at increasing distances away from any point within the data field. It can be thought of as a plot of the difference in value between some starting point and points increasing distances away from that starting point. A typical variogram is shown in figure D1 . The underlying assumption of the variogram, and hence the kriging process itself, is that the values of neighboring points influence each other. In other words, we expect the variogram to show the smallest differences at points closest to the starting point (represented on the variogram by

the origin), with the differences becoming greater at increasing distances away from the starting point. Data that behaves in this manner is termed regionalized. As distance increases, the regionalized influence diminishes and variation increases until some plateau is reached where differences become purely random.

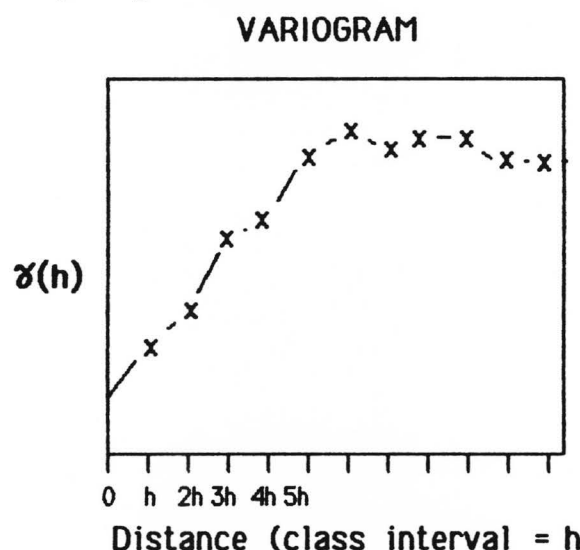


Figure D1. A typical variogram where the x's are $\gamma(h)$'s plotted at each multiple of h .

The variogram is an integral part of the Kriging process. In essence the kriging program I used constructs a regularly spaced grid over the data field and then finds values for the center of each block in the grid. It determines the value of each block center by averaging the surrounding data points. This is not a simple average because the kriging process weights each surrounding datum separately. This weight is based on its distance from the block center. The program knows how to weight the data by referring to the variogram as shown in figure D2. In general, points closest to the block center receive greater weight than those

farther away. In actuality, the kriging program cannot read the variogram directly. Instead, it must be given an equation of a curve which closely approximates the variogram. This curve is shown in figure D2 by the heavier line. The details of fitting a curve are supplied in the section on NEWPLOT.

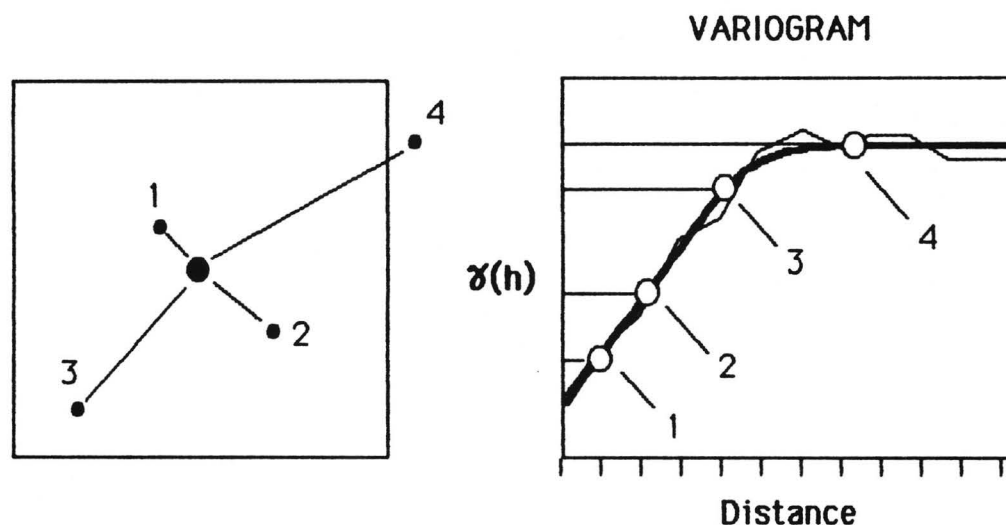


Figure D2 shows how a variogram is used to krig a block center (left). The data points 1-4 around the block center are weighted according to the variogram (right). The weighting values are derived from the γ 's read off the variogram, which in turn are approximated by fitting a curve to the actual variogram. The weighted values are then averaged yielding a value for the block center.

The variogram is constructed by surveying within a set distance around each data point, finding the average difference of the value between the point being surveyed and the points that occur within the set distance from it, and then averaging this result with those of all the other data points in the data field. This number is then plotted on the

variogram, with the average difference plotted along the y axis and the distance surveyed plotted along the x axis. The process is then repeated for multiples of the original surveying distance. The distance used in the original survey is set arbitrarily and is termed the *class interval*. Finding the optimum class interval is a trial and error process. The investigator guesses at the class interval until a proper looking variogram such as shown in figure D2 results. Note in the figure that the y axis is labeled gamma. Gamma is actually the standard deviation squared. The squaring eliminates the problem of dealing with both positive and negative variations. If this were not done, the average differences would be near zero as a result of the positive and negative variations cancelling each other. As one can now see, there is a staggering amount of calculation involved in finding all these averages for each data point, and then averaging these together. This technique would not be practical without the aid of a computer.

The above method is stated in equation form as follows:

$$\gamma[h] = \frac{1}{2N} \sum_{j=1}^N [f(x_j) - f(x_j + h)]^2$$

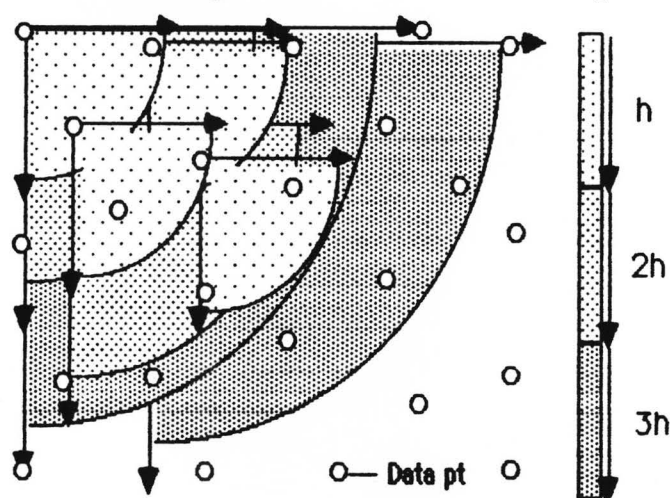
where N is the number of data points, γ is the standard deviation squared, and h is the class interval. To construct the variogram, γ is plotted against distances h and multiples of h.

VARIOGRAM COMPUTER PROGRAM

THE SEARCH WINDOW

When calculating γ at multiples of class interval h , it is possible to control the direction of search. In other words, one need not necessarily sweep 360 degrees around each data point. In fact, since every data point is used as a starting point of search for each class interval, the search angle, also termed the *window*, need only be 90 degrees to cover the whole data base in all directions as shown in figure D3. By decreasing the window angle further and specifying a direction, such as N, NW, SE ect., a variogram showing behavior of the data in one direction is created (figure D4). This variogram can be compared to variograms resulting from different window orientations for the purpose of estimating anisotropy. The section Fitting A Curve contains some additional information on this. It is necessary to introduce the concept of the search window here since the following program requires that the window be specified.

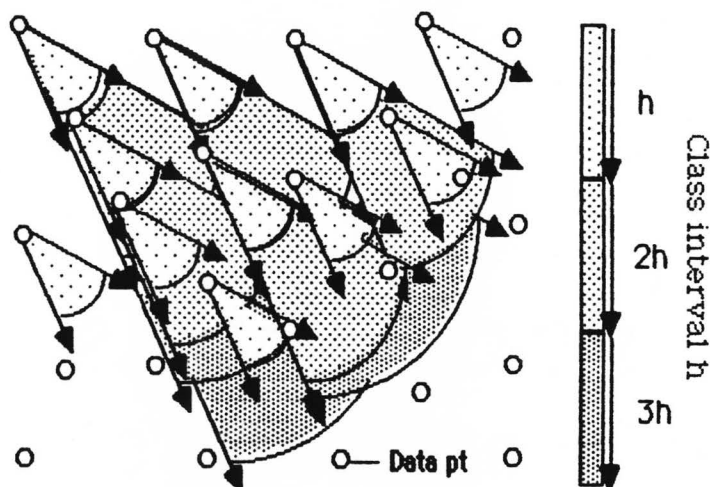
Constructing a Nondirection Variogram



90° Windows

Figure D3 schematically showing sweeps of the data field for h , $2h$ and $3h$ intervals for a nondirectional variogram.

CONSTRUCTING A NW-SE DIRECTIONAL VARIOGRAM



45° Windows

Figure D4 schematically showing sweeps of the data field for h , $2h$ and $3h$ intervals for a directional variogram.

THE UTM COORDINATE SYSTEM

For all the programs used to analyze my data, a cartesian coordinate system is required as the means to locate the data points. The usual latitude and longitude will not work as these are not cartesian. Another mapping system was developed for this need and it is called the *Universal Transverse Mercator* system or *UTM* for short. Unfortunately, there are some problems with UTM that can make it difficult to use. UTM does not account for earth curvature, and for a very large area, say the size of the state of Texas, this can be a problem, depending on how accurately the data need to be located. For my study this was not a problem. Another problem occurs at central meridians, which occur every six degrees of longitude, where the UTM system is discontinuous. This is a concern only if the Bureau's printed UTM listings are used. This may be convenient for some applications, but if data cross over one of these meridians, a program which the Bureau calls LLTRANB should be used to get your UTM coordinates.

VARIOGRAM PROGRAM GAMM8

The following instructions apply to the use of the FORTRAN program GAMM8B which is the property of the Bureau of Economic Geology. This and other programs used in my thesis are not in the public domain and are not listed here for that reason. A knowledge of FORTRAN formatting is assumed.

To use this program, one must create an input deck. The input deck contains information which controls the various options of the program as well as supplying basic information about the data. The data are then placed after this initial information. When running the program, the input deck must be specified as the input file. For those not familiar with constructing an input deck: The term *card* in the listing below refers to the now seldom used computer punch cards. An entry line on a terminal display is the modern equivalent of a card. There are 80 columns to a card and each column corresponds to one character. An input deck should have the following form (**Bold** type indicates the usual selections):

CARD	COLUMN	FORMAT	SELECTION DESCRIPTION
Card 1	1-80	8A10	Run Name
Card 2	3	11	1= Log Transform, 0 = no
	4	11	1= Average variogram option: yes 0 = " " " " ; no
	5	11	1= Punch for NEWPLOT, 0 = no
	6	11	1= Moment center, 0 = $\chi(h)$
	7	11	1=Auto Scale, 0 = no
	8	11	2= divide by variance 1= divide by mean square 0 =neither
	9	11	1- tape input, 0 = system
	10	11	No. of directions (max. 8)
	11-20	F10.0	Class interval: 20 ×

	41-50	F10.0	class = max range Max vertical value to allow
Card 3	You will need one of these cards for each Direction		
	1-10	F10.0	Direction
			0. = E-W 45. = NE-SW
			90. = N-S -45. = NW-SE
	11-20	F10.0	Window in degrees
Card 4	1	I1	1,2, or 3 (specifies order of occurrence of value being Kriged on data line)
	2	I1	1,2, or 3 (specifies occur. of y coord. ie. Northing)
	3	I1	1,2 or 3 (specifies occur. of x coord. ie. Easting)
	4	I1	1=flags definition of rectangular sub-area to confine variogram to follow: 0=no sub-area defined
	11-20	F10.0	Minimum Northing (y)
	21-30	F10.0	Maximum Northing
	31-40	F10.0	Minimum Easting (x)
	41-50	F10.0	Maximum Easting
Card 5	1-80	8A10	Format of input on system enclosed in (paren.)
Card 6	1-80	8A10	Title of Variogram for output
Card 7→n	Data. . .		

Each data point is represented by its own data card (7→n). Three values are being read in from each of these. They are the Easting, Northing,

and the value to be kriged in that order. Other numbers may exist on the data line as well, but formatting should ensure that just the three numbers being used are read in.

The output from the program is written to a file named OUTPUT and also to one named TAPE8. OUTPUT includes information on the variogram(s) just run and a character built plot of the variogram which comes off the line printer and provides a good, if not pretty, variogram plot. To get a copy of the results printed on an actual plotter, TAPE8 provides unformatted plotting output. NEWPLOT, the next program described, will produce a ZETA plot file from the TAPE8 file, which in turn can be sent directly to a ZETA plotter for draft-quality plots.

PLOTTING PROGRAM NEWPLOT

Before running this program, output file TAPE8 discussed above must be modified. To run several variograms at once with GAMM8B, TAPE8 must contain plotting information for each variogram, each being separated by a blank card (line). For each plot, one card free-formatted with the following data must be present:

- Variogram value at the top of the graph
- Variogram value at the bottom of the graph
- Variogram value at the right of the graph
- Variogram value at the left of the graph
- Length of the x axis (distance) in inches
- Length of the y axis (γ) in inches

These values should be separated by spaces but contained within one line

and should be listed in order. Do not leave a blank card between this card and the TAPE8 group (plot). Each plot should have one of these cards at the top. Leave only one blank card between the end of one group and this top card for the following one.

To then run NEWPLTB on the CDC (Control Data Dual Cyber) you will need to first execute the following command (this procedure as of May, 1985):

```
LDSET,PRESET=ZERO,LIB=ZETLIBF
```

The computer then prompts you with:

```
LDR>
```

You then type:

```
NEWPLTB,TAPE8
```

The program then generates a file called PLOTR, which can be sent to a ZETA plotter by issuing the command:

```
DISPOSE PLOT=PLOTR,ID=id site
```

where *id site* is some printing site identification number, eg. Taylor Hall is 44.

Fitting A Curve. As mentioned in the section THE VARIOGRAM, and as shown in figure D2, you will need to fit a curve to the variogram. This curve will be the weighting curve that the kriging program will actually use. The kriging program presented in this appendix uses the so-called *spherical model* for this curve. To fit the curve, three characteristic measurements of the variogram plot must be estimated. These are the *nugget value*, *Co*; the *C value*; and the *range*, *a*. To read these values off the variogram, the variogram must first be plotted.

The line-printed plot will do (this can be done on a terminal screen).

Figure D5 illustrates the procedure. Draw a straight line along the slope described by the first three or four points. Then draw another straight line across the points that indicate a plateau (its normal for this plateau to slope back down at the far end of the x axis.) This second line should be nearly horizontal and should pass through or near the highest γ values on the graph. If the shape of your variogram does not comply with this procedure and does not resemble those shown in the various figures in this appendix, then the variogram needs to be redone choosing a different class interval. The y-intercept of the first sloping line is the nugget value (C_0) and can be thought of as an indication of the accuracy of the data. Ideally this intercept should intersect the origin, though it almost never does. The y-intercept of the horizontal line is the sill value which is given as $C + C_0$, C being the difference between the sill and the nugget values. The sill indicates the maximum variance of the data. The range (a) is calculated as $3/2$ times the distance (measured off the x axis) from the origin of the intersection of the two lines which have been traced.

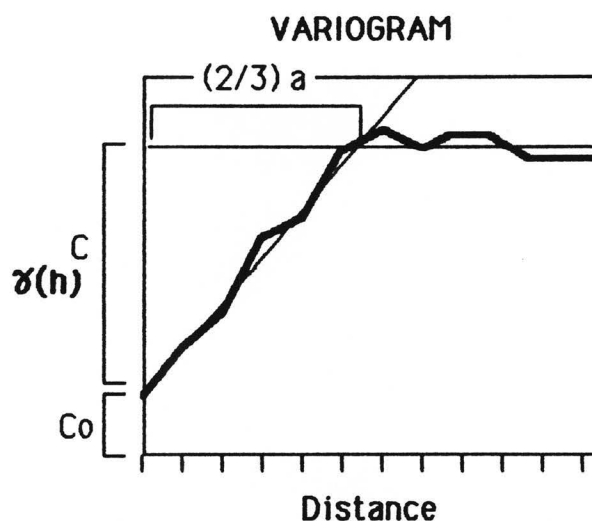


Figure D5. Obtaining C , C_0 and a by construction. The sill is given by $C + C_0$.

These three values alone are used to describe the fitted curve. Record these for future use in the kriging program. NEWPLOTB can plot the fitted curve on the variogram by further adding the following cards to the end of each data group on TAPE8:

CARD	COLUMN	FORMAT	DESCRIPTION
Card1	Blank		
Card 2	1	A1	Type a "3" in this column
Card 3	1-10	F10.0	Nugget value (C_0)
	11-20	F10.0	Sill value (C)
	21-30	F10.0	Range (a)

The Anisotropy Ratio. If several variograms have been run in different directions as discussed in previous sections, the *anisotropy*

ratio can be calculated using the sill values. The ratio is given by the ratio of the maximum directional sill value to the minimum directional sill value. The anisotropy ratio is used by the kriging program. Since the program only allows integers for this ratio, if the ratio is less than 1.5, then the anisotropy should be ignored.

THE KRIGING PROGRAM NEWUKB

The goal of constructing a variogram is to obtain the nugget, sill and range values for the nondirectional variogram, and also the anisotropy ratio if it is not less than 1.5. These will now be used as part of the input deck of the kriging program.

Kriging produces a rectangular grid of rows and columns. The centers of the blocks in this grid are the points for which the kriging program calculates values. One must therefore decide what size blocks to use. Considerations include how much resolution is desired, how many data points there are, and how much computer time can be dedicated to the procedure. In general, smaller blocks yield better resolution, but also require more computer time. However, a block will not be kriged with less than two points. If there are less than two points within the radius of search you specify, no value will be generated for that block.

The theory of kriging has been covered in the previous sections detailing the construction of the variogram. The procedure for executing the kriging program NEWUKB follows the same form as the previous programs. Here is the form for the input deck:

CARD	COLUMN	FORMAT	SELECTION DESCRIPTION
Card 1	1-5	I5	Starting Column
	6-10	I5	Ending Column
	11-15	I5	Starting Row
	16-20	I5	Ending Row
	21-30	F10.0	Max. Northing (Row 1)
	31-40	F10.0	Min. Easting (Col 1)
	41-50	F10.0	Block Width-North direction
	51-60	F10.0	Block Width-East direction
Card 2	1-5	F5.0	Max. dist. from block for hole to be included in kriging. This should be about half the distance across the data field.
	6-10	I5	Max. no. of points to be used for each block (≤ 50 , usually 10 is enough).
	11	I1	Debug output (1=yes, 0=no)
	12	I1	List drill holes (1=yes, 0=no)
Card 3	Card three is for mining applications and is used for kriging mineral-grade. You should set everything except the anisotropy ratio equal to zero. The ratio should be set at one. Card four, however is identical and was originally intended for vein thickness. For thickness, substitute the value you want to krig, eg. hydraulic heads. You should select the appropriate settings for card four. Bear in mind that in the output, the kriged values will be printed under the heading "thickness".		
	1-10	F10.0	Nugget Value (Co)
	11-20	F10.0	Sill Value (C)
	21-30	F10.0	Range (a)
	31-40	F10.0	Angle of anisotropy (0 if none) 0. = E-W 45. = NE-SW 90. = N-S -45. = NW-SE
	41-50	F10.0	Anisotropy Ratio (1 if none)

	51-60	F10.0	Vertical anisotropy (for mining, set to 0 for 2D)
Card 4	Same as Card 3. Use this card for you kriging values.		
Card 5	Format Card for your data (bracketted by paren.). You must format for the following variables in order (use tab format if necessary): Well identification X coordinate (Easting) Y coordinate (Northing) Z coordinate (supply a dummy variable, blanks should work; however, if you have problems, you might have to supply zeros) Grade-thickness (another dummy) Thickness Value (this is the value you are kriging)		
Card	6 → n	Your Data (one data point per card).	

To then run NEWUKB on the CDC the following command needs to be executed (this procedure as of May, 1985):

LDSET,PRESET=ZERO

The computer than prompts with:

LDR>

You then type:

NEWUKB, *datafile*

Where *datafile* is the name given to the input deck. The program then generates a file called OUTPUT, which contains the kriged grid, giving the kriged value for the block center under the column "thickness" and the

row and column for each block center. These block-centers can be displayed by using the following program MAPWB, which provides a crude line-printed map of the output. However, it will not handle more than approximately 30 blocks across. The output file can be used as input for a plotting package (eg. CPS1).

MAPWB

For a line-printed map of the kriged output used the output file in conjunction with the following cards:

CARD	COLUMN	FORMAT	DESCRIPTION
Card 1	1-10	F10.0	North coord of upper edge (row1)
	11-20	F10.0	East coord of left edge (col 1)
	21-25	F5.0	Block size y dimension
	26-30	F5.0	Block size x dimension
	31-35	I5	Beginning col.no. of area to be mapped
	36-40	I5	Ending col.no. of area to be mapped
	41-45	I5	Beginning row.no. of area to be mapped
	46-50	I5	Ending row.no. of area to be mapped
Card 2	1-10	F10.0	Scale parameter to make variable1 an integer (usually 1)
	11-20	F10.0	Scale parameter to make variable2 an integer (usually 1)
	21-30	F10.0	Scale parameter to make variable3 an integer (usually 1)

Card 3 Format Card for your data (bracketted by paren.).
You must format for the following variables in
order (use tab format if necessary):

East coordinate (as produced by
 NEWUKB)

North coordinate (as produced by
 NEWUKB)

Variable 1 (thickness)

 Variable 2 (dummy)

 Variable 3 (dummy)

Card 4→ n Data (from output of NEWUKB)

APPENDIX E-CPS1

The following files were used to execute the CPS1 graphics subroutines on the Dual Cyber computer. The command file was used to submit the job batch, ie., run CPS1 as a job on a queue independant of an interactive user. The command file is written in a command language which is specific to the operating system (called Taurus) on the CDC at the University of Texas. The verb files contain the commands (which they call verbs) to execute the CPS1 subroutines.

CPS1 turned out to be computation intensive. The contouring programs, for example, required an average of 300 seconds of central processing time. The isometric plots (3-D) took even longer. The boldface print indicates file names or variables that can be changed depending on the desired options and which files the user desires to be used for data input, etc.. . Because of the complexity of using CPS1, these files are presented without explanation. For descriptions of their construction and use, the reader is refered to a Taurus manual for the command files, and a CPS1 manual for the verb files.

POSTING DATA

COMMAND FILE

READPF 4170 T400L50
 READPF 4170 POSTYRB
 RENAME T400L50 TAPE1
 RFL 220000
 CPS1 POSTYRB
 CPS1TRN,,PLOT.
 SKIPCC.
 EXIT.
 REVIEW L=DAY
 SAYEPF 4170 5682 DAY OUTPUT P400L50=PLOT
 ZAP

VERB FILE

\$YC JOB=1 \$END
WELL LOCATIONS FOR K400L50
 \$YC AOI=1,YMIN=3014056.1,YMAX=3237605.5,XMIN=450501.3,XMAX=718654.3,XINC=1.,YINC=1. \$END
 \$YC PDEF=1,XSCL=23000,YSCL=23000,IOP=0 \$END
 \$YC POLY=1,NPC=27,ISMB=0,ISCL=1,IFMT=2 \$END
 (20X,F20.1,T1,F20.1)
 3014594.7 451204.7
 3014007.3 675797.4
 3055372.3 674940.5
 ect.....
 3112021.3 475036.1
 3112309.0 450669.8
 3014594.7 451204.7
 \$YC FENC=1,NPC=27,IFMT=2,ISCL=1 \$END
 (20X,F20.1,T1,F20.1)
 3014594.7 451204.7
 3014007.3 675797.4
 3055372.3 674940.5
 ect.....
 3112021.3 475036.1
 3112309.0 450669.8
 3014594.7 451204.7
 \$CD FILD=0 \$END
 \$YC LGPT=1,ISMB=-6,SIZE=0.06,LUN=1,MODE=2 \$END
 \$CD FILD=1,SIZE=0.06,IPSN=1,NCLS=5 \$END
 (19X,2F10.0,T1,I1,T4,5A1)
 \$YC NWPG=1 \$END
 \$YC STOP=1 \$END

CONTOURING

COMMAND FILE

```

READPF 5576 KRGCONB KRG350
RENAME KRG350 TAPE1
REWALLX
RFL,250000
CPS1 KRGCONB
CPS1TRN,,PLTCNT.
SKIPCC.
EXIT.
REVIEW,L=DAYCNT
SAVEPF 5576 8908 DAYCNT,K350L50=PLTCNT,OUTKRGB=OUTPUT
ZAP

```

VERB FILE

```

$YC JOB=1 $END
K250L30
$YC AOI=1,XMIN=450501.3,XMAX=718654.3,YMIN=3014056.1,YMAX=
3237605.5,XINC=5000.,YINC=5000. $END
$YC MGRD=1,NCP=0,IFMT=2,LUN=1 $END
(19X,2F10.0,T56,F8.0)
$YC STAT=1 $END
$YC PDEF=1,XSCL=23000.,YSCL=23000.,MODE=1 $END
$YC POLY=1,NPC=27,ISMB=0,ISCL=1,IFMT=2 $END
(20X,F20.1,T1,F20.1)
    3014594.7    451204.7
    3014007.3    675797.4
    3055372.3    674940.5
    ect...
    3112021.3    475036.1
    3112309.0    450669.8
    3014594.7    451204.7
$YC FENC=1,NPC=27,IFMT=2,ISCL=1 $END
(20X,F20.1,T1,F20.1)
    3014594.7    451204.7
    3014007.3    675797.4
    3055372.3    674940.5
    ect...
    3112021.3    475036.1
    3112309.0    450669.8
    3014594.7    451204.7
$CD FILD=0 $END
$YC CTYP=1,NREF=0,BOLD=0,DIS1=4,DIS2=2,HSR=0,LABR=1,SIZE=.08 $END
$YC CLY4=1,ZIDA=1,CLMN=-5000.,DEL=-500.,NCY=30,ICRT=1 $END
$YC STOP=1 $END

```

ISOMETRIC PROJECTIONCOMMAND FILE

READPF 5576 YRBISOC KRG350
RENAME KRG350 TAPE1
REWALLX
RFL,371000
CPS1 YRBISOC
CPS1TRN,,I350L50
SKIPCC
EXIT
REVIEW,L=DAYISO
SAYEPF 5576 8908 DAYISO,I350L50,OUTISO=OUTPUT
ZAP

VERB FILE

\$YC JOB=1 \$END
LOOKING NORTHEAST,DEPTH TO 250 DEG F
\$YC AOI=1,YMIN=3014056.1,YMAX=3237605.5,XMIN=450501.3,XMAX=718654.3,
XINC=8000.,YINC=8000. \$END
\$YC MGRD=1,IFMT=2,LUN=1,NCP=0,NZ=1,IANL=1 \$END
(18X,F10.0,F10.0,T56,F8.0)
\$YC STAT=1 \$END
\$YC PDEF=1,XSCL=50000.,YSCL=50000. \$END
\$YC RFIT=1,NREF=1,ZIDA=1,ITYP=1,ZIDB=0 \$END
\$YC ISOM=1,ZIDA=1,ZSCL=2500.,ZMIN=-20000.,ZMAX=1000.,
ISID=1,THET=65. \$END
\$YC NWPG=1 \$END
\$YC STOP=1 \$END

APPENDIX F – FLOW VELOCITY VECTOR FIELDS

The flow velocity vectors are listed for each triangular element. They appear by column, each column taking three lines in this listing. The columns are listed in order, as they appear from the left to right side of the mesh. The vectors for each element are listed in order starting at the bottom or basement of the mesh, and working toward the top. There are twenty-two elements in each column, and the vectors of the top two elements are all that appear on the third line for each column. The columns can be distinguished by this partial line.

SIMULATION ONE-Horizontal component

0.64E-12	0.50E-12	0.72E-10	0.53E-10	0.76E-10	0.79E-10	0.93E-10	0.80E-10	0.86E-10	0.87E-10
0.10E-09	0.69E-10	0.72E-10	0.50E-10	0.50E-10	0.45E-10	0.51E-10	0.18E-10	0.58E-09	0.20E-09
0.18E-09	0.00E+00								
0.11E-11	-0.15E-11	0.48E-10	0.27E-10	0.51E-10	0.34E-10	0.50E-10	0.34E-10	0.47E-10	0.38E-10
0.40E-10	0.31E-10	0.33E-10	0.21E-10	0.26E-10	0.15E-10	0.16E-10	0.98E-11	0.32E-09	0.16E-09
0.15E-09	0.00E+00								
0.26E-11	0.34E-11	0.53E-11	-0.32E-11	0.25E-10	-0.11E-10	0.38E-10	0.50E-11	0.33E-10	0.15E-10
0.34E-10	0.18E-10	0.27E-10	0.16E-10	0.17E-10	0.81E-11	0.97E-11	0.42E-11	0.18E-09	0.78E-10
0.13E-09	0.25E-10								
0.43E-12	0.11E-10	0.43E-11	-0.54E-11	0.61E-11	-0.27E-11	0.19E-10	0.24E-11	0.18E-10	0.72E-11
0.18E-10	0.83E-11	0.13E-10	0.51E-11	0.10E-10	0.33E-11	0.56E-11	0.20E-11	0.95E-10	0.50E-10
0.45E-10	0.00E+00								
-0.15E-10	0.49E-10	0.50E-11	-0.24E-11	0.65E-11	-0.44E-11	0.36E-10	-0.41E-11	0.29E-10	0.55E-11
0.22E-10	0.31E-11	0.15E-10	0.35E-11	0.82E-11	0.12E-11	0.26E-11	0.32E-12	0.10E-10	-0.89E-11
0.14E-10	0.00E+00								
-0.84E-10	0.16E-09	-0.12E-11	-0.11E-10	-0.14E-11	-0.65E-11	0.51E-10	0.28E-12	0.35E-10	0.24E-11
0.27E-10	0.90E-12	0.14E-10	-0.63E-12	0.76E-11	-0.26E-11	0.12E-11	-0.13E-11	-0.41E-10	-0.42E-10
-0.38E-10	-0.54E-11								
-0.32E-09	0.57E-09	-0.73E-10	-0.40E-10	-0.40E-10	0.17E-10	0.10E-10	-0.18E-11	0.28E-10	-0.52E-11
0.21E-10	-0.63E-11	0.84E-11	-0.79E-11	0.19E-11	-0.11E-10	-0.52E-11	-0.77E-11	-0.20E-09	-0.14E-09
-0.13E-09	0.31E-10								
-0.97E-09	0.79E-09	-0.96E-11	-0.14E-11	-0.16E-09	-0.43E-10	-0.11E-09	0.41E-10	0.48E-10	-0.45E-12
0.33E-10	-0.12E-10	0.14E-10	-0.14E-10	0.39E-11	-0.21E-10	-0.12E-10	-0.22E-10	-0.58E-09	-0.59E-09
-0.48E-09	-0.52E-10								
-0.78E-09	0.41E-09	-0.73E-11	-0.11E-11	-0.34E-11	0.10E-11	-0.52E-10	0.12E-09	0.11E-09	0.55E-10
0.92E-10	0.32E-10	0.49E-10	0.10E-10	0.25E-10	-0.27E-10	-0.87E-11	-0.51E-10	-0.14E-08	-0.26E-08
-0.20E-08	-0.56E-08								
-0.52E-09	0.68E-10	-0.68E-11	0.42E-11	0.11E-11	0.25E-11	0.41E-10	0.15E-09	0.35E-09	0.37E-09
0.35E-09	0.30E-09	0.29E-09	0.26E-09	0.24E-09	0.22E-09	0.17E-09	0.15E-09	0.45E-08	0.25E-08
0.22E-08	-0.58E-08								
-0.46E-09	-0.20E-09	-0.22E-11	0.69E-11	0.18E-11	0.34E-11	0.17E-11	0.82E-13	-0.13E-09	0.97E-10
0.69E-09	0.79E-09	0.69E-09	0.86E-09	0.68E-09	0.79E-09	0.60E-09	0.73E-09	0.20E-07	0.24E-07
0.20E-07	0.31E-07								
-0.44E-09	-0.40E-09	0.76E-12	-0.90E-12	-0.41E-11	0.55E-11	-0.17E-11	0.25E-11	0.12E-10	0.89E-10
-0.89E-09	-0.47E-10	-0.28E-09	0.13E-08	0.64E-09	0.66E-09	0.50E-09	0.58E-09	0.16E-07	0.22E-07
0.19E-07	0.34E-07								
-0.45E-09	-0.61E-09	0.20E-12	-0.14E-10	-0.73E-11	-0.13E-10	-0.15E-10	-0.65E-13	-0.47E-11	0.14E-12
-0.18E-09	0.43E-10	-0.68E-09	-0.13E-09	-0.50E-09	-0.65E-09	-0.45E-09	-0.66E-09	-0.18E-07	-0.28E-07
-0.22E-07	-0.51E-07								
-0.27E-09	-0.76E-09	-0.42E-11	-0.15E-10	-0.10E-10	-0.18E-10	-0.14E-10	-0.22E-10	-0.23E-10	-0.82E-11

-0.12E-10	-0.28E-11	-0.31E-09	0.31E-10	-0.55E-09	-0.64E-09	-0.44E-09	-0.52E-09	-0.15E-07	-0.19E-07
-0.17E-07	-0.21E-07								
-0.11E-09	-0.88E-09	-0.41E-11	-0.16E-10	-0.12E-10	-0.19E-10	-0.15E-10	-0.23E-10	-0.19E-10	-0.23E-10
-0.24E-10	-0.53E-11	-0.24E-09	0.52E-10	-0.42E-09	-0.51E-09	-0.35E-09	-0.42E-09	-0.12E-07	-0.14E-07
-0.13E-07	-0.23E-07								
0.54E-10	-0.91E-09	-0.20E-11	-0.87E-11	-0.62E-11	-0.89E-11	-0.68E-11	-0.91E-11	-0.66E-11	-0.82E-11
-0.71E-11	-0.42E-11	-0.25E-09	0.77E-10	-0.31E-09	-0.32E-09	-0.21E-09	-0.22E-09	-0.65E-08	-0.71E-08
-0.62E-08	-0.71E-08								
0.13E-08	-0.74E-09	-0.26E-11	-0.62E-11	-0.39E-11	-0.83E-11	-0.49E-11	-0.80E-11	-0.55E-11	-0.51E-11
-0.38E-11	-0.32E-11	-0.20E-09	0.98E-11	-0.92E-10	0.18E-10	0.28E-10	0.81E-10	0.18E-08	0.43E-08
0.37E-08	0.12E-07								
0.25E-08	-0.62E-09	-0.18E-11	-0.19E-11	-0.14E-11	0.56E-12	0.17E-11	-0.12E-12	0.20E-11	-0.19E-11
-0.19E-12	-0.27E-11	-0.12E-09	0.18E-10	-0.44E-10	0.27E-10	0.16E-09	0.17E-09	0.39E-08	0.76E-08
0.65E-08	0.23E-07								
0.28E-08	-0.60E-09	-0.37E-11	-0.30E-11	-0.19E-11	-0.23E-11	-0.21E-11	-0.75E-11	-0.43E-11	-0.19E-11
-0.16E-11	-0.11E-11	-0.85E-10	0.73E-10	-0.62E-10	0.60E-10	0.14E-09	0.10E-09	0.31E-08	0.48E-08
0.55E-08	0.00E+00								
0.31E-08	-0.48E-09	-0.19E-11	-0.22E-11	-0.20E-11	-0.48E-12	-0.43E-12	0.33E-11	0.30E-11	0.74E-12
0.61E-12	-0.87E-12	-0.49E-10	0.50E-10	-0.24E-10	0.41E-10	0.18E-09	0.55E-10	0.23E-08	0.42E-08
0.34E-09	0.00E+00								
0.30E-08	-0.37E-09	-0.42E-14	0.14E-11	0.13E-11	0.19E-11	0.17E-11	0.15E-11	0.13E-11	0.86E-12
0.72E-12	0.50E-13	0.18E-11	0.11E-09	-0.22E-10	0.90E-10	0.51E-09	0.52E-09	0.43E-09	0.36E-09
0.22E-08	0.54E-08								
0.28E-09	-0.20E-09	-0.23E-14	0.21E-12	0.19E-12	0.29E-12	0.26E-12	0.26E-12	0.23E-12	0.21E-12
0.18E-12	0.11E-13	0.74E-11	0.83E-10	-0.76E-11	0.11E-09	0.68E-09	0.70E-09	0.53E-09	0.61E-09
0.14E-07	0.19E-07								
0.28E-08	-0.37E-10	-0.43E-15	-0.18E-12	-0.17E-12	-0.35E-12	-0.32E-12	-0.43E-12	-0.38E-12	-0.64E-12
-0.55E-12	-0.83E-12	-0.85E-11	-0.38E-10	-0.29E-10	0.20E-10	0.62E-09	0.59E-09	0.49E-09	0.44E-09
0.11E-07	0.71E-09								
0.27E-08	0.15E-09	0.17E-14	-0.76E-12	-0.70E-12	-0.17E-11	-0.16E-11	-0.26E-11	-0.24E-11	-0.37E-11
-0.33E-11	-0.49E-11	-0.43E-11	-0.62E-11	-0.51E-11	0.34E-12	0.62E-09	0.61E-09	0.48E-09	0.48E-09
0.11E-07	0.15E-07								
0.24E-08	0.33E-09	0.39E-14	0.44E-12	0.41E-12	0.84E-12	0.79E-12	0.11E-11	0.98E-12	0.12E-11
0.10E-11	0.12E-11	0.10E-11	0.98E-12	0.83E-12	0.27E-12	0.51E-09	0.50E-09	0.41E-09	0.40E-09
0.10E-07	0.26E-07								
0.22E-03	0.44E-09	0.51E-14	0.17E-11	0.16E-11	0.33E-11	0.30E-11	0.43E-11	0.39E-11	0.52E-11
0.45E-11	0.58E-11	0.48E-11	0.54E-11	0.45E-11	0.23E-12	0.43E-09	0.41E-09	0.33E-09	0.30E-09
0.77E-09	-0.13E-07								
0.21E-08	0.50E-09	0.57E-14	0.25E-11	0.22E-11	0.46E-11	0.41E-11	0.59E-11	0.53E-11	0.76E-11
0.65E-11	0.10E-10	0.84E-11	0.18E-10	0.15E-10	0.20E-12	0.38E-09	0.39E-09	0.31E-09	0.32E-09
0.82E-09	0.26E-07								

0.20E-09	0.04E-09	0.65E-10	0.33E-11	0.29E-11	0.58E-11	0.53E-11	0.71E-11	0.65E-11	0.82E-11
0.72E-11	0.84E-11	0.71E-11	0.89E-11	0.50E-11	0.20E-12	0.39E-09	0.38E-09	0.30E-09	0.30E-09
0.78E-08	0.14E-07								
0.18E-08	0.65E-09	0.75E-14	0.30E-11	0.45E-11	0.89E-11	0.80E-11	0.11E-10	0.96E-11	0.12E-10
0.11E-10	0.13E-10	0.11E-10	0.48E-11	0.42E-11	0.18E-12	0.35E-09	0.32E-09	0.26E-09	0.22E-09
0.56E-09	0.00E+00								
0.16E-08	0.74E-09	0.84E-14	0.85E-11	0.75E-11	0.15E-10	0.13E-10	0.18E-10	0.15E-10	0.19E-10
0.16E-10	0.95E-11	0.81E-11	0.16E-12	0.31E-09	0.30E-09	0.25E-09	0.26E-09	0.21E-09	0.21E-09
0.54E-09	0.00E+00								
0.14E-08	0.77E-09	0.31E-12	0.13E-10	0.12E-10	0.23E-10	0.21E-10	0.31E-10	0.26E-10	0.14E-10
0.12E-10	0.24E-11	0.14E-09	0.11E-10	0.38E-09	0.37E-09	0.31E-09	0.32E-09	0.26E-09	0.26E-09
0.68E-08	0.00E+00								
0.13E-09	0.77E-09	0.12E-10	0.19E-10	0.17E-10	0.26E-10	0.23E-10	0.83E-11	0.73E-11	0.18E-11
0.85E-10	0.20E-09	0.18E-09	0.13E-10	0.46E-09	0.45E-09	0.38E-09	0.39E-09	0.32E-09	0.36E-09
0.94E-08	0.26E-07								
0.11E-08	0.70E-09	0.11E-10	0.92E-11	0.87E-11	0.70E-11	0.61E-11	0.30E-11	0.26E-11	0.13E-11
0.62E-10	0.13E-10	0.47E-09	0.44E-09	0.37E-09	0.37E-09	0.31E-09	0.30E-09	0.24E-09	0.25E-09
0.63E-08	0.48E-08								
0.70E-09	0.53E-09	0.84E-11	0.58E-11	0.52E-11	0.26E-11	0.23E-11	0.16E-11	0.14E-11	0.88E-12
0.43E-10	0.95E-11	0.33E-09	0.31E-09	0.26E-09	0.25E-09	0.21E-09	0.20E-09	0.16E-09	0.14E-09
0.36E-08	0.33E-08								
0.39E-09	0.39E-09	0.61E-11	0.43E-11	0.38E-11	0.19E-11	0.16E-11	0.12E-11	0.10E-11	0.63E-12
0.31E-10	0.57E-11	0.20E-09	0.17E-09	0.15E-09	0.12E-09	0.10E-09	0.86E-10	0.69E-10	0.48E-10
0.12E-08	0.36E-10								
0.13E-09	0.29E-09	0.45E-11	0.57E-11	0.51E-11	0.20E-11	0.18E-11	0.11E-11	0.98E-12	0.46E-12
0.23E-10	0.25E-11	0.88E-10	0.72E-10	0.61E-10	0.50E-10	0.43E-10	0.32E-10	0.26E-10	0.14E-10
0.34E-09	0.17E-11								

SIMULATION ONE-Vertical component

-0.44E-13	-0.45E-13	-0.37E-11	-0.82E-11	-0.68E-11	-0.50E-11	0.46E-11	-0.11E-11	-0.10E-11	0.81E-12
0.21E-10	0.53E-11	0.15E-10	0.81E-11	0.11E-10	0.95E-11	0.50E-10	0.11E-10	0.35E-10	0.15E-10
0.25E-10	0.16E-10								
-0.26E-13	-0.74E-13	0.91E-11	0.18E-12	0.54E-11	-0.24E-11	0.40E-11	-0.13E-11	0.53E-11	0.26E-11
0.53E-11	0.38E-11	0.70E-11	0.52E-11	0.88E-11	0.65E-11	0.10E-10	0.65E-11	0.13E-10	0.84E-11
0.15E-10	0.18E-10								
-0.48E-13	-0.33E-13	0.17E-13	-0.15E-12	0.41E-11	-0.10E-10	0.49E-11	-0.53E-11	0.34E-11	-0.25E-11
0.38E-11	0.16E-12	0.41E-11	0.28E-11	0.39E-11	0.26E-11	0.46E-11	0.27E-11	0.59E-11	0.38E-11
0.79E-11	0.58E-11								
0.11E-14	0.31E-12	-0.14E-12	-0.50E-12	0.56E-11	-0.11E-11	0.29E-11	-0.57E-11	0.19E-11	-0.51E-11
0.12E-11	-0.22E-11	0.79E-12	-0.22E-11	0.15E-11	-0.81E-12	0.18E-11	-0.49E-12	0.22E-11	0.11E-11
0.35E-11	0.14E-11								

0.39E-12	0.19E-11	-0.40E-12	-0.66E-12	0.22E-13	-0.36E-12	0.41E-11	-0.14E-10	0.35E-11	-0.88E-11
0.19E-13	-0.66E-11	0.27E-12	-0.36E-11	-0.11E-11	-0.20E-11	-0.12E-11	-0.12E-11	-0.14E-11	-0.43E-12
0.24E-12	0.13E-11								
0.21E-11	0.83E-11	-0.65E-12	-0.88E-12	-0.34E-12	-0.51E-12	0.91E-11	-0.16E-10	0.51E-11	-0.14E-10
0.33E-11	-0.82E-11	0.12E-11	-0.52E-11	0.37E-13	-0.25E-11	-0.34E-12	0.58E-13	0.55E-13	0.21E-11
0.15E-12	0.52E-11								
0.69E-11	0.23E-10	-0.86E-12	-0.75E-13	-0.49E-12	0.16E-11	0.13E-10	0.79E-11	0.11E-10	-0.11E-10
0.60E-11	-0.65E-11	0.41E-11	-0.46E-11	0.23E-11	-0.16E-11	0.71E-12	0.17E-11	0.21E-11	0.69E-11
0.28E-11	0.14E-10								
0.26E-10	0.27E-10	0.25E-14	0.74E-11	0.18E-11	0.58E-11	0.66E-13	0.45E-11	0.14E-10	-0.18E-10
0.96E-11	-0.13E-10	0.71E-11	-0.84E-11	0.44E-11	-0.46E-11	0.20E-11	-0.19E-11	0.71E-11	0.91E-11
0.13E-10	0.44E-10								
0.10E-10	0.10E-10	0.80E-11	-0.13E-10	0.64E-11	0.15E-10	0.48E-11	0.10E-10	0.17E-10	-0.21E-10
0.99E-11	-0.21E-10	0.45E-11	-0.23E-10	-0.18E-12	-0.24E-10	-0.21E-11	-0.28E-10	0.78E-11	-0.82E-10
0.42E-10	-0.17E-09								
0.11E-10	0.11E-10	0.15E-10	0.23E-10	0.12E-10	0.14E-10	0.11E-10	0.14E-10	0.30E-10	0.49E-10
0.10E-10	-0.10E-10	-0.39E-11	-0.21E-10	-0.18E-10	-0.25E-10	-0.33E-10	-0.44E-10	-0.85E-10	-0.23E-09
-0.18E-09	-0.62E-09								
0.13E-10	0.13E-10	0.86E-11	0.11E-10	0.16E-10	0.18E-10	0.15E-10	0.13E-10	0.18E-12	0.90E-11
0.33E-10	0.95E-10	0.12E-10	0.15E-09	-0.21E-10	0.41E-10	-0.46E-10	0.52E-10	-0.23E-09	0.15E-10
-0.63E-09	0.12E-10								
0.14E-10	0.14E-10	0.12E-10	0.11E-10	0.67E-11	0.12E-10	0.11E-10	0.16E-10	0.97E-11	0.13E-10
0.19E-12	0.92E-12	0.26E-12	0.18E-11	0.48E-10	0.68E-10	0.49E-10	0.12E-09	0.14E-10	0.49E-09
0.98E-11	0.13E-08								
0.14E-10	0.14E-10	0.12E-10	0.95E-11	0.12E-10	0.10E-10	0.58E-11	0.10E-10	0.10E-10	0.15E-10
0.10E-11	0.69E-11	0.19E-11	0.22E-11	0.56E-10	0.19E-11	0.12E-09	0.14E-11	0.49E-09	-0.12E-11
0.13E-08	-0.34E-09								
0.15E-10	0.15E-10	0.98E-11	0.87E-11	0.11E-10	0.87E-11	0.11E-10	0.89E-11	0.53E-11	0.86E-11
0.55E-11	0.13E-10	0.22E-11	0.64E-11	-0.74E-14	-0.17E-13	-0.37E-11	-0.19E-10	-0.56E-11	-0.13E-09
-0.35E-09	-0.46E-09								
0.15E-10	0.15E-10	0.87E-11	0.80E-11	0.90E-11	0.80E-11	0.93E-11	0.85E-11	0.88E-11	0.83E-11
0.43E-11	0.70E-11	0.65E-11	0.85E-11	-0.25E-13	-0.22E-13	-0.24E-10	-0.26E-10	-0.13E-09	-0.17E-09
-0.46E-09	-0.59E-09								
0.14E-10	0.15E-10	0.81E-11	0.72E-11	0.84E-11	0.76E-11	0.87E-11	0.82E-11	0.86E-11	0.82E-11
0.72E-11	0.80E-11	0.64E-11	0.11E-10	-0.29E-13	-0.16E-13	-0.29E-10	-0.25E-10	-0.17E-09	-0.19E-09
-0.60E-09	-0.62E-09								
0.15E-10	0.15E-10	0.74E-11	0.67E-11	0.78E-11	0.64E-11	0.85E-11	0.75E-11	0.85E-11	0.86E-11
0.82E-11	0.84E-11	0.35E-11	0.48E-11	-0.18E-13	0.20E-11	-0.29E-10	-0.33E-11	-0.19E-09	-0.88E-10
-0.52E-09	-0.27E-09								
0.15E-10	0.15E-10	0.69E-11	0.70E-11	0.65E-11	0.70E-11	0.77E-11	0.73E-11	0.89E-11	0.80E-11
0.85E-11	0.81E-11	0.48E-11	0.53E-11	0.20E-11	0.28E-11	-0.12E-10	-0.78E-12	-0.93E-10	-0.22E-10
-0.27E-09	0.15E-09								

0.15E-10	0.14E-10	0.71E-11	0.70E-11	0.71E-11	0.71E-11	0.74E-11	0.68E-11	0.83E-11	0.86E-11
0.81E-11	0.82E-11	0.54E-11	0.57E-11	0.28E-11	0.36E-11	-0.22E-11	0.29E-11	-0.25E-10	-0.60E-11
0.15E-09	0.95E-10								
0.16E-10	0.16E-10	0.73E-11	0.72E-11	0.72E-11	0.76E-11	0.70E-11	0.80E-11	0.88E-11	0.82E-11
0.83E-11	0.81E-11	0.59E-11	0.63E-11	0.37E-11	0.46E-11	-0.12E-11	-0.16E-10	-0.88E-11	0.23E-10
0.70E-10	0.17E-10								
0.16E-10	0.16E-10	0.75E-11	0.80E-11	0.78E-11	0.81E-11	0.83E-11	0.82E-11	0.85E-11	0.83E-11
0.84E-11	0.82E-11	0.66E-11	0.74E-11	0.48E-11	0.75E-11	-0.18E-10	-0.72E-11	0.11E-10	-0.11E-10
0.13E-10	-0.14E-09								
0.16E-10	0.17E-10	0.82E-11	0.84E-11	0.83E-11	0.84E-11	0.84E-11	0.85E-11	0.85E-11	0.85E-11
0.85E-11	0.84E-11	0.78E-11	0.84E-11	0.78E-11	0.11E-10	-0.45E-11	0.10E-10	-0.15E-10	0.17E-10
-0.19E-09	-0.73E-10								
0.17E-10	0.17E-10	0.85E-11	0.85E-11	0.85E-11	0.85E-11	0.86E-11	0.86E-11	0.86E-11	0.86E-11
0.86E-11	0.85E-11	0.87E-11	0.85E-11	0.11E-10	0.13E-10	0.12E-10	0.50E-11	0.19E-10	0.11E-10
-0.79E-10	-0.19E-09								
0.17E-10	0.17E-10	0.85E-11	0.80E-11	0.85E-11	0.78E-11	0.86E-11	0.76E-11	0.86E-11	0.74E-11
0.86E-11	0.71E-11	0.85E-11	0.69E-11	0.32E-11	0.64E-11	0.99E-11	0.11E-10	0.69E-11	0.15E-10
-0.17E-09	-0.65E-10								
0.17E-10	0.16E-10	0.80E-11	0.78E-11	0.78E-11	0.76E-11	0.76E-11	0.73E-11	0.74E-11	0.70E-11
0.71E-11	0.67E-11	0.67E-11	0.63E-11	0.64E-11	0.60E-11	0.48E-11	0.97E-11	0.12E-10	0.18E-10
-0.77E-10	0.17E-09								
0.16E-10	0.16E-10	0.78E-11	0.79E-11	0.75E-11	0.77E-11	0.73E-11	0.75E-11	0.70E-11	0.72E-11
0.67E-11	0.69E-11	0.63E-11	0.63E-11	0.60E-11	0.48E-11	0.74E-11	0.88E-11	0.13E-10	0.12E-10
0.17E-09	-0.36E-10								
0.16E-10	0.16E-10	0.79E-11	0.81E-11	0.77E-11	0.79E-11	0.74E-11	0.76E-11	0.72E-11	0.74E-11
0.69E-11	0.73E-11	0.63E-11	0.74E-11	0.49E-11	0.29E-11	0.49E-11	0.11E-10	0.67E-11	0.14E-10
-0.47E-10	0.55E-10								
0.16E-10	0.15E-10	0.81E-11	0.84E-11	0.78E-11	0.82E-11	0.76E-11	0.79E-11	0.74E-11	0.76E-11
0.73E-11	0.73E-11	0.74E-11	0.69E-11	0.12E-10	0.69E-11	0.68E-11	0.94E-11	0.77E-11	0.16E-10
0.46E-10	0.12E-09								
0.15E-10	0.15E-10	0.84E-11	0.99E-11	0.81E-11	0.95E-11	0.79E-11	0.91E-11	0.76E-11	0.86E-11
0.73E-11	0.82E-11	0.69E-11	0.43E-11	0.70E-11	0.62E-14	0.77E-11	-0.22E-11	0.10E-10	0.98E-12
0.11E-09	0.12E-10								
0.15E-10	0.15E-10	0.98E-11	0.12E-10	0.95E-11	0.12E-10	0.90E-11	0.11E-10	0.86E-11	0.10E-10
0.82E-11	0.57E-11	0.13E-10	0.47E-11	0.55E-12	0.40E-11	-0.87E-11	0.49E-12	-0.55E-11	-0.33E-11
0.72E-12	-0.77E-10								
0.15E-10	0.15E-10	0.12E-10	0.15E-10	0.12E-10	0.14E-10	0.11E-10	0.14E-10	0.10E-10	0.68E-11
0.17E-10	0.10E-10	0.63E-11	0.41E-11	-0.48E-13	0.41E-11	-0.65E-11	0.11E-12	-0.11E-10	-0.59E-11
-0.87E-10	-0.15E-09								
0.15E-10	0.13E-10	0.15E-10	0.15E-10	0.14E-10	0.17E-10	0.14E-10	0.75E-11	0.20E-10	0.14E-10
0.14E-10	0.17E-10	0.42E-11	0.10E-13	-0.58E-12	0.59E-11	-0.67E-11	0.78E-11	-0.12E-10	0.14E-10
-0.16E-09	0.12E-09								

0.13E-10	0.11E-10	0.14E-10	0.15E-10	0.16E-10	0.16E-10	0.22E-10	0.16E-10	0.19E-10	0.15E-10
0.18E-10	0.15E-10	0.23E-10	0.35E-11	0.78E-11	0.12E-10	0.12E-10	0.15E-10	0.15E-10	0.25E-10
0.11E-09	0.82E-10								
0.11E-10	0.10E-10	0.16E-10	0.19E-10	0.16E-10	0.14E-10	0.16E-10	0.14E-10	0.16E-10	0.14E-10
0.16E-10	0.14E-10	0.27E-10	0.10E-10	0.27E-10	0.15E-10	0.29E-10	0.19E-10	0.28E-10	0.23E-10
0.78E-10	0.76E-10								
0.10E-10	0.10E-10	0.14E-10	0.13E-10	0.14E-10	0.13E-10	0.14E-10	0.13E-10	0.14E-10	0.13E-10
0.19E-10	0.13E-10	0.23E-10	0.11E-10	0.26E-10	0.14E-10	0.27E-10	0.17E-10	0.27E-10	0.21E-10
0.71E-10	0.48E-10								
0.10E-10	0.11E-10	0.13E-10	0.14E-10	0.13E-10	0.11E-10	0.13E-10	0.11E-10	0.13E-10	0.11E-10
0.13E-10	0.12E-10	0.18E-10	0.15E-10	0.20E-10	0.19E-10	0.21E-10	0.19E-10	0.22E-10	0.20E-10
0.36E-10	0.35E-10								

SIMULATION TWO-Horizontal component

0.98E-13	0.37E-12	0.40E-10	0.38E-10	0.47E-10	0.43E-10	0.59E-10	0.48E-10	0.58E-10	0.52E-10
0.63E-10	0.46E-10	0.50E-10	0.31E-10	0.34E-10	0.24E-10	0.96E-09	0.41E-09	0.38E-09	0.14E-09
0.13E-09	0.00E+00								
-0.45E-12	0.23E-12	0.27E-10	-0.87E-11	0.23E-10	0.55E-11	0.24E-10	0.85E-11	0.23E-10	0.15E-10
0.19E-10	0.14E-10	0.16E-10	0.11E-10	0.15E-10	0.12E-10	0.48E-09	0.30E-09	0.28E-09	0.14E-09
0.13E-09	0.00E+00								
-0.17E-11	0.83E-12	-0.35E-11	-0.54E-10	-0.34E-11	-0.36E-10	0.12E-10	-0.15E-10	0.13E-10	-0.31E-12
0.17E-10	0.78E-11	0.16E-10	0.11E-10	0.13E-10	0.99E-11	0.35E-09	0.19E-09	0.25E-09	0.96E-10
0.15E-09	0.12E-10								
-0.58E-11	0.63E-11	-0.17E-10	-0.29E-10	-0.34E-12	-0.15E-10	0.97E-11	-0.32E-11	0.12E-10	0.37E-11
0.12E-10	0.73E-11	0.12E-10	0.66E-11	0.10E-10	0.69E-11	0.28E-09	0.16E-09	0.18E-09	0.94E-10
0.16E-10	0.00E+00								
-0.27E-10	0.48E-11	-0.16E-10	0.86E-11	0.59E-10	0.15E-10	0.55E-10	0.13E-10	0.44E-10	0.19E-10
0.34E-10	0.14E-10	0.25E-10	0.11E-10	0.15E-10	0.83E-11	0.29E-09	0.15E-09	0.14E-09	0.64E-10
0.87E-10	0.00E+00								
-0.66E-10	0.28E-10	-0.24E-10	0.35E-10	0.11E-09	0.85E-10	0.12E-09	0.46E-10	0.80E-10	0.34E-10
0.59E-10	0.19E-10	0.33E-10	0.11E-10	0.20E-10	0.99E-11	0.38E-09	0.19E-09	0.17E-09	0.11E-09
0.56E-10	0.41E-10								
-0.19E-09	0.85E-10	-0.78E-10	0.36E-10	-0.28E-10	0.28E-10	0.17E-09	0.74E-10	0.12E-09	0.63E-10
0.91E-10	0.39E-10	0.93E-10	0.28E-10	0.36E-10	0.20E-10	0.71E-09	0.27E-09	0.29E-09	0.12E-09
0.11E-09	-0.63E-10								
-0.41E-09	0.11E-09	-0.17E-09	0.99E-10	-0.15E-10	0.78E-10	0.13E-09	0.95E-10	0.12E-09	0.67E-10
0.67E-10	0.59E-10	0.72E-10	0.91E-10	0.62E-10	0.48E-10	0.15E-09	0.53E-09	0.55E-09	0.68E-10
0.14E-09	0.41E-10								

-0.47E-09	0.60E-10	-0.75E-11	0.18E-09	-0.42E-10	0.11E-09	-0.01E-09	0.70E-09	0.18E-09	0.15E-09
0.17E-09	0.14E-09	0.14E-09	0.13E-09	0.12E-09	0.11E-09	0.32E-08	0.12E-08	0.11E-08	-0.13E-08
-0.76E-09	-0.54E-08								
-0.45E-09	-0.28E-09	-0.76E-11	0.39E-11	-0.43E-10	-0.38E-10	-0.17E-09	0.96E-10	0.38E-09	0.38E-09
0.37E-09	0.33E-09	0.31E-09	0.29E-09	0.27E-09	0.26E-09	0.73E-08	0.61E-08	0.52E-08	0.25E-08
0.23E-08	-0.57E-08								
-0.34E-09	-0.34E-09	-0.52E-11	-0.48E-11	-0.89E-11	0.28E-11	-0.15E-09	0.53E-10	-0.12E-09	0.88E-10
0.68E-09	0.67E-09	0.59E-09	0.54E-09	0.47E-09	0.45E-09	0.12E-07	0.17E-07	0.13E-07	0.20E-07
0.17E-07	0.31E-07								
-0.33E-09	-0.50E-09	-0.68E-11	-0.74E-11	-0.50E-11	-0.38E-11	-0.88E-11	0.62E-11	-0.61E-10	0.16E-09
-0.23E-09	0.11E-08	0.43E-09	0.35E-09	0.33E-09	0.30E-09	0.84E-08	0.13E-07	0.11E-07	0.20E-07
0.17E-07	0.36E-07								
-0.33E-09	-0.74E-09	-0.95E-11	-0.25E-10	-0.17E-10	-0.20E-10	-0.15E-10	-0.17E-10	-0.20E-10	0.56E-11
-0.27E-09	0.18E-09	0.57E-10	-0.13E-09	-0.39E-10	-0.13E-09	-0.17E-08	-0.10E-07	-0.60E-08	-0.25E-07
-0.18E-07	-0.52E-07								
-0.29E-09	-0.83E-09	-0.14E-10	-0.45E-10	-0.33E-10	-0.37E-10	-0.31E-10	-0.28E-10	-0.73E-10	-0.81E-10
-0.29E-10	0.17E-10	-0.56E-13	-0.41E-11	-0.49E-10	-0.11E-09	-0.14E-08	-0.80E-08	-0.63E-08	-0.15E-07
-0.13E-07	-0.21E-07								
-0.39E-09	-0.82E-09	-0.17E-10	-0.49E-10	-0.33E-10	-0.63E-10	-0.49E-10	-0.26E-10	-0.77E-10	-0.11E-10
-0.14E-09	0.14E-11	0.12E-11	0.61E-16	0.69E-10	-0.50E-10	-0.53E-09	-0.59E-08	-0.46E-08	-0.11E-07
-0.10E-07	-0.23E-07								
-0.73E-09	-0.62E-09	-0.12E-10	-0.13E-10	-0.81E-11	-0.16E-10	-0.59E-11	-0.28E-10	-0.59E-10	0.79E-11
-0.91E-10	0.23E-09	0.32E-11	0.85E-12	0.86E-10	-0.91E-11	0.57E-09	-0.25E-08	-0.26E-08	-0.48E-08
-0.42E-08	-0.65E-08								
-0.57E-11	-0.47E-10	-0.38E-11	-0.41E-11	-0.18E-11	-0.61E-11	0.26E-11	-0.10E-10	0.17E-10	0.13E-10
0.80E-10	0.37E-09	0.44E-11	0.29E-11	0.39E-11	0.23E-11	0.29E-08	0.28E-08	0.13E-08	0.49E-08
0.42E-08	0.12E-07								
0.68E-09	0.54E-09	0.67E-11	0.13E-10	0.11E-10	0.15E-10	0.17E-10	0.14E-10	0.93E-10	-0.83E-11
0.14E-09	0.30E-11	0.42E-13	0.13E-13	0.16E-11	0.44E-11	0.45E-08	0.60E-08	0.23E-08	0.84E-08
0.69E-08	0.23E-07								
0.37E-09	0.77E-09	0.92E-11	0.21E-10	0.19E-10	0.29E-10	0.89E-10	0.99E-10	0.13E-09	0.71E-11
0.99E-10	-0.45E-10	-0.55E-09	-0.53E-09	-0.43E-09	0.49E-10	0.55E-08	0.67E-08	0.39E-08	0.66E-08
0.61E-08	0.00E+00								
0.53E-09	0.87E-09	0.13E-10	0.21E-10	0.19E-10	0.27E-10	0.83E-10	0.86E-10	0.79E-10	0.18E-10
0.25E-09	-0.50E-09	-0.62E-09	-0.96E-10	-0.92E-10	0.40E-10	0.52E-08	0.74E-08	0.47E-08	0.65E-08
0.51E-08	0.00E+00								
0.29E-09	0.70E-09	0.11E-10	0.17E-10	0.16E-10	0.20E-10	0.62E-10	0.67E-10	0.60E-10	0.41E-10
0.56E-09	-0.37E-09	-0.42E-09	-0.11E-09	-0.14E-09	0.11E-09	0.44E-08	0.89E-08	0.55E-08	0.81E-08
0.63E-08	0.53E-08								
0.93E-09	0.57E-09	0.93E-11	0.19E-10	0.18E-10	0.29E-10	0.88E-10	0.61E-10	0.55E-10	0.41E-10
0.58E-09	0.48E-11	-0.29E-09	-0.35E-09	-0.19E-10	0.14E-09	0.61E-09	0.55E-09	0.14E-07	0.14E-07
0.86E-08	0.18E-07								

0.83E-09	0.56E-09	0.92E-11	0.89E-11	0.82E-11	0.72E-11	0.22E-10	0.22E-10	0.21E-10	0.17E-10
0.25E-09	0.67E-11	0.19E-09	0.17E-09	0.33E-11	0.13E-10	0.40E-09	0.40E-09	0.11E-07	0.12E-07
0.10E-07	0.72E-08								
0.89E-09	0.52E-09	0.87E-11	0.19E-11	0.18E-11	-0.52E-11	-0.17E-10	0.47E-11	0.45E-11	0.15E-11
0.24E-10	0.95E-11	0.34E-09	0.76E-09	0.28E-12	0.20E-12	0.36E-09	0.36E-09	0.99E-08	0.13E-07
0.98E-08	0.16E-07								
0.73E-09	0.51E-09	0.85E-11	0.32E-11	0.36E-11	-0.38E-12	-0.12E-11	-0.82E-11	-0.77E-11	-0.25E-11
-0.39E-10	0.11E-10	0.40E-09	0.56E-09	0.48E-09	0.31E-09	0.26E-09	0.27E-09	0.77E-08	0.12E-07
0.90E-08	0.26E-07								
0.67E-09	0.52E-09	0.83E-11	0.51E-11	0.46E-11	0.12E-11	0.37E-11	-0.65E-11	-0.57E-11	0.22E-10
0.19E-10	0.29E-10	0.40E-09	0.57E-09	0.48E-09	0.80E-11	0.27E-09	0.19E-09	0.56E-08	0.91E-08
0.68E-08	-0.13E-07								
0.65E-09	0.53E-09	0.84E-11	0.63E-11	0.56E-11	0.29E-11	0.87E-11	0.74E-11	0.65E-11	0.24E-10
0.20E-10	0.30E-10	0.42E-09	0.60E-09	0.51E-09	0.38E-10	0.13E-08	0.14E-08	0.12E-08	0.56E-09
0.15E-07	0.26E-07								
0.68E-09	0.53E-09	0.87E-11	0.77E-11	0.70E-11	0.69E-11	0.64E-11	0.65E-11	0.20E-10	0.27E-10
0.24E-10	0.30E-10	0.43E-09	0.34E-10	0.12E-08	0.77E-09	0.67E-09	0.68E-09	0.56E-09	0.51E-09
0.13E-07	0.14E-07								
0.65E-09	0.54E-09	0.87E-11	0.91E-11	0.82E-11	0.89E-11	0.80E-11	0.88E-11	0.26E-10	0.36E-10
0.30E-10	0.96E-12	0.54E-09	0.12E-08	0.48E-12	0.24E-12	0.47E-09	0.45E-09	0.37E-09	0.34E-09
0.88E-08	0.00E+00								
0.61E-09	0.55E-09	0.87E-11	0.94E-11	0.84E-11	0.81E-11	0.72E-11	0.57E-11	0.50E-11	0.11E-12
0.20E-09	0.51E-09	0.20E-12	0.24E-12	0.47E-09	0.43E-09	0.36E-09	0.36E-09	0.29E-09	0.28E-09
0.74E-08	0.00E+00								
0.60E-09	0.55E-09	0.87E-11	0.12E-10	0.11E-10	0.15E-10	0.13E-10	0.15E-10	0.14E-10	0.35E-11
0.17E-09	0.33E-09	0.53E-11	0.21E-12	0.41E-09	0.42E-09	0.36E-09	0.38E-09	0.31E-09	0.31E-09
0.82E-08	0.00E+00								
0.60E-09	0.53E-09	0.86E-11	0.15E-10	0.14E-10	0.22E-10	0.20E-10	0.14E-10	0.13E-10	0.11E-10
0.94E-11	0.71E-11	0.35E-09	0.15E-10	0.53E-09	0.53E-09	0.46E-09	0.46E-09	0.38E-09	0.39E-09
0.10E-07	0.26E-07								
0.53E-09	0.47E-09	0.73E-11	0.68E-11	0.60E-11	0.48E-11	0.42E-11	0.21E-11	0.18E-11	0.78E-12
0.37E-10	0.15E-10	0.52E-09	0.49E-09	0.41E-09	0.40E-09	0.34E-09	0.32E-09	0.26E-09	0.26E-09
0.67E-08	0.48E-08								
0.39E-09	0.36E-09	0.56E-11	0.39E-11	0.34E-11	0.17E-11	0.15E-11	0.10E-11	0.88E-12	0.57E-12
0.28E-10	0.96E-11	0.33E-09	0.31E-09	0.26E-09	0.25E-09	0.21E-09	0.20E-09	0.16E-09	0.15E-09
0.37E-08	0.33E-08								
0.25E-09	0.29E-09	0.45E-11	0.32E-11	0.28E-11	0.13E-11	0.12E-11	0.79E-12	0.70E-12	0.44E-12
0.22E-10	0.58E-11	0.20E-09	0.17E-09	0.14E-09	0.12E-09	0.10E-09	0.82E-10	0.67E-10	0.49E-10
0.12E-08	0.32E-10								
0.82E-10	0.25E-09	0.39E-11	0.53E-11	0.47E-11	0.17E-11	0.15E-11	0.88E-12	0.78E-12	0.37E-12
0.18E-10	0.26E-11	0.91E-10	0.73E-10	0.62E-10	0.50E-10	0.43E-10	0.32E-10	0.26E-10	0.14E-10
0.34E-09	0.12E-11								

SIMULATION TWO-Vertical component

-0.68E-14	0.21E-14	-0.20E-11	-0.25E-10	-0.38E-11	-0.14E-10	0.97E-11	-0.48E-11	0.60E-11	-0.26E-11
0.31E-10	0.56E-11	0.27E-10	0.91E-11	0.24E-10	0.94E-11	0.40E-10	0.11E-10	0.27E-10	0.15E-10
0.20E-10	0.15E-10								
0.26E-13	0.46E-13	0.32E-10	-0.19E-10	0.19E-10	-0.74E-11	0.10E-10	-0.62E-11	0.10E-10	0.33E-11
0.38E-11	0.49E-11	0.39E-11	0.67E-11	0.61E-11	0.89E-11	0.93E-11	0.54E-11	0.13E-10	0.83E-11
0.13E-10	0.90E-11								
0.82E-13	0.14E-12	0.20E-10	-0.43E-10	0.12E-10	-0.32E-10	0.12E-10	-0.16E-10	0.42E-11	-0.82E-11
0.32E-11	-0.13E-11	0.19E-11	0.49E-11	-0.81E-13	0.27E-11	0.22E-11	-0.20E-12	0.58E-11	0.19E-11
0.67E-11	0.35E-11								
0.19E-12	0.63E-12	0.23E-10	0.16E-11	0.16E-10	-0.22E-10	0.73E-11	-0.14E-10	0.32E-11	-0.13E-10
-0.46E-12	-0.50E-11	-0.33E-11	-0.63E-11	-0.20E-11	-0.28E-11	-0.19E-11	-0.53E-11	0.12E-12	-0.41E-11
0.12E-11	-0.44E-11								
0.74E-12	0.16E-11	0.42E-13	0.80E-12	0.24E-10	-0.70E-10	0.10E-10	-0.51E-10	0.85E-11	-0.35E-10
-0.25E-11	-0.28E-10	-0.16E-11	-0.18E-10	-0.59E-11	-0.83E-11	-0.71E-11	-0.97E-11	-0.66E-11	-0.90E-11
-0.57E-11	-0.88E-11								
0.19E-11	0.49E-11	0.93E-12	0.27E-11	0.54E-10	0.56E-11	0.35E-10	-0.87E-10	0.20E-10	-0.75E-10
0.13E-10	-0.51E-10	0.52E-11	-0.30E-10	-0.32E-11	-0.12E-10	-0.88E-11	-0.12E-10	-0.89E-11	-0.12E-10
-0.10E-10	-0.11E-10								
0.57E-11	0.14E-10	0.32E-11	0.66E-11	0.12E-12	0.27E-11	0.76E-10	-0.11E-09	0.64E-10	-0.72E-10
0.42E-10	-0.49E-10	0.24E-10	-0.26E-10	0.71E-11	-0.13E-10	-0.11E-10	-0.22E-10	-0.12E-10	-0.25E-10
-0.13E-10	-0.23E-10								
0.98E-11	0.18E-10	0.74E-11	0.14E-10	0.31E-11	0.71E-11	0.11E-09	0.67E-10	0.74E-10	-0.40E-10
0.51E-10	-0.23E-10	0.28E-10	-0.88E-11	0.13E-10	0.23E-11	-0.21E-10	-0.47E-10	-0.25E-10	-0.55E-10
-0.25E-10	-0.22E-10								
0.20E-10	0.26E-10	0.93E-11	0.16E-10	0.76E-11	0.15E-10	0.92E-13	0.15E-11	0.53E-10	-0.13E-10
0.34E-10	-0.92E-12	0.17E-10	-0.42E-12	0.21E-11	-0.31E-11	-0.52E-10	-0.11E-09	-0.57E-10	-0.24E-09
-0.32E-10	-0.31E-09								
0.95E-11	0.10E-10	0.55E-11	0.85E-11	0.15E-10	0.16E-10	0.16E-11	0.11E-10	0.31E-10	0.28E-10
0.40E-11	-0.56E-10	-0.76E-11	-0.34E-10	-0.15E-10	-0.17E-10	-0.12E-09	-0.15E-09	-0.25E-09	-0.43E-09
-0.32E-09	-0.76E-09								
0.12E-10	0.12E-10	0.93E-11	0.94E-11	0.34E-11	0.92E-11	0.72E-11	0.12E-10	0.14E-12	0.99E-11
0.57E-10	0.49E-10	0.17E-10	-0.11E-09	-0.11E-10	-0.20E-10	-0.15E-09	0.14E-10	-0.44E-09	0.77E-10
-0.77E-09	0.44E-10								
0.13E-10	0.13E-10	0.11E-10	0.11E-10	0.11E-10	0.11E-10	0.45E-11	0.11E-10	0.70E-11	0.14E-10
0.18E-12	0.16E-11	0.10E-09	-0.75E-10	0.51E-10	0.30E-10	0.25E-10	0.22E-09	0.86E-10	0.89E-09
0.41E-10	0.16E-08								
0.14E-10	0.13E-10	0.12E-10	0.93E-11	0.13E-10	0.12E-10	0.13E-10	0.13E-10	0.52E-11	0.14E-10
0.12E-11	0.99E-11	0.21E-09	-0.28E-10	0.84E-10	0.14E-10	0.23E-09	0.34E-10	0.89E-09	0.21E-10
0.16E-08	-0.33E-09								

0.42E-10	0.40E-10	0.10E-10	0.52E-11	0.13E-10	0.12E-10	0.13E-10	0.14E-10	0.15E-10	0.15E-10
0.36E-11	0.15E-10	0.95E-10	0.11E-10	0.81E-11	0.88E-11	0.25E-10	-0.50E-10	0.16E-10	-0.24E-09
-0.34E-09	-0.54E-09								
0.41E-10	0.40E-10	0.21E-10	0.15E-10	0.12E-10	0.82E-11	0.15E-10	0.17E-10	0.15E-10	0.18E-10
0.16E-10	0.16E-10	0.23E-10	-0.17E-11	-0.69E-11	-0.52E-11	-0.59E-10	-0.81E-10	-0.24E-09	-0.33E-09
-0.54E-09	-0.70E-09								
0.68E-10	0.70E-10	0.28E-10	0.28E-10	0.26E-10	0.19E-10	0.17E-10	0.12E-10	0.18E-10	0.24E-10
0.16E-10	0.18E-10	0.40E-11	-0.53E-10	0.27E-11	-0.34E-10	-0.83E-10	-0.11E-09	-0.34E-09	-0.41E-09
-0.71E-09	-0.78E-09								
0.72E-10	0.72E-10	0.29E-10	0.29E-10	0.34E-10	0.27E-10	0.39E-10	0.27E-10	0.25E-10	0.25E-10
0.19E-10	0.21E-10	0.60E-10	0.11E-10	0.52E-10	0.19E-10	-0.11E-09	-0.95E-10	-0.42E-09	-0.27E-09
-0.79E-09	-0.47E-09								
0.72E-10	0.72E-10	0.30E-10	0.34E-10	0.28E-10	0.33E-10	0.29E-10	0.27E-10	0.26E-10	0.19E-10
0.21E-10	0.21E-10	0.87E-10	0.69E-10	0.72E-10	0.19E-09	-0.10E-09	-0.75E-10	-0.28E-09	-0.17E-09
-0.47E-09	-0.56E-10								
0.43E-10	0.43E-10	0.35E-10	0.39E-10	0.33E-10	0.38E-10	0.27E-10	0.28E-10	0.19E-10	0.15E-10
0.21E-10	0.20E-10	0.50E-10	0.19E-09	0.90E-13	0.36E-11	-0.81E-10	-0.59E-10	-0.17E-09	-0.15E-09
-0.59E-10	-0.12E-09								
0.14E-10	0.15E-10	0.23E-10	0.27E-10	0.23E-10	0.27E-10	0.28E-10	0.28E-10	0.44E-10	0.29E-10
0.20E-10	0.18E-10	0.61E-13	0.10E-10	0.37E-11	0.58E-11	-0.69E-10	-0.36E-10	-0.15E-09	-0.12E-09
-0.13E-09	-0.24E-09								
0.14E-10	0.16E-10	0.26E-10	0.33E-10	0.27E-10	0.32E-10	0.28E-10	0.31E-10	0.29E-10	0.21E-10
0.17E-10	0.11E-10	0.22E-11	0.44E-11	0.58E-11	0.14E-10	-0.47E-10	0.41E-10	-0.13E-09	-0.70E-10
-0.24E-09	-0.28E-09								
0.16E-10	0.14E-10	0.11E-10	0.14E-10	0.11E-10	0.15E-10	0.10E-10	0.59E-11	0.21E-10	0.14E-10
0.34E-10	0.22E-10	0.14E-10	0.14E-10	0.27E-11	0.38E-11	0.68E-10	0.21E-10	-0.83E-10	-0.69E-10
-0.29E-09	0.24E-10								
0.14E-10	0.13E-10	0.14E-10	0.14E-10	0.15E-10	0.14E-10	0.30E-10	0.29E-10	0.24E-10	0.21E-10
0.37E-10	0.28E-10	0.23E-10	0.23E-10	0.39E-11	0.38E-11	-0.78E-12	0.23E-11	-0.77E-10	-0.44E-10
0.22E-10	-0.34E-10								
0.13E-10	0.10E-10	0.14E-10	0.10E-10	0.14E-10	0.10E-10	0.58E-11	0.95E-11	0.12E-10	0.10E-10
0.17E-10	0.15E-10	0.13E-10	0.13E-10	0.38E-11	0.36E-11	0.69E-11	0.73E-11	-0.53E-10	0.35E-10
-0.42E-10	0.11E-09								
0.10E-10	0.94E-11	0.10E-10	0.85E-11	0.10E-10	0.84E-11	0.95E-11	0.81E-11	0.99E-11	0.11E-10
0.50E-11	0.42E-11	0.43E-11	0.39E-11	0.36E-11	0.34E-11	-0.15E-10	0.17E-10	0.32E-10	0.12E-09
0.10E-09	0.38E-09								
0.93E-11	0.90E-11	0.86E-11	0.80E-11	0.84E-11	0.77E-11	0.81E-11	0.72E-11	0.36E-11	0.56E-11
0.49E-11	0.54E-11	0.39E-11	0.45E-11	0.34E-11	0.12E-11	0.83E-11	-0.58E-10	0.12E-09	0.17E-09
0.38E-09	0.19E-09								
0.89E-11	0.89E-11	0.79E-11	0.78E-11	0.77E-11	0.74E-11	0.72E-11	0.72E-11	0.57E-11	0.63E-11
0.55E-11	0.59E-11	0.46E-11	0.49E-11	0.12E-11	0.15E-12	-0.67E-10	0.48E-11	0.33E-09	0.10E-09
0.17E-09	0.24E-09								

0.38E-11	0.84E-11	0.77E-11	0.74E-11	0.74E-11	0.72E-11	0.72E-11	0.70E-11	0.64E-11	0.66E-11
0.59E-11	0.41E-11	0.49E-11	0.31E-11	0.75E-12	0.53E-12	-0.98E-11	0.33E-10	0.77E-10	0.60E-10
0.23E-09	0.26E-09								
0.83E-11	0.79E-11	0.74E-11	0.75E-11	0.72E-11	0.74E-11	0.70E-11	0.73E-11	0.66E-11	0.79E-11
0.61E-11	0.22E-11	0.31E-11	0.32E-11	0.54E-12	0.28E-13	0.31E-10	0.25E-10	0.40E-10	0.35E-10
0.25E-09	0.92E-10								
0.78E-11	0.77E-11	0.74E-11	0.75E-11	0.73E-11	0.72E-11	0.73E-11	0.67E-11	0.79E-11	0.60E-11
0.66E-11	0.68E-11	0.99E-11	0.10E-10	0.63E-10	0.23E-10	0.14E-10	0.24E-10	0.15E-10	0.23E-10
0.91E-10	-0.23E-10								
0.76E-11	0.75E-11	0.76E-11	0.82E-11	0.72E-11	0.78E-11	0.67E-11	0.72E-11	0.60E-11	0.33E-11
0.68E-11	0.89E-11	0.17E-10	0.11E-10	0.22E-11	0.22E-10	0.42E-12	0.13E-10	-0.26E-11	0.74E-11
-0.39E-10	-0.12E-09								
0.74E-11	0.70E-11	0.82E-11	0.98E-11	0.78E-11	0.10E-10	0.72E-11	0.47E-11	0.98E-11	0.77E-11
0.15E-10	0.10E-10	0.11E-10	-0.12E-13	-0.46E-11	-0.69E-11	-0.10E-10	0.55E-12	-0.17E-10	0.11E-10
-0.13E-09	0.13E-09								
0.70E-11	0.67E-11	0.98E-11	0.95E-11	0.10E-10	0.95E-11	0.14E-10	0.95E-11	0.13E-10	0.93E-11
0.11E-10	0.94E-11	0.10E-10	-0.45E-10	0.16E-11	-0.19E-10	0.57E-11	-0.90E-11	0.98E-11	0.18E-10
0.13E-09	0.84E-10								
0.68E-11	0.67E-11	0.98E-11	0.85E-11	0.98E-11	0.85E-11	0.98E-11	0.84E-11	0.98E-11	0.83E-11
0.99E-11	0.63E-11	0.52E-10	-0.18E-10	0.41E-10	-0.71E-11	0.40E-10	0.26E-11	0.28E-10	0.15E-10
0.81E-10	0.72E-10								
0.68E-11	0.71E-11	0.87E-11	0.91E-11	0.86E-11	0.79E-11	0.86E-11	0.78E-11	0.85E-11	0.77E-11
0.96E-11	0.76E-11	0.52E-10	-0.10E-10	0.44E-10	-0.24E-11	0.37E-10	0.44E-11	0.26E-10	0.13E-10
0.67E-10	0.43E-10								
0.71E-11	0.82E-11	0.81E-11	0.89E-11	0.79E-11	0.62E-11	0.78E-11	0.63E-11	0.78E-11	0.63E-11
0.78E-11	0.68E-11	0.28E-10	0.97E-11	0.26E-10	0.14E-10	0.23E-10	0.14E-10	0.18E-10	0.16E-10
0.32E-10	0.31E-10								

SIMULATION THREE-Horizontal component

0.10E-12	0.36E-12	0.40E-10	0.28E-10	0.47E-10	0.43E-10	0.59E-10	0.48E-10	0.58E-10	0.52E-10
0.63E-10	0.46E-10	0.50E-10	0.31E-10	0.34E-10	0.24E-10	0.96E-09	0.41E-09	0.38E-09	0.14E-09
0.13E-09	0.00E+00								
-0.41E-12	0.16E-12	0.26E-10	-0.85E-11	0.23E-10	0.56E-11	0.24E-10	0.87E-11	0.23E-10	0.15E-10
0.19E-10	0.14E-10	0.17E-10	0.12E-10	0.15E-10	0.12E-10	0.48E-09	0.31E-09	0.28E-09	0.14E-09
0.13E-09	0.00E+00								
-0.16E-11	0.70E-12	-0.23E-11	-0.53E-10	-0.25E-11	-0.35E-10	0.13E-10	-0.15E-10	0.13E-10	-0.21E-13
0.18E-10	0.79E-11	0.16E-10	0.11E-10	0.13E-10	0.90E-11	0.35E-09	0.19E-09	0.25E-09	0.95E-10
0.15E-09	0.12E-10								

-0.54E-11	0.75E-11	-0.15E-10	-0.27E-10	0.13E-11	-0.14E-10	0.11E-10	-0.25E-11	0.12E-10	0.41E-11
0.14E-10	0.74E-11	0.12E-10	0.66E-11	0.11E-10	0.6PE-11	0.27E-09	0.16E-09	0.18E-09	0.94E-10
0.95E-10	0.00E+00								
-0.25E-10	0.46E-11	-0.15E-10	0.83E-11	0.62E-10	0.17E-10	0.57E-10	0.14E-10	0.45E-10	0.20E-10
0.35E-10	0.17E-10	0.25E-10	0.11E-10	0.15E-10	0.81E-11	0.29E-09	0.15E-09	0.14E-09	0.64E-10
0.96E-10	0.00E+00								
-0.62E-10	0.27E-10	-0.22E-10	0.33E-10	0.10E-09	0.81E-10	0.11E-09	0.44E-10	0.78E-10	0.33E-10
0.58E-10	0.19E-10	0.33E-10	0.11E-10	0.20E-10	0.10E-10	0.38E-09	0.19E-09	0.17E-09	0.11E-09
0.98E-10	0.42E-10								
-0.18E-09	0.80E-10	-0.74E-10	0.34E-10	-0.26E-10	0.26E-10	0.16E-09	0.68E-10	0.12E-09	0.60E-10
0.87E-10	0.37E-10	0.52E-10	0.28E-10	0.36E-10	0.21E-10	0.72E-09	0.28E-09	0.30E-09	0.12E-09
0.11E-09	-0.64E-10								
-0.39E-09	0.10E-09	-0.16E-09	0.93E-10	-0.14E-10	0.73E-10	0.12E-09	0.93E-10	0.12E-09	0.67E-10
0.95E-10	0.50E-10	0.72E-10	0.52E-10	0.62E-10	0.49E-10	0.15E-08	0.54E-09	0.55E-09	0.72E-10
0.15E-09	0.41E-10								
-0.44E-09	0.55E-10	-0.23E-09	0.13E-09	-0.39E-10	0.11E-09	-0.75E-09	0.68E-09	0.18E-09	0.15E-09
0.17E-09	0.14E-09	0.14E-09	0.13E-09	0.12E-09	0.11E-09	0.32E-08	0.12E-08	0.11E-08	-0.13E-08
-0.77E-09	-0.54E-08								
-0.42E-09	-0.19E-09	-0.71E-11	0.37E-11	-0.39E-10	-0.34E-10	-0.16E-09	0.91E-10	0.39E-09	0.38E-09
0.38E-09	0.33E-09	0.31E-09	0.29E-09	0.27E-09	0.26E-09	0.72E-08	0.61E-08	0.52E-08	0.25E-08
0.23E-08	-0.57E-08								
-0.32E-09	-0.32E-09	-0.48E-11	-0.45E-11	-0.83E-11	0.27E-11	-0.14E-09	0.51E-10	-0.11E-09	0.84E-10
0.68E-09	0.60E-09	0.58E-09	0.53E-09	0.47E-09	0.45E-09	0.12E-07	0.17E-07	0.13E-07	0.20E-07
0.17E-07	0.31E-07								
-0.31E-09	-0.47E-09	-0.64E-11	-0.70E-11	-0.47E-11	-0.36E-11	-0.83E-11	0.58E-11	-0.57E-10	0.15E-09
-0.20E-09	0.10E-08	0.42E-09	0.35E-09	0.32E-09	0.30E-09	0.84E-08	0.13E-07	0.11E-07	0.20E-07
0.17E-07	0.36E-07								
-0.30E-09	-0.70E-09	-0.88E-11	-0.24E-10	-0.16E-10	-0.18E-10	-0.15E-10	-0.16E-10	-0.19E-10	0.51E-11
-0.25E-09	0.16E-09	0.31E-10	-0.14E-09	-0.46E-10	-0.13E-09	-0.17E-08	-0.10E-07	-0.60E-08	-0.25E-07
-0.18E-07	-0.52E-07								
-0.26E-09	-0.78E-09	-0.13E-10	-0.42E-10	-0.31E-10	-0.34E-10	-0.29E-10	-0.27E-10	-0.69E-10	-0.79E-10
-0.28E-10	0.16E-10	-0.45E-12	-0.40E-11	-0.52E-10	-0.11E-09	-0.14E-08	-0.79E-08	-0.62E-08	-0.15E-07
-0.13E-07	-0.21E-07								
-0.35E-09	-0.77E-09	-0.16E-10	-0.45E-10	-0.31E-10	-0.58E-10	-0.46E-10	-0.25E-10	-0.73E-10	-0.13E-10
-0.15E-09	0.75E-12	0.62E-12	-0.25E-12	0.57E-10	-0.47E-10	-0.54E-09	-0.58E-08	-0.46E-08	-0.11E-07
-0.99E-08	-0.23E-07								
-0.65E-09	-0.58E-09	-0.11E-10	-0.13E-10	-0.78E-11	-0.15E-10	-0.58E-11	-0.26E-10	-0.56E-10	0.36E-11
-0.10E-09	0.22E-09	0.24E-11	0.61E-12	0.73E-10	-0.34E-11	0.57E-09	-0.24E-08	-0.26E-08	-0.48E-08
-0.42E-08	-0.65E-08								
0.15E-09	-0.41E-10	-0.32E-11	-0.42E-11	-0.20E-11	-0.62E-11	0.18E-11	-0.10E-10	0.12E-10	0.59E-11
0.42E-10	0.30E-09	0.42E-11	0.30E-11	0.38E-11	0.25E-11	0.29E-08	0.28E-08	0.13E-08	0.49E-08
0.42E-08	0.12E-07								

0.94E-09	0.48E-09	0.62E-11	0.11E-10	0.95E-11	0.12E-10	0.14E-10	0.10E-10	0.77E-10	-0.16E-10
0.74E-10	0.20E-11	0.29E-13	0.55E-14	0.11E-11	0.46E-11	0.46E-08	0.60E-08	0.23E-08	0.84E-08
0.69E-08	0.23E-07								
0.71E-09	0.67E-09	0.84E-11	0.18E-10	0.16E-10	0.23E-10	0.71E-10	0.79E-10	0.11E-09	0.16E-11
0.21E-10	-0.98E-10	-0.58E-09	-0.56E-09	-0.45E-09	0.53E-10	0.57E-08	0.67E-08	0.38E-08	0.65E-08
0.51E-08	0.00E+00								
0.92E-09	0.70E-09	0.12E-10	0.22E-10	0.19E-10	0.19E-10	0.57E-10	0.66E-10	0.59E-10	0.10E-10
0.14E-09	-0.57E-09	-0.67E-09	-0.11E-09	-0.11E-09	0.44E-10	0.54E-08	0.73E-08	0.47E-08	0.65E-08
0.51E-08	0.00E+00								
0.52E-09	0.40E-09	0.36E-09	0.13E-09	0.12E-09	0.74E-10	0.41E-11	0.36E-10	0.32E-10	0.30E-10
0.41E-09	-0.44E-09	-0.48E-09	-0.14E-09	-0.17E-09	0.12E-09	0.46E-08	0.89E-08	0.55E-08	0.81E-08
0.62E-08	0.53E-08								
0.85E-09	0.38E-09	0.35E-09	0.17E-09	0.16E-09	0.11E-09	0.60E-11	0.55E-10	0.49E-10	0.40E-10
0.57E-09	0.43E-11	-0.40E-09	-0.42E-09	-0.21E-10	0.16E-09	0.63E-09	0.56E-09	0.14E-07	0.14E-07
0.86E-08	0.18E-07								
0.65E-09	0.34E-09	0.33E-09	0.13E-09	0.13E-09	0.27E-09	0.16E-10	0.20E-10	0.19E-10	0.16E-10
0.25E-09	0.14E-10	0.17E-09	0.17E-09	0.34E-11	0.13E-10	0.40E-09	0.40E-09	0.11E-07	0.12E-07
0.10E-07	0.73E-08								
0.61E-09	0.61E-09	0.10E-10	0.19E-10	0.18E-10	0.23E-10	0.74E-10	0.21E-10	0.21E-10	0.42E-11
0.67E-10	0.11E-10	0.41E-09	0.84E-09	0.32E-12	0.19E-12	0.36E-09	0.35E-09	0.99E-08	0.13E-07
0.98E-08	0.16E-07								
0.53E-09	0.58E-09	0.96E-11	0.85E-11	0.79E-11	0.54E-11	0.17E-10	0.10E-12	0.98E-13	-0.13E-12
-0.21E-11	0.17E-10	0.47E-09	0.62E-09	0.53E-09	0.31E-09	0.26E-09	0.27E-09	0.77E-08	0.12E-07
0.90E-08	0.26E-07								
0.51E-09	0.57E-09	0.91E-11	0.75E-11	0.67E-11	0.43E-11	0.13E-10	0.29E-11	0.26E-11	0.31E-10
0.27E-10	0.33E-10	0.47E-09	0.63E-09	0.53E-09	0.80E-11	0.27E-09	0.19E-09	0.56E-08	0.91E-08
0.68E-08	-0.13E-07								
0.51E-09	0.57E-09	0.90E-11	0.80E-11	0.71E-11	0.49E-11	0.15E-10	0.14E-10	0.12E-10	0.31E-10
0.26E-10	0.35E-10	0.48E-09	0.66E-09	0.55E-09	0.37E-10	0.13E-08	0.14E-08	0.12E-08	0.56E-09
0.15E-07	0.26E-07								
0.55E-09	0.56E-09	0.91E-11	0.90E-11	0.82E-11	0.89E-11	0.82E-11	0.84E-11	0.26E-10	0.32E-10
0.28E-10	0.34E-10	0.49E-09	0.37E-10	0.13E-08	0.77E-09	0.67E-09	0.68E-09	0.56E-09	0.51E-09
0.13E-07	0.14E-07								
0.55E-09	0.56E-09	0.89E-11	0.97E-11	0.87E-11	0.96E-11	0.86E-11	0.95E-11	0.28E-10	0.37E-10
0.31E-10	0.11E-11	0.60E-09	0.13E-08	0.52E-12	0.24E-12	0.46E-09	0.45E-09	0.36E-09	0.34E-09
0.88E-08	0.00E+00								
0.53E-09	0.55E-09	0.87E-11	0.94E-11	0.84E-11	0.82E-11	0.73E-11	0.58E-11	0.51E-11	0.11E-12
0.20E-09	0.55E-09	0.21E-12	0.24E-12	0.47E-09	0.43E-09	0.37E-09	0.36E-09	0.29E-09	0.28E-09
0.74E-08	0.00E+00								
0.53E-09	0.54E-09	0.86E-11	0.12E-10	0.11E-10	0.14E-10	0.13E-10	0.15E-10	0.13E-10	0.36E-11
0.17E-09	0.35E-09	0.56E-11	0.21E-12	0.42E-09	0.43E-09	0.37E-09	0.38E-09	0.31E-09	0.31E-09
0.82E-08	0.00E+00								

0.54E-09	0.54E-09	0.84E-11	0.15E-10	0.13E-10	0.22E-10	0.20E-10	0.14E-10	0.13E-10	0.11E-10
0.96E-11	0.72E-11	0.37E-09	0.15E-10	0.53E-09	0.53E-09	0.46E-09	0.46E-09	0.38E-09	0.39E-09
0.10E-07	0.26E-07								
0.49E-09	0.46E-09	0.71E-11	0.66E-11	0.58E-11	0.47E-11	0.41E-11	0.21E-11	0.18E-11	0.77E-12
0.37E-10	0.15E-10	0.52E-09	0.49E-09	0.41E-09	0.40E-09	0.34E-09	0.32E-09	0.26E-09	0.26E-09
0.67E-08	0.48E-08								
0.36E-09	0.35E-09	0.55E-11	0.39E-11	0.33E-11	0.16E-11	0.14E-11	0.98E-12	0.86E-12	0.56E-12
0.27E-10	0.95E-11	0.33E-09	0.31E-09	0.26E-09	0.24E-09	0.21E-09	0.19E-09	0.16E-09	0.15E-09
0.37E-08	0.33E-08								
0.23E-09	0.28E-09	0.44E-11	0.31E-11	0.27E-11	0.13E-11	0.11E-11	0.78E-12	0.68E-12	0.43E-12
0.21E-10	0.57E-11	0.20E-09	0.17E-09	0.14E-09	0.12E-09	0.10E-09	0.82E-10	0.66E-10	0.49E-10
0.12E-08	0.32E-10								
0.77E-10	0.25E-09	0.39E-11	0.53E-11	0.46E-11	0.17E-11	0.15E-11	0.88E-12	0.78E-12	0.37E-12
0.18E-10	0.26E-11	0.90E-10	0.73E-10	0.62E-10	0.50E-10	0.42E-10	0.32E-10	0.26E-10	0.14E-10
0.34E-09	0.12E-11								

SIMULATION THREE-Vertical component

-0.72E-14	0.11E-14	-0.20E-11	-0.25E-10	-0.38E-11	-0.13E-10	0.97E-11	-0.47E-11	0.60E-11	-0.25E-11
0.31E-10	0.57E-11	0.27E-10	0.91E-11	0.24E-10	0.94E-11	0.40E-10	0.11E-10	-0.27E-10	0.15E-10
0.20E-10	0.14E-10								
0.25E-13	0.42E-13	0.32E-10	-0.18E-10	0.18E-10	-0.74E-11	0.98E-11	-0.62E-11	0.10E-10	0.32E-11
0.37E-11	0.48E-11	0.38E-11	0.66E-11	0.60E-11	0.88E-11	0.92E-11	0.53E-11	0.13E-10	0.82E-11
0.13E-10	0.90E-11								
0.78E-13	0.13E-12	0.20E-10	-0.44E-10	0.11E-10	-0.32E-10	0.12E-10	-0.16E-10	0.40E-11	-0.86E-11
0.31E-11	-0.14E-11	0.18E-11	0.45E-11	-0.21E-12	0.26E-11	0.21E-11	-0.28E-12	0.58E-11	0.19E-11
0.67E-11	0.34E-11								
0.18E-12	0.59E-12	0.23E-10	0.15E-11	0.16E-10	-0.23E-10	0.74E-11	-0.15E-10	0.32E-11	-0.14E-10
-0.43E-12	-0.56E-11	-0.33E-11	-0.67E-11	-0.21E-11	-0.31E-11	-0.19E-11	-0.53E-11	0.95E-13	-0.41E-11
0.11E-11	-0.43E-11								
0.70E-12	0.15E-11	0.44E-13	0.75E-12	0.25E-10	-0.71E-10	0.12E-10	-0.52E-10	0.94E-11	-0.36E-10
-0.18E-11	-0.28E-10	-0.11E-11	-0.18E-10	-0.56E-11	-0.81E-11	-0.70E-11	-0.94E-11	-0.66E-11	-0.89E-11
-0.56E-11	-0.87E-11								
0.18E-11	0.46E-11	0.89E-12	0.25E-11	0.56E-10	0.53E-11	0.37E-10	-0.84E-10	0.22E-10	-0.72E-10
0.14E-10	-0.48E-10	0.59E-11	-0.28E-10	-0.26E-11	-0.10E-10	-0.85E-11	-0.12E-10	-0.87E-11	-0.12E-10
-0.10E-10	-0.11E-10								
0.54E-11	0.13E-10	0.30E-11	0.62E-11	0.12E-12	0.25E-11	0.73E-10	-0.10E-09	0.62E-10	-0.67E-10
0.41E-10	-0.45E-10	0.23E-10	-0.24E-10	0.66E-11	-0.12E-10	-0.11E-10	-0.22E-10	-0.12E-10	-0.25E-10
-0.13E-10	-0.23E-10								
0.92E-11	0.17E-10	0.69E-11	0.14E-10	0.29E-11	0.66E-11	0.10E-09	0.62E-10	0.68E-10	-0.41E-10
0.46E-10	-0.24E-10	0.24E-10	-0.11E-10	0.10E-10	0.52E-12	-0.22E-10	-0.48E-10	-0.26E-10	-0.55E-10

0.19E-10	0.28E-10	0.60E-11	0.15E-10	0.72E-11	0.14E-10	0.89E-13	0.14E-11	0.52E-10	-0.19E-10
0.32E-10	-0.59E-11	0.16E-10	-0.38E-11	0.93E-12	-0.45E-11	-0.53E-10	-0.11E-09	-0.57E-10	-0.24E-09
-0.33E-10	-0.31E-09								
0.90E-11	0.97E-11	0.52E-11	0.80E-11	0.14E-10	0.15E-10	0.16E-11	0.10E-10	0.36E-10	0.27E-10
0.79E-11	-0.56E-10	-0.52E-11	-0.34E-10	-0.15E-10	-0.17E-10	-0.12E-09	-0.15E-09	-0.25E-09	-0.43E-09
-0.32E-09	-0.76E-09								
0.11E-10	0.11E-10	0.88E-11	0.88E-11	0.33E-11	0.87E-11	0.68E-11	0.11E-10	0.14E-12	0.93E-11
0.56E-10	0.47E-10	0.18E-10	-0.99E-10	-0.11E-10	-0.17E-10	-0.15E-09	0.15E-10	-0.44E-09	0.77E-10
-0.77E-09	0.44E-10								
0.12E-10	0.12E-10	0.10E-10	0.10E-10	0.10E-10	0.11E-10	0.42E-11	0.11E-10	0.66E-11	0.13E-10
0.17E-12	0.15E-11	0.98E-10	-0.62E-10	0.98E-10	0.35E-10	0.24E-10	0.22E-09	0.86E-10	0.89E-09
0.41E-10	0.16E-08								
0.13E-10	0.12E-10	0.11E-10	0.86E-11	0.12E-10	0.11E-10	0.12E-10	0.12E-10	0.48E-11	0.13E-10
0.11E-11	0.93E-11	0.19E-09	-0.22E-10	0.74E-10	0.13E-10	0.23E-09	0.31E-10	0.88E-09	0.19E-10
0.16E-08	-0.33E-09								
0.39E-10	0.37E-10	0.93E-11	0.63E-11	0.12E-10	0.11E-10	0.12E-10	0.13E-10	0.14E-10	0.14E-10
0.34E-11	0.14E-10	0.79E-10	0.91E-11	0.42E-11	0.79E-11	0.24E-10	-0.50E-10	0.14E-10	-0.24E-09
-0.34E-09	-0.54E-09								
0.38E-10	0.37E-10	0.20E-10	0.15E-10	0.11E-10	0.75E-11	0.13E-10	0.16E-10	0.14E-10	0.16E-10
0.15E-10	0.15E-10	0.24E-10	0.69E-11	-0.48E-11	0.13E-11	-0.59E-10	-0.80E-10	-0.24E-09	-0.33E-09
-0.54E-09	-0.70E-09								
0.62E-10	0.64E-10	0.25E-10	0.25E-10	0.23E-10	0.17E-10	0.16E-10	0.11E-10	0.17E-10	0.22E-10
0.15E-10	0.17E-10	0.11E-12	-0.33E-10	-0.17E-12	-0.23E-10	-0.84E-10	-0.11E-09	-0.34E-09	-0.41E-09
-0.71E-09	-0.78E-09								
0.65E-10	0.63E-10	0.26E-10	0.25E-10	0.30E-10	0.23E-10	0.35E-10	0.24E-10	0.23E-10	0.22E-10
0.18E-10	0.20E-10	0.43E-10	0.12E-10	0.42E-10	0.18E-10	-0.11E-09	-0.99E-10	-0.42E-09	-0.27E-09
-0.78E-09	-0.47E-09								
0.63E-10	0.59E-10	0.26E-10	0.29E-10	0.25E-10	0.27E-10	0.25E-10	0.23E-10	0.23E-10	0.17E-10
0.20E-10	0.20E-10	0.79E-10	0.70E-10	0.67E-10	0.21E-09	-0.11E-09	-0.80E-10	-0.28E-09	-0.17E-09
-0.47E-09	-0.59E-10								
0.35E-10	0.35E-10	0.29E-10	0.33E-10	0.28E-10	0.32E-10	0.23E-10	0.24E-10	0.17E-10	0.13E-10
0.20E-10	0.20E-10	0.62E-10	0.20E-09	0.99E-13	0.39E-11	-0.86E-10	-0.66E-10	-0.18E-09	-0.15E-09
-0.61E-10	-0.13E-09								
0.12E-10	0.11E-10	0.20E-10	0.24E-10	0.19E-10	0.18E-10	0.24E-10	0.26E-10	0.39E-10	0.28E-10
0.19E-10	0.17E-10	0.89E-13	0.11E-10	0.39E-11	0.63E-11	-0.75E-10	-0.43E-10	-0.16E-09	-0.12E-09
-0.13E-09	-0.24E-09								
0.33E-10	0.31E-10	0.40E-10	0.32E-10	0.31E-10	0.29E-10	0.27E-10	0.42E-10	0.28E-10	0.27E-10
0.18E-10	0.12E-10	0.24E-11	0.50E-11	0.66E-11	0.15E-10	-0.44E-10	0.40E-10	-0.13E-09	-0.69E-10
-0.25E-09	-0.28E-09								
0.31E-10	0.25E-10	0.32E-10	0.26E-10	0.29E-10	0.27E-10	0.14E-10	0.22E-10	0.27E-10	0.23E-10
0.38E-10	0.25E-10	0.16E-10	0.16E-10	0.31E-11	0.44E-11	0.85E-10	0.20E-10	-0.82E-10	-0.68E-10
-0.29E-09	0.27E-10								

0.24E-10	0.21E-10	0.21E-10	0.19E-10	0.27E-10	0.30E-10	0.22E-10	0.21E-10	0.23E-10	0.21E-10
0.43E-10	0.33E-10	0.27E-10	0.25E-10	0.45E-11	0.43E-11	0.45E-11	0.45E-11	-0.76E-10	-0.42E-10
0.25E-10	-0.31E-10								
0.69E-11	0.68E-11	0.38E-11	0.80E-11	0.60E-11	0.85E-11	0.21E-10	0.10E-10	0.21E-10	0.11E-10
0.20E-10	0.16E-10	0.15E-10	0.14E-10	0.43E-11	0.40E-11	0.63E-11	0.77E-11	-0.52E-10	0.37E-10
-0.40E-10	0.12E-09								
0.68E-11	0.67E-11	0.80E-11	0.74E-11	0.85E-11	0.74E-11	0.10E-10	0.75E-11	0.11E-10	0.99E-11
0.55E-11	0.51E-11	0.47E-11	0.43E-11	0.40E-11	0.37E-11	-0.15E-10	0.18E-10	0.32E-10	0.12E-09
0.11E-09	0.39E-09								
0.67E-11	0.64E-11	0.74E-11	0.71E-11	0.74E-11	0.69E-11	0.75E-11	0.67E-11	0.33E-11	0.54E-11
0.51E-11	0.54E-11	0.43E-11	0.49E-11	0.37E-11	0.13E-11	0.68E-11	-0.58E-10	0.12E-09	0.17E-09
0.38E-09	0.19E-09								
0.67E-11	0.68E-11	0.70E-11	0.70E-11	0.69E-11	0.67E-11	0.67E-11	0.67E-11	0.54E-11	0.61E-11
0.55E-11	0.58E-11	0.49E-11	0.53E-11	0.14E-11	0.18E-12	-0.68E-10	0.51E-11	0.33E-09	0.10E-09
0.18E-09	0.24E-09								
0.68E-11	0.67E-11	0.70E-11	0.68E-11	0.67E-11	0.67E-11	0.67E-11	0.66E-11	0.61E-11	0.64E-11
0.59E-11	0.61E-11	0.52E-11	0.32E-11	0.91E-12	0.62E-12	-0.99E-11	0.33E-10	0.77E-10	0.60E-10
0.23E-09	0.26E-09								
0.67E-11	0.67E-11	0.68E-11	0.70E-11	0.66E-11	0.69E-11	0.65E-11	0.69E-11	0.64E-11	0.76E-11
0.60E-11	0.21E-11	0.32E-11	0.34E-11	0.64E-12	0.31E-13	0.31E-10	0.28E-10	0.41E-10	0.37E-10
0.25E-09	0.94E-10								
0.66E-11	0.67E-11	0.70E-11	0.72E-11	0.69E-11	0.68E-11	0.69E-11	0.63E-11	0.76E-11	0.56E-11
0.62E-11	0.63E-11	0.10E-10	0.10E-10	0.70E-10	0.23E-10	0.17E-10	0.24E-10	0.17E-10	0.24E-10
0.83E-10	-0.27E-10								
0.67E-11	0.67E-11	0.71E-11	0.77E-11	0.68E-11	0.74E-11	0.63E-11	0.67E-11	0.56E-11	0.31E-11
0.63E-11	0.87E-11	0.17E-10	0.11E-10	0.28E-11	0.22E-10	0.11E-11	0.13E-10	-0.18E-11	0.76E-11
-0.38E-10	-0.12E-09								
0.66E-11	0.64E-11	0.77E-11	0.93E-11	0.73E-11	0.96E-11	0.67E-11	0.44E-11	0.92E-11	0.73E-11
0.14E-10	0.10E-10	0.11E-10	-0.11E-13	-0.44E-11	-0.65E-11	-0.98E-11	0.96E-12	-0.17E-10	0.11E-10
-0.13E-09	0.13E-09								
0.64E-11	0.62E-11	0.93E-11	0.91E-11	0.96E-11	0.91E-11	0.13E-10	0.90E-11	0.12E-10	0.89E-11
0.10E-10	0.89E-11	0.10E-10	-0.43E-10	0.11E-11	-0.17E-10	0.54E-11	-0.81E-11	0.97E-11	0.18E-10
0.13E-09	0.84E-10								
0.64E-11	0.63E-11	0.93E-11	0.91E-11	0.93E-11	0.80E-11	0.93E-11	0.79E-11	0.93E-11	0.78E-11
0.94E-11	0.78E-11	0.50E-10	-0.17E-10	0.39E-10	-0.65E-11	0.39E-10	0.29E-11	0.27E-10	0.15E-10
0.81E-10	0.72E-10								
0.64E-11	0.68E-11	0.83E-11	0.75E-11	0.82E-11	0.75E-11	0.81E-11	0.74E-11	0.81E-11	0.72E-11
0.81E-11	0.72E-11	0.49E-10	-0.10E-10	0.42E-10	-0.26E-11	0.36E-10	0.42E-11	0.25E-10	0.13E-10
0.67E-10	0.43E-10								
0.67E-11	0.78E-11	0.77E-11	0.85E-11	0.75E-11	0.58E-11	0.74E-11	0.58E-11	0.73E-11	0.58E-11
0.74E-11	0.63E-11	0.27E-10	0.92E-11	0.25E-10	0.13E-10	0.22E-10	0.13E-10	0.17E-10	0.15E-10
0.31E-10	0.30E-10								

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